Title: What drives weather changes?

Date: Apr 27, 2016 02:00 PM

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Abstract: <p>Winds are driven by the gradients of solar heating. Vertical gradients cause thermal convection on the scale of the troposphere depth (less than 10 km). Horizontal gradients excite motions on a planetary (10000 km) and smaller scales. Weather is mostly determined by the flows at intermediate scale (hundreds of kilometers). Where these flows get their energy from? The puzzle is that three-dimensional small-scale motions cannot transfer energy to larger scales while large-scale planar motions cannot transfer energy to smaller scales. In the talk, I'll describe experimental and observational data that suggest one possible resolution of this puzzle. I also describe some puzzling properties of two-dimensional turbulence including conformal invariance of statistics.</p>

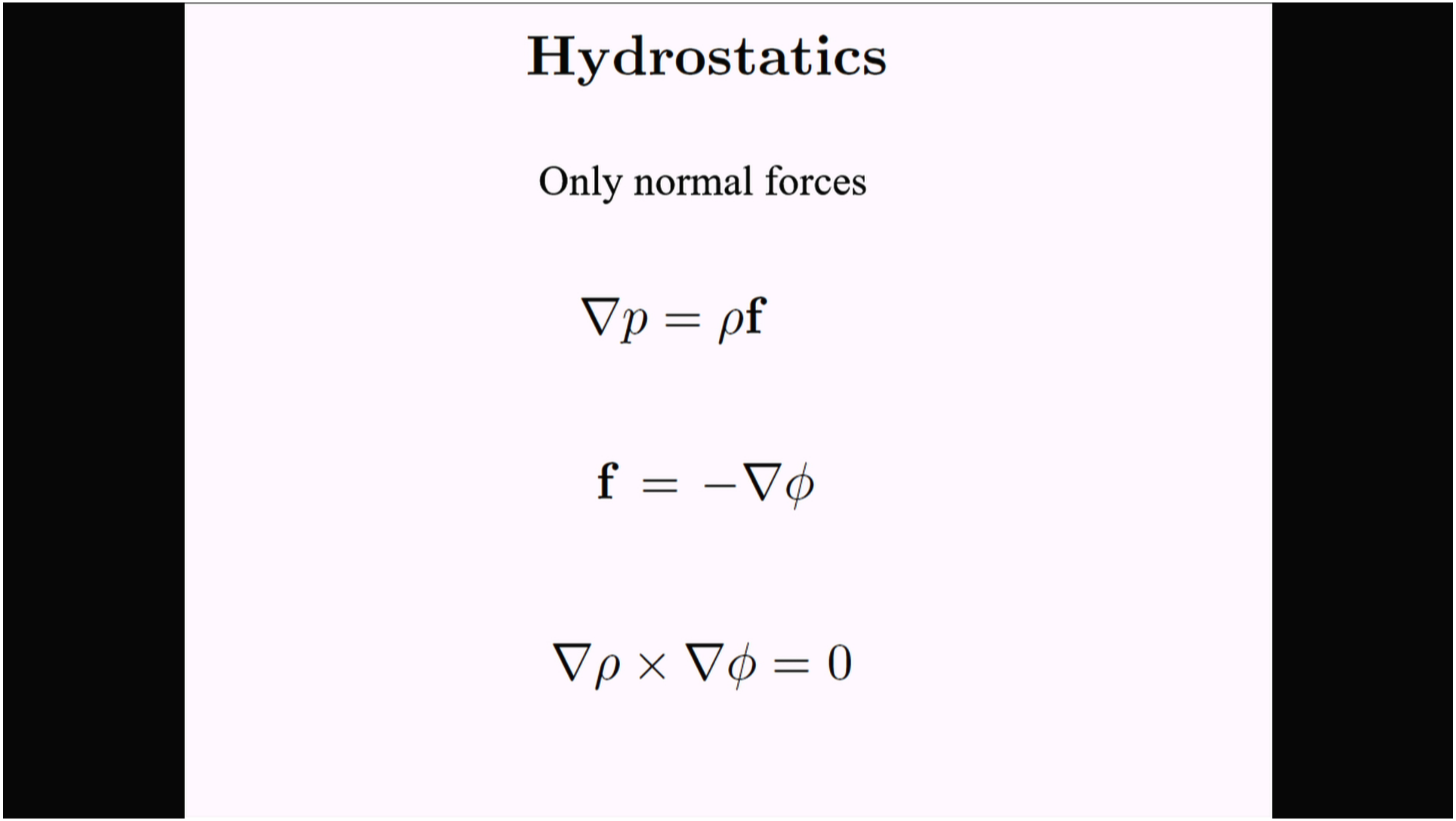
What drives weather changes

Gregory Falkovich

Dept of Physics, Weizmann Institute of Science

April 27, 2016, Perimeter Institute

“The answer is blowing in the wind”



Hydrostatics

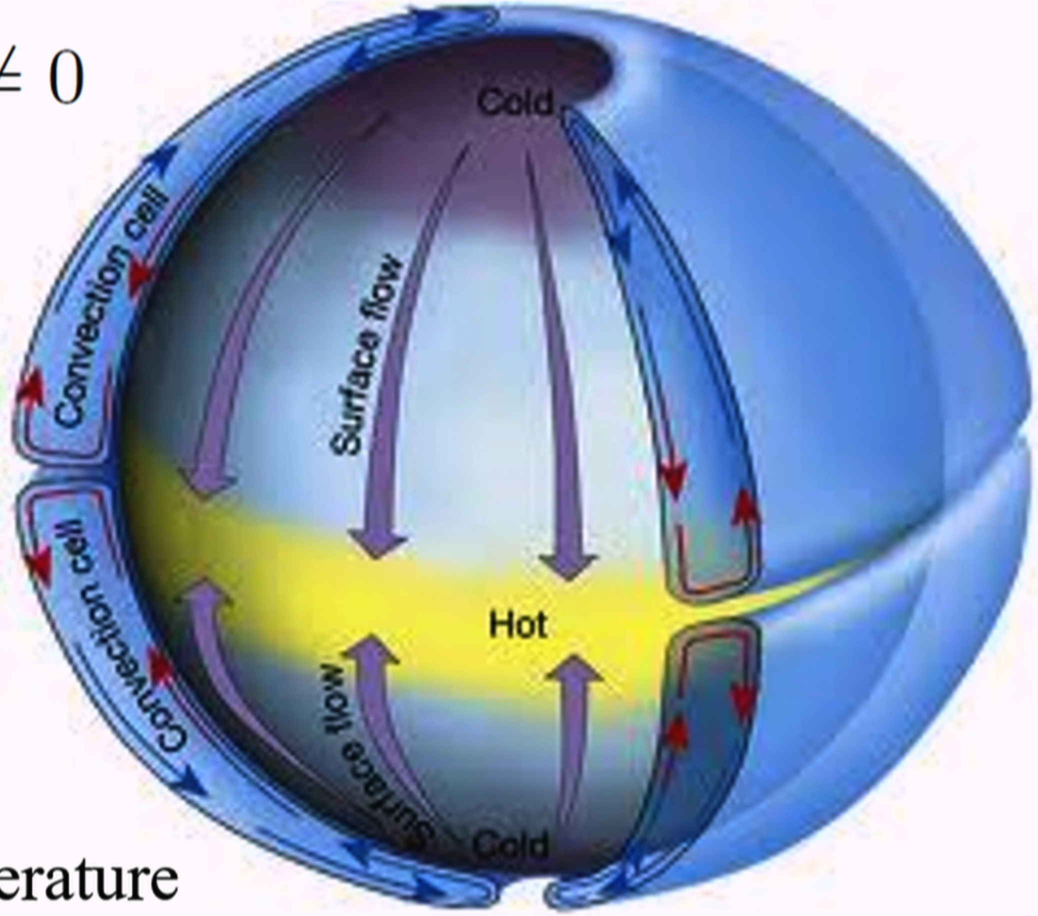
Only normal forces

$$\nabla p = \rho \mathbf{f}$$

$$\mathbf{f} \; = \; -\nabla\phi$$

$$\nabla\rho \times \nabla\phi = 0$$

$$\nabla\rho \times \nabla\phi \neq 0$$



Horizontal temperature gradient causes wind

What about the vertical temperature gradient?

$$T(z) = T_0 - \alpha z$$

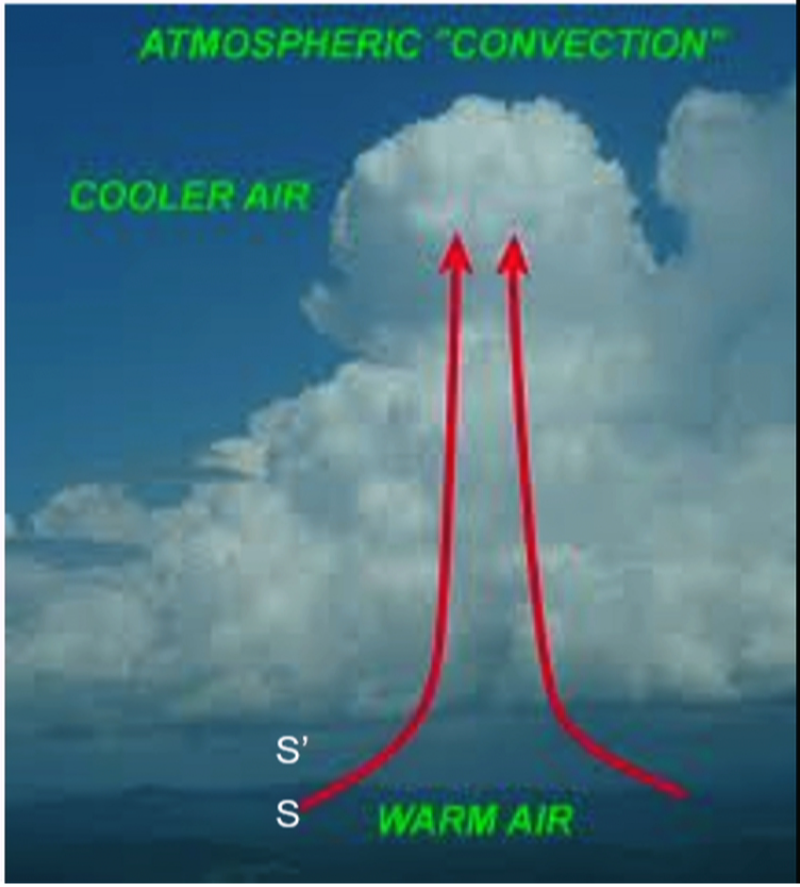
$$p(z) = p(0)(1 - \alpha z/T_0)^{mg/\alpha}$$

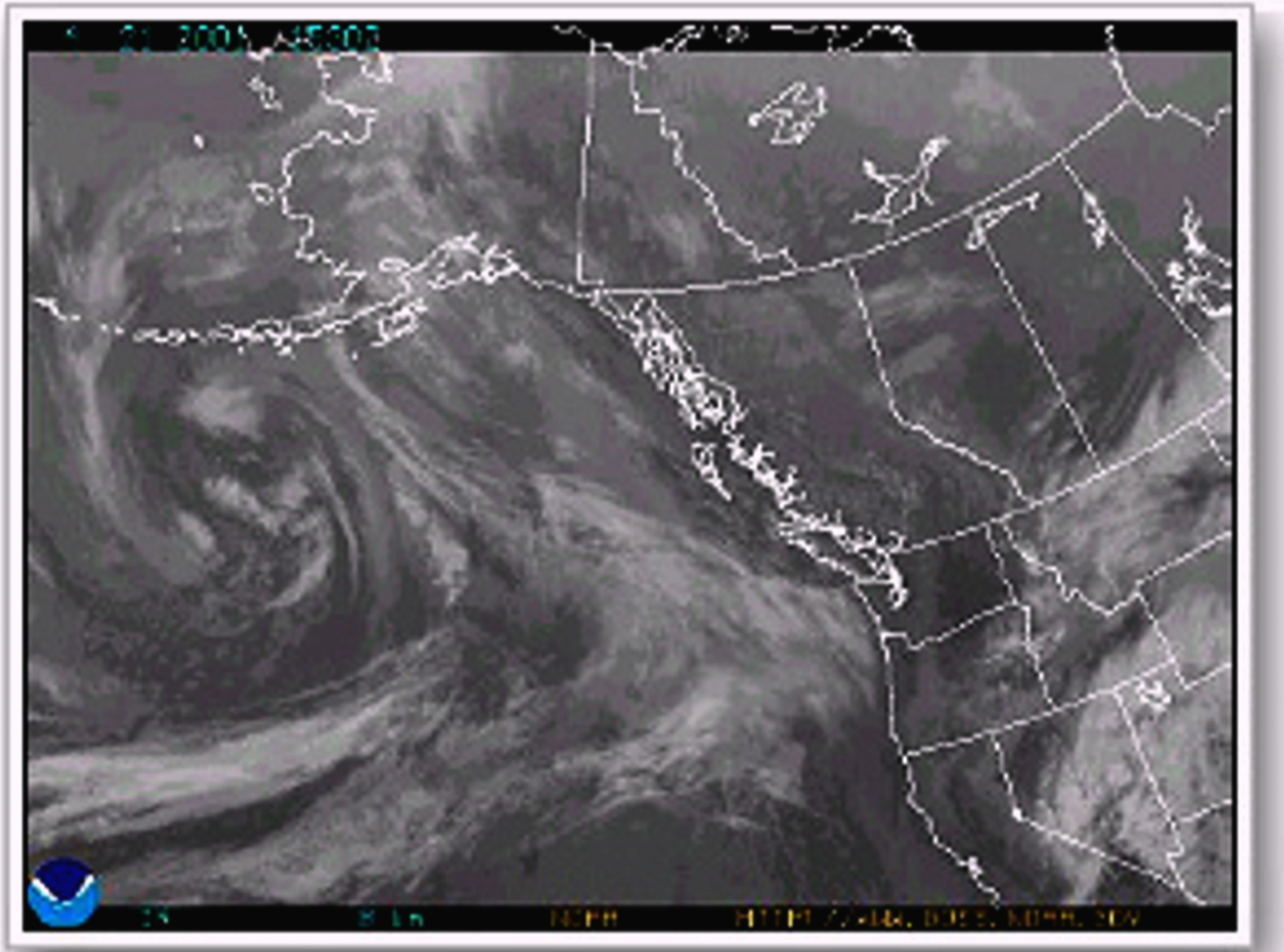
$$\alpha \simeq 6.5^\circ/km$$

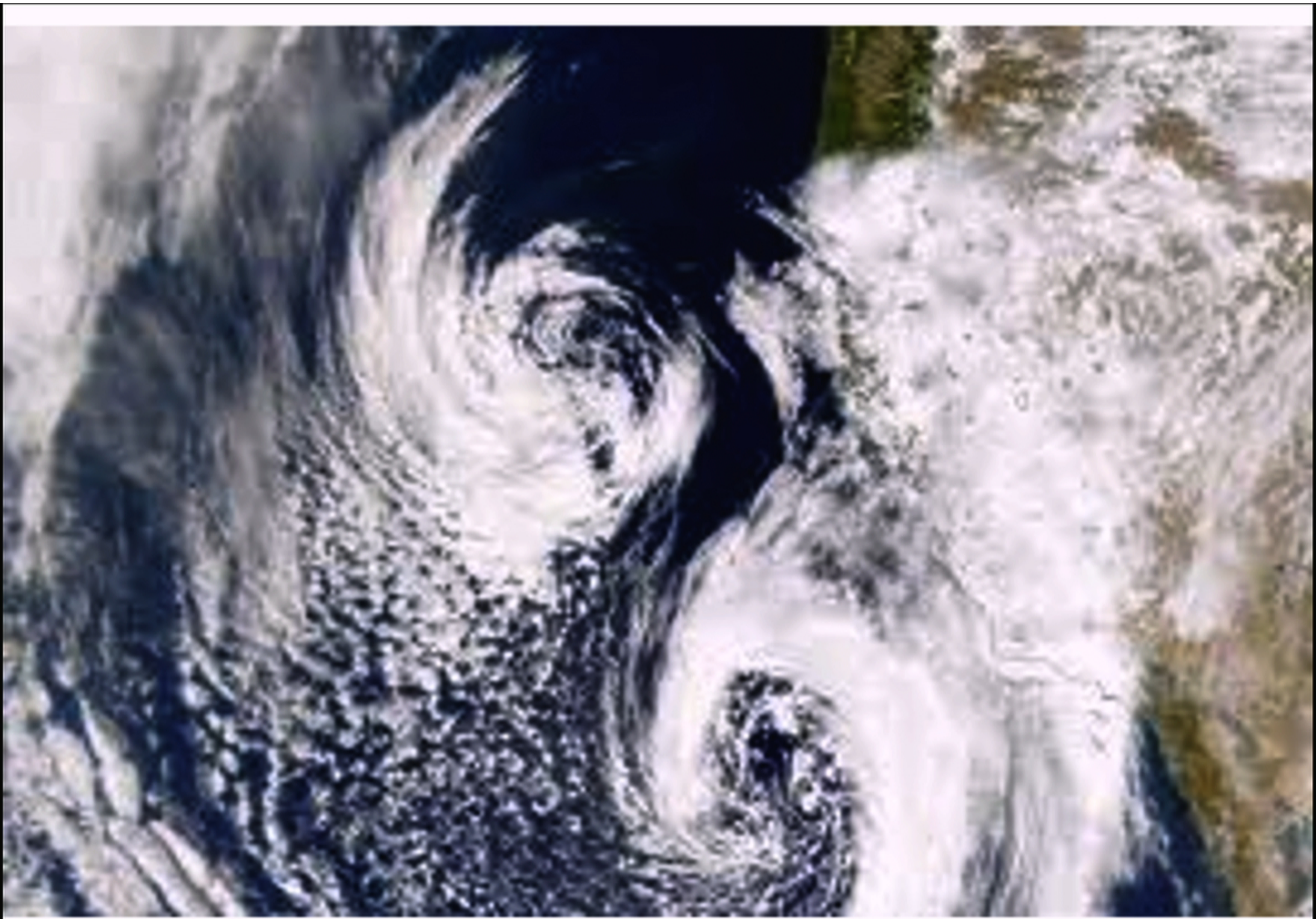
$$\rho(p', s) > \rho(p', s') \Rightarrow \left(\frac{\partial \rho}{\partial s}\right)_p \frac{ds}{dz} < 0$$

$$C_p dT < g dz$$

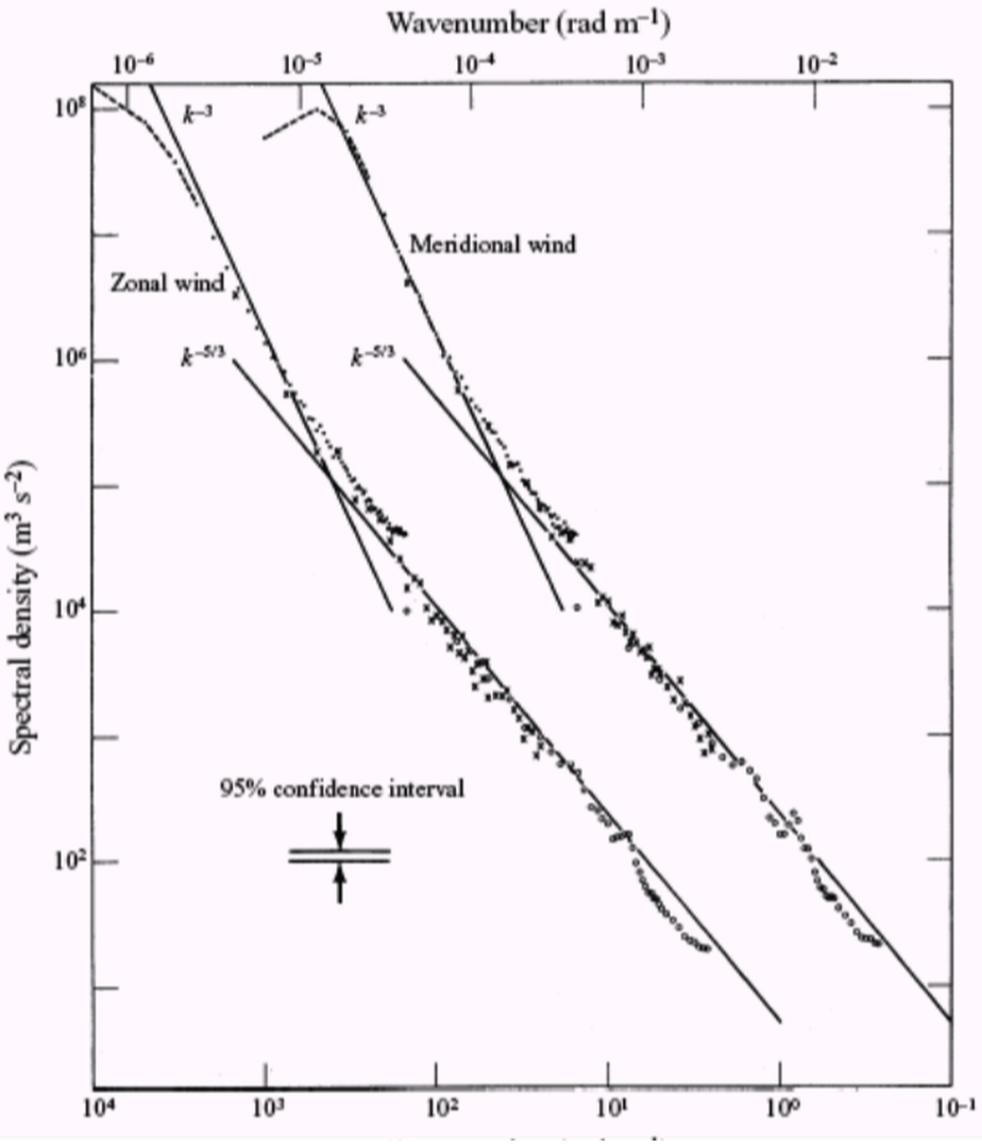
$$-\frac{dT}{dz} < \frac{g}{c_p} \sim 10^\circ/km$$



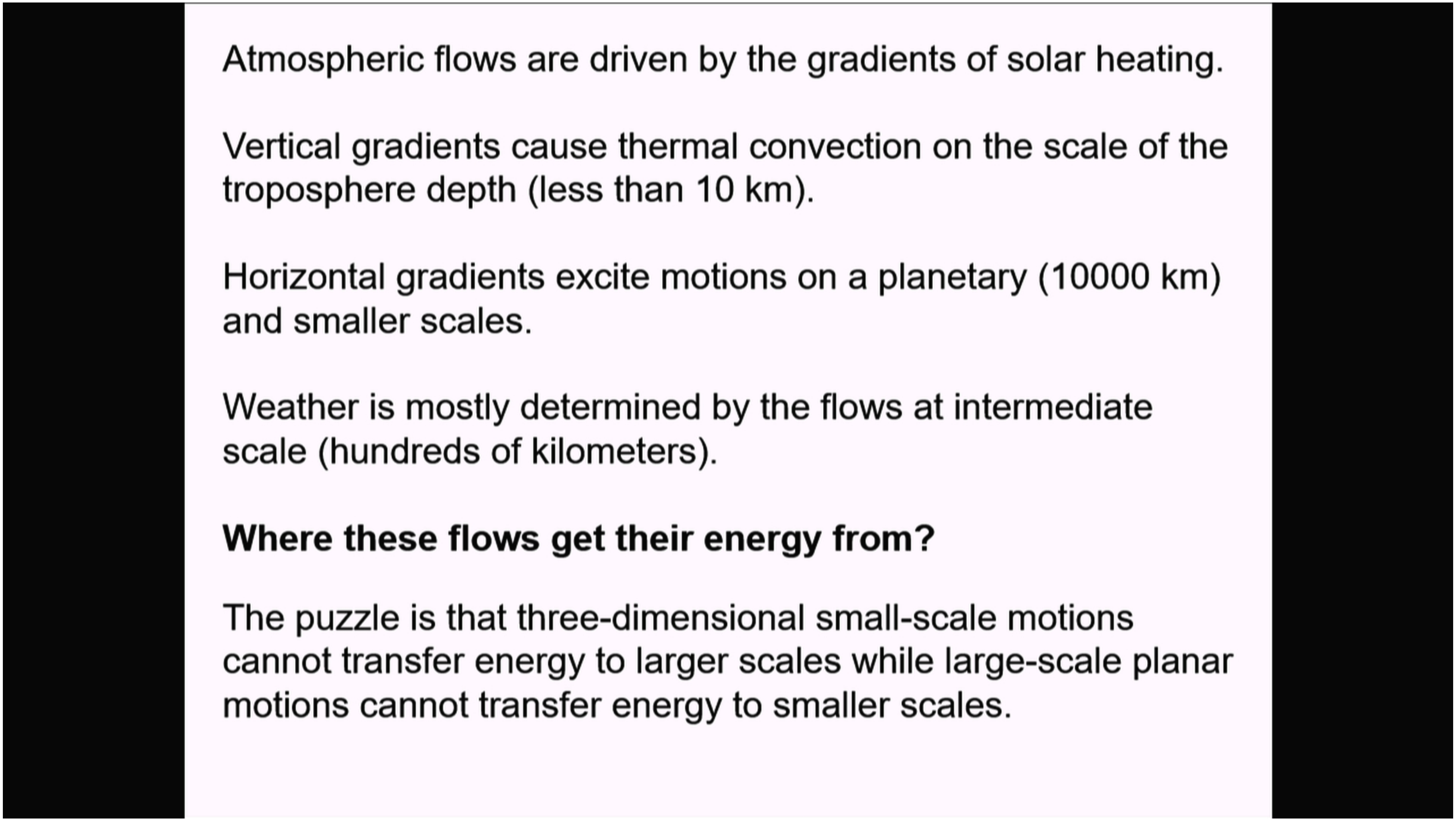

IR
8 km
NOAA



Atmospheric spectrum



Nastrom, Gage,
J. Atmosph. Sci. 1985



Atmospheric flows are driven by the gradients of solar heating.

Vertical gradients cause thermal convection on the scale of the troposphere depth (less than 10 km).

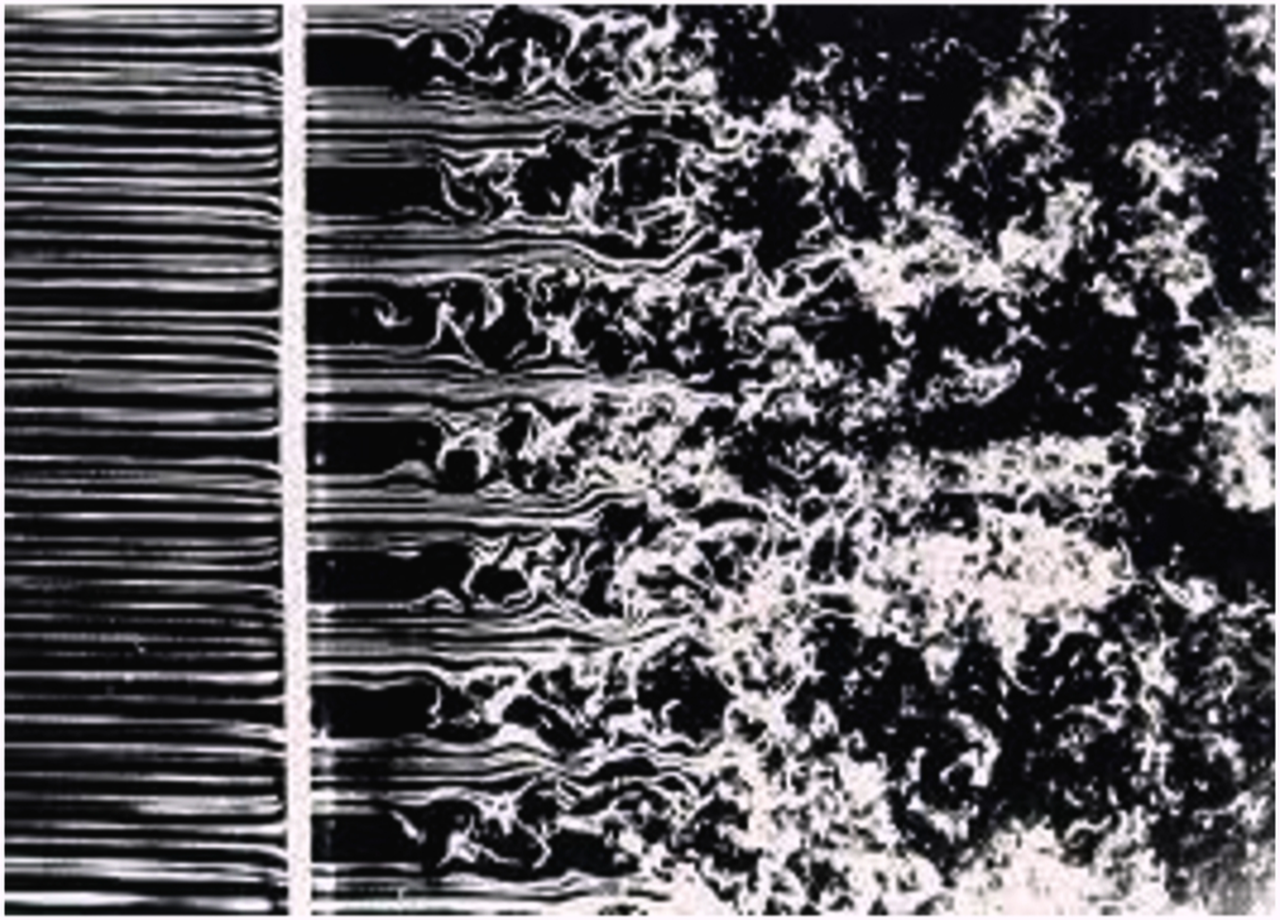
Horizontal gradients excite motions on a planetary (10000 km) and smaller scales.

Weather is mostly determined by the flows at intermediate scale (hundreds of kilometers).

Where these flows get their energy from?

The puzzle is that three-dimensional small-scale motions cannot transfer energy to larger scales while large-scale planar motions cannot transfer energy to smaller scales.

Three- dimensional turbulence: fragmentation and energy transfer to smaller scales



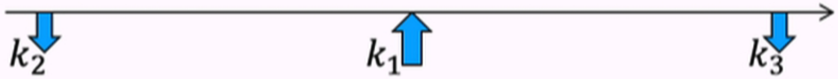
Two- dimensional turbulence: vortex clustering and merging, and energy transfer to smaller scales

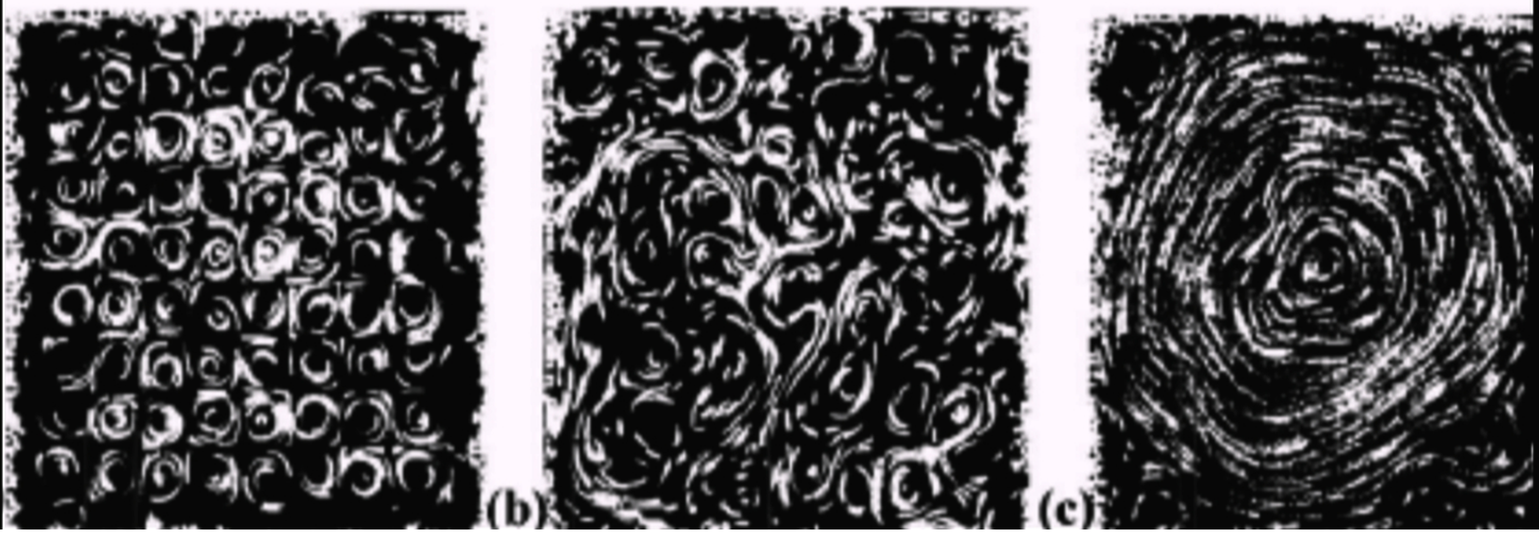
Two conservation laws: energy $v^2/2$ and squared vorticity $(\nabla \times v)^2$

$$v_1^2 = v_2^2 + v_3^2$$

$$k_1^2 v_1^2 = k_2^2 v_2^2 + k_3^2 v_3^2$$

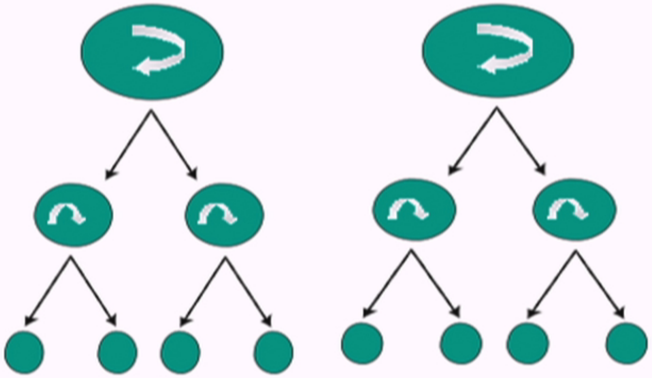
$$v_2^2 = v_1^2 \frac{k_3^2 - k_1^2}{k_3^2 - k_2^2} \qquad v_3^2 = v_1^2 \frac{k_1^2 - k_2^2}{k_3^2 - k_2^2}$$

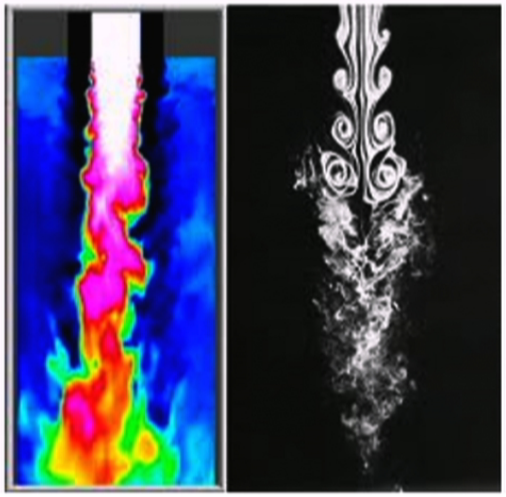


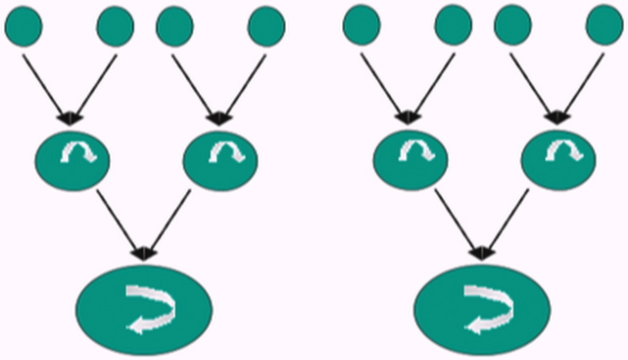


Direct cascade

Inverse cascade







Kolmogorov energy cascade

$$\frac{\text{kinetic energy } (\delta v_r)^2}{\text{time } r/\delta v_r} = \text{energy flux } \epsilon$$

$$\epsilon = \nu \langle |\nabla v|^2 \rangle \simeq v_{rms}^3 / L$$

$$(\delta v_r)^3 \sim \epsilon r$$

$$S_3 = \langle (\delta v_r)^3 \rangle = -\frac{12 \epsilon r}{d(d+2)}$$

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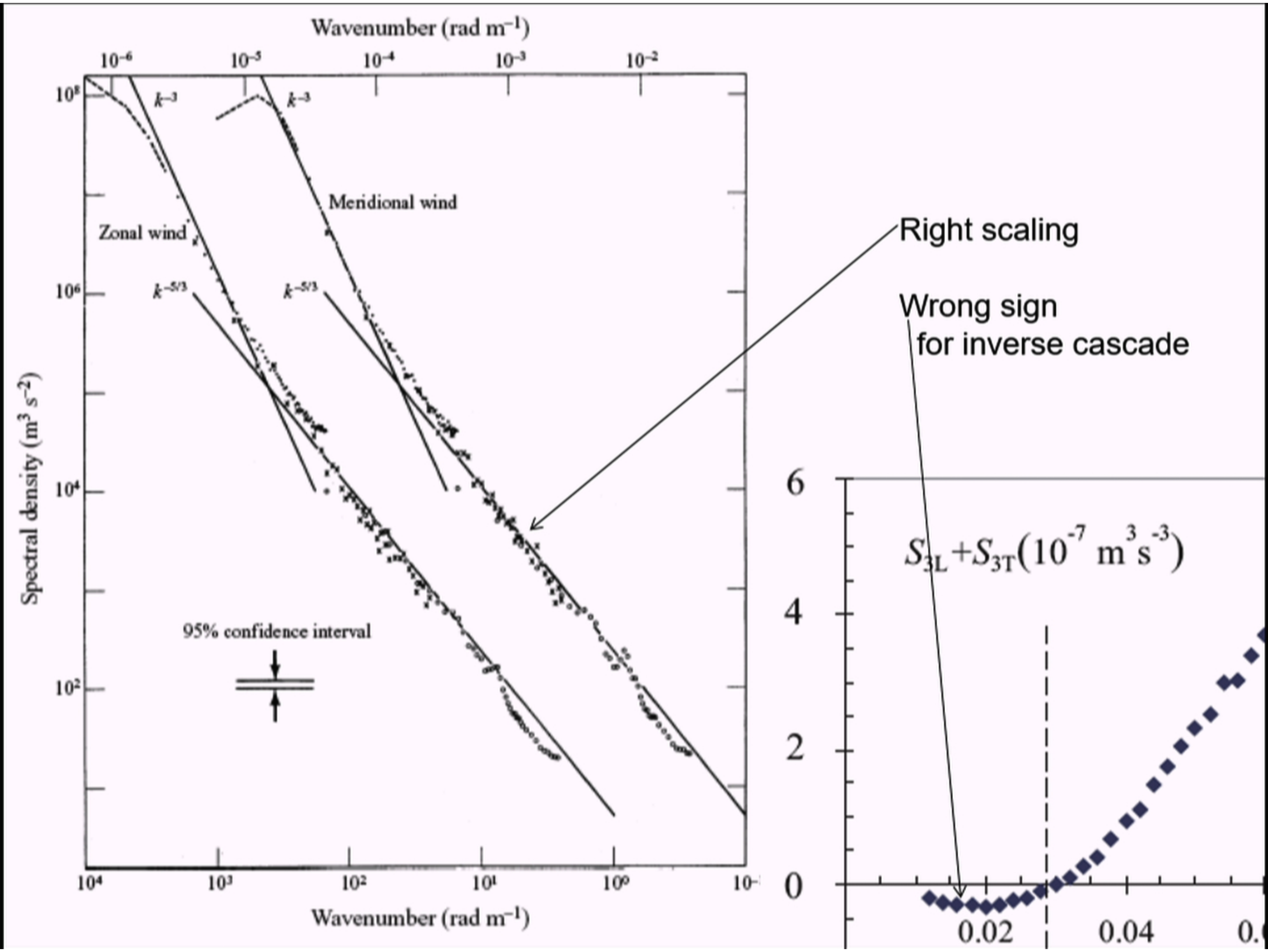


Right scaling

Wrong sign
for inverse cascade

Laboratory experiments

J. Paret and P. Tabeling 3127

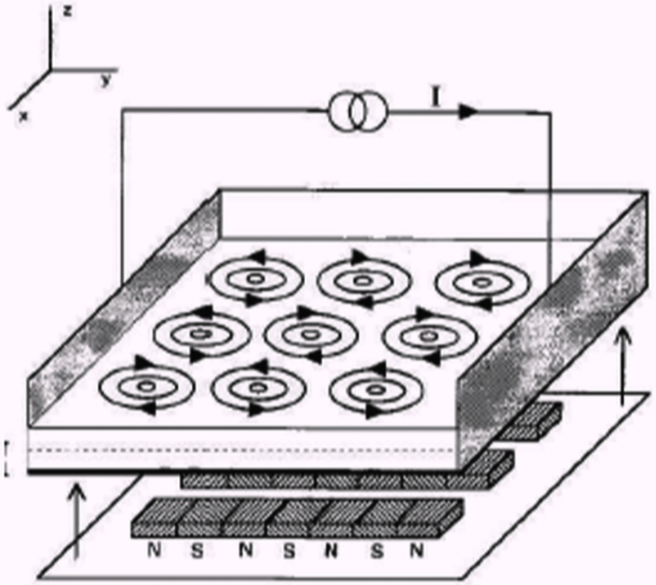
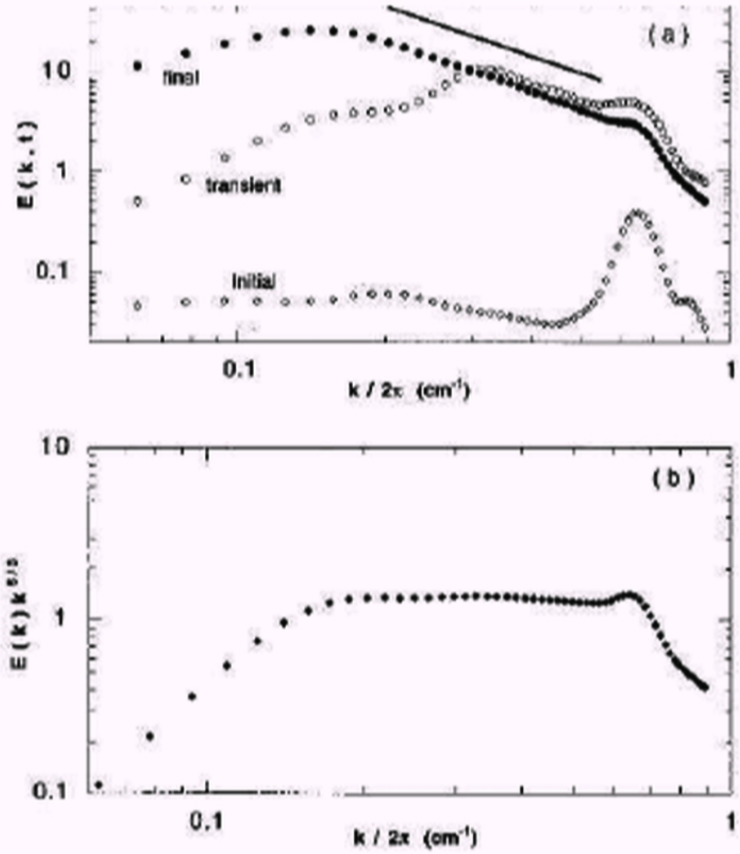


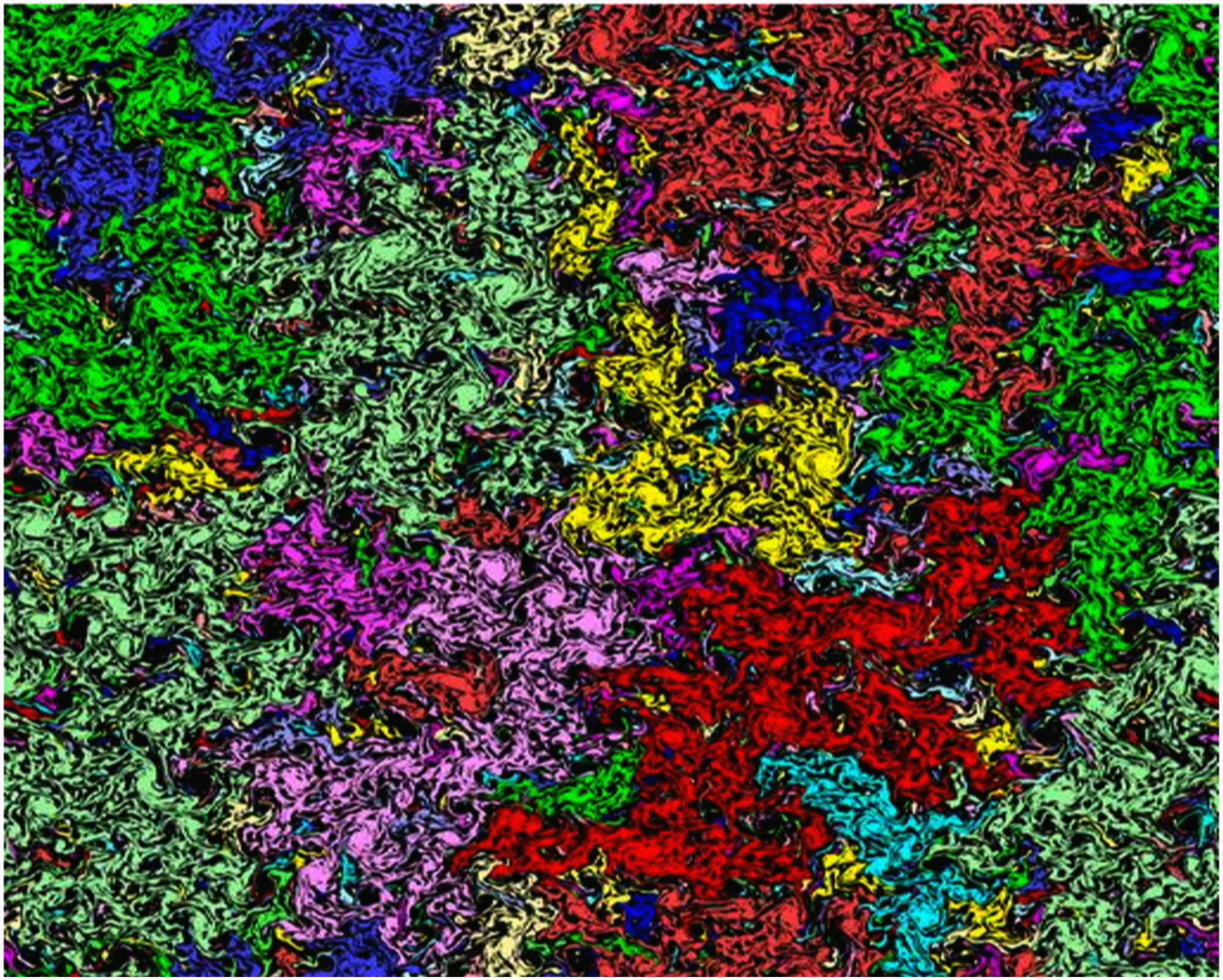
FIG. 1. The experimental set-up.

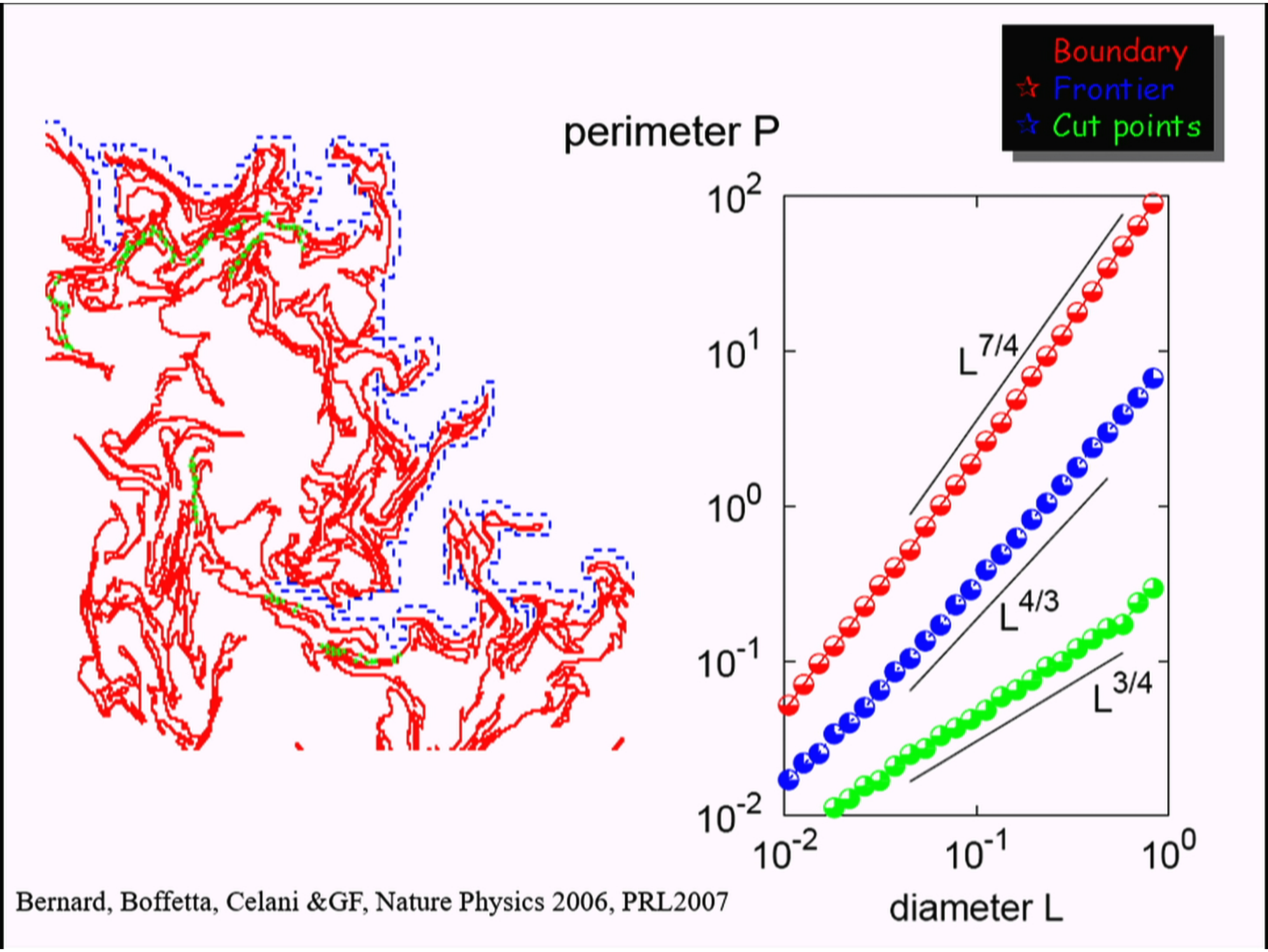


We expect from turbulence
fragmentation, **mixing** and **loss of coherence**.

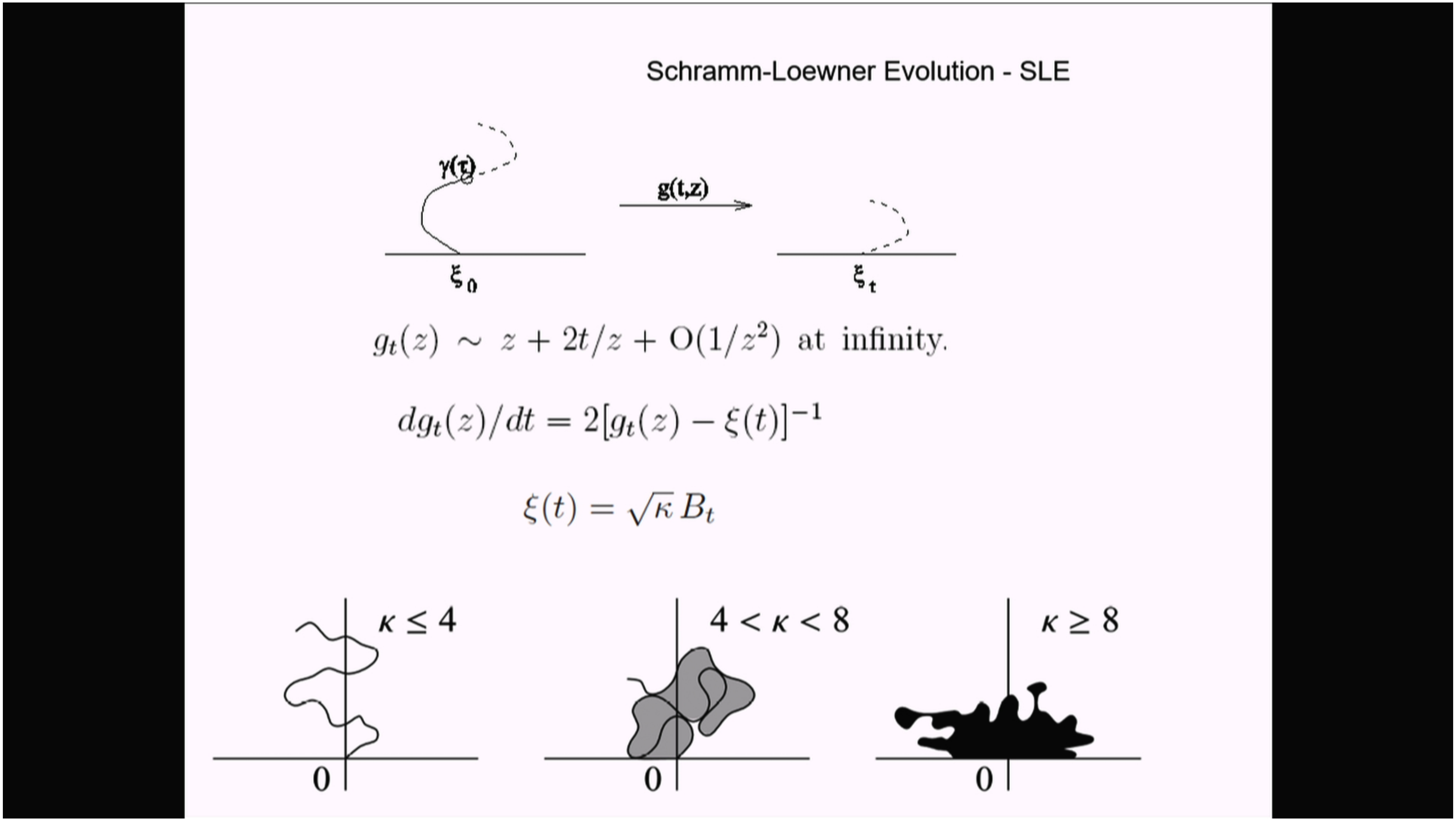
However,
an inverse turbulent cascade proceeds from small to large scales and brings some **self-organization**

and eventually appearance of
a **coherent system-size condensate**.





Bernard, Boffetta, Celani &GF, Nature Physics 2006, PRL2007



Schramm-Loewner Evolution - SLE

$$g_t(z) \sim z + 2t/z + \mathrm{O}(1/z^2) \text{ at infinity.}$$

$$dg_t(z)/dt = 2[g_t(z) - \xi(t)]^{-1}$$

$$\xi(t) = \sqrt{\kappa}\, B_t$$

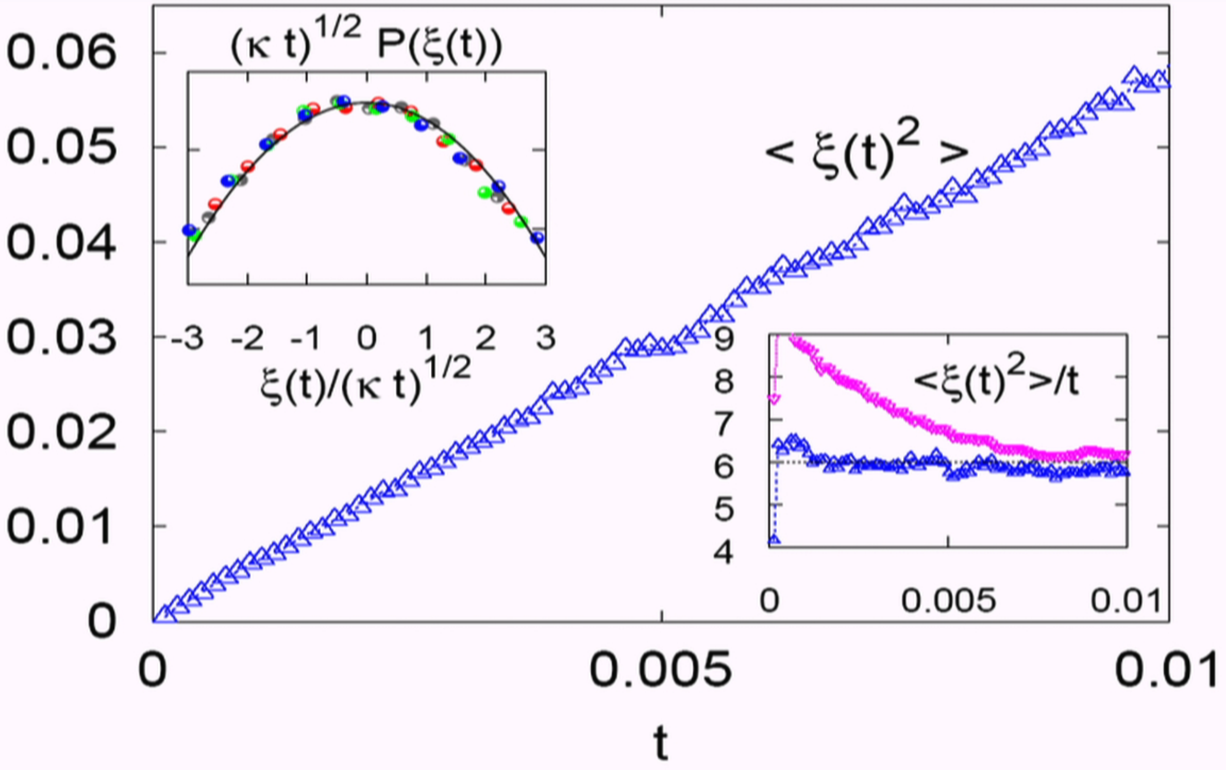
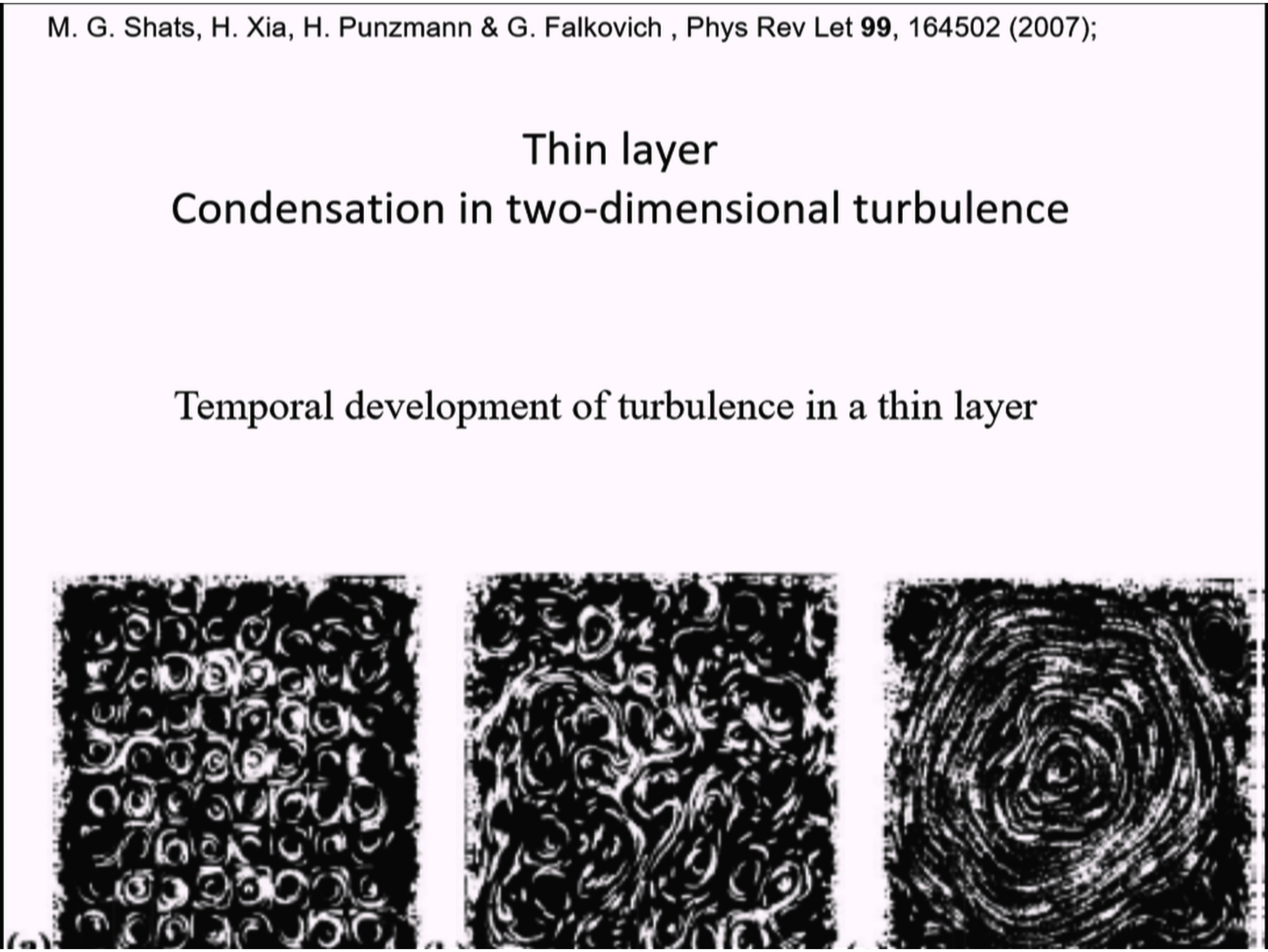


FIG. 4: The driving function is an effective diffusion process with diffusion coefficent $\kappa = 6 \pm 0.3$. The inverse cascade range corresponds to $5 \cdot 10^{-5} < t < 10^{-2}$. *Main frame*: the linear behaviour of $\langle \xi(t)^2 \rangle$. *Lower-right inset* : Diffusivity: blue for vorticity isolines, pink for the field with randomized phases. *Upper-left inset*: the probability density function of the rescaled driving function $\xi(t)/\sqrt{\kappa t}$ at four different times $t = 0.0012, 0.003, 0.006, 0.009$; the solid line is the Gaussian distribution $g(x) = (2\pi)^{-1/2}\exp(-x^2/2)$.

Different systems producing SLE

- Critical phenomena with local Hamiltonians
- Random walks, non necessarily local
- Inverse cascades in turbulence
- Nodal lines of wave functions in chaotic systems
- Spin glasses
- Rocky coastlines



M. G. Shats, H. Xia, H. Punzmann & G. Falkovich , Phys Rev Let **99**, 164502 (2007);

Thin layer
Condensation in two-dimensional turbulence

Temporal development of turbulence in a thin layer

$$E=E_0e^{-at}$$

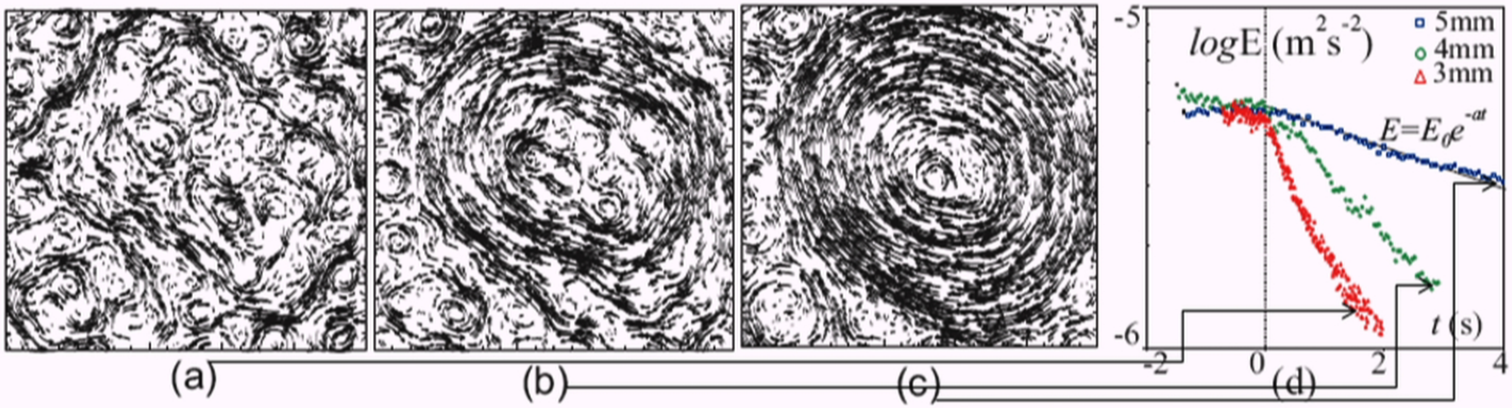
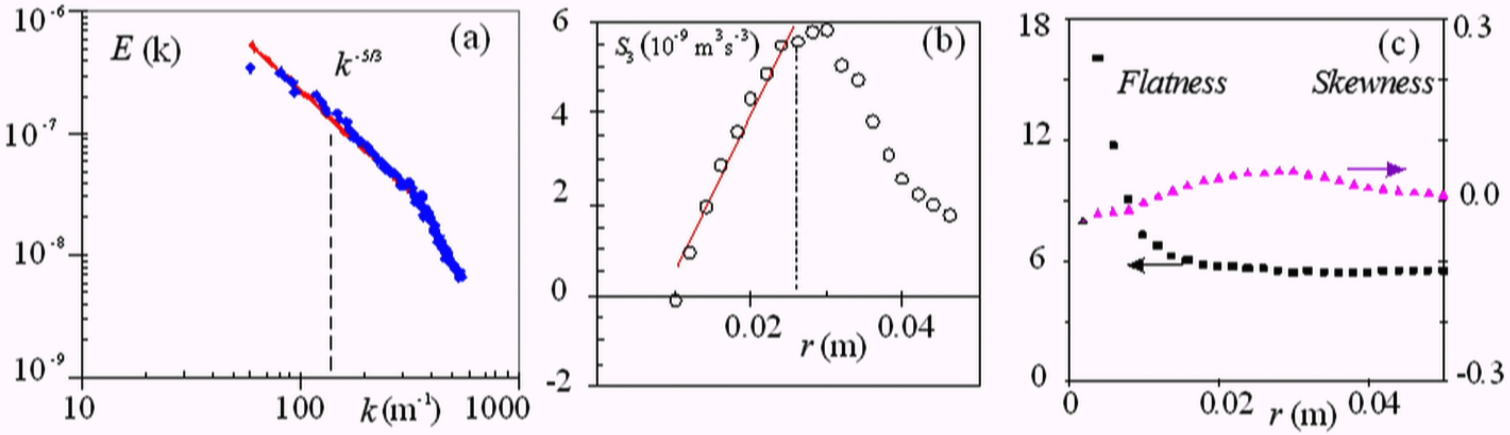


FIGURE 1. Time-averaged velocity field of the condensate in the square box of $L \approx 0.10$ m at different thicknesses of the bottom fluid (Fluorinert FC-77): (a) $\Delta h_b = 3$ mm, $\alpha = 0.3$ s^{-1}, (b) $\Delta h_b = 4$ mm, $\alpha = 0.15$ s^{-1} and (c) $\Delta h_b = 5$ mm, $\alpha = 0.05$ s^{-1}. (d) Decay of the total kinetic energy for cases (a-c).

Mean subtraction recovers isotropic turbulence

1. Compute time-average velocity field (400 snapshots)

2. Subtract from 400 instantaneous velocity fields

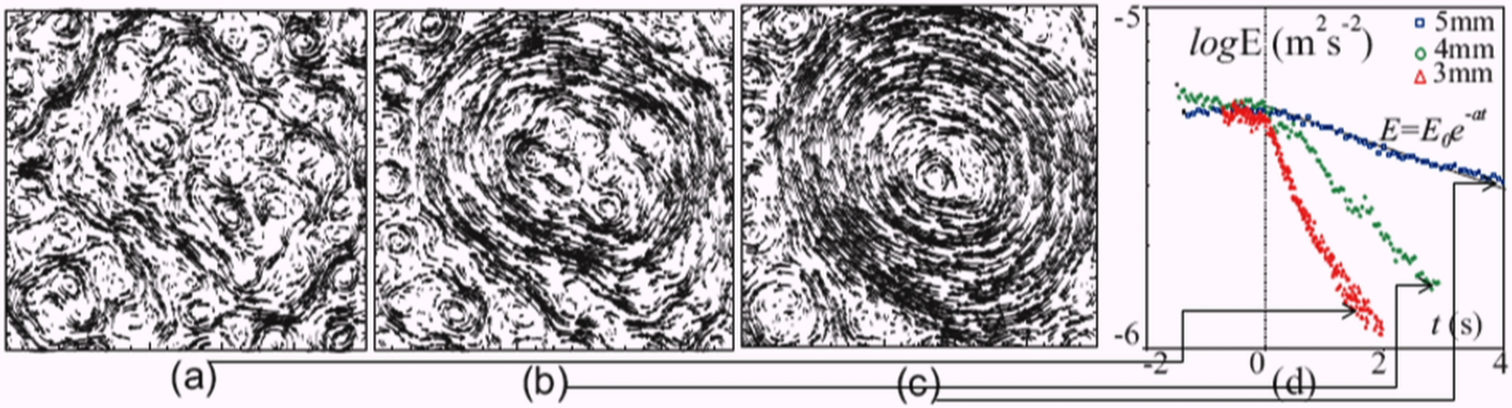


Recover ~ $k^{-5/3}$ spectrum in the energy range

Kolmogorov law – linear S3 (r) dependence in the “turbulence range”;

Kolmogorov constant C≈7

$$E = E_0 e^{-at}$$



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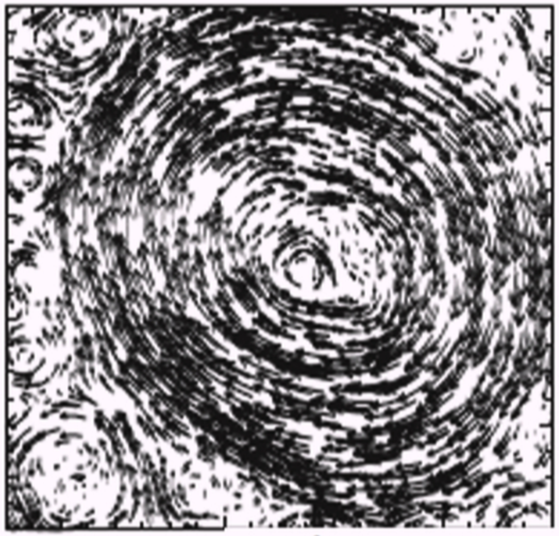
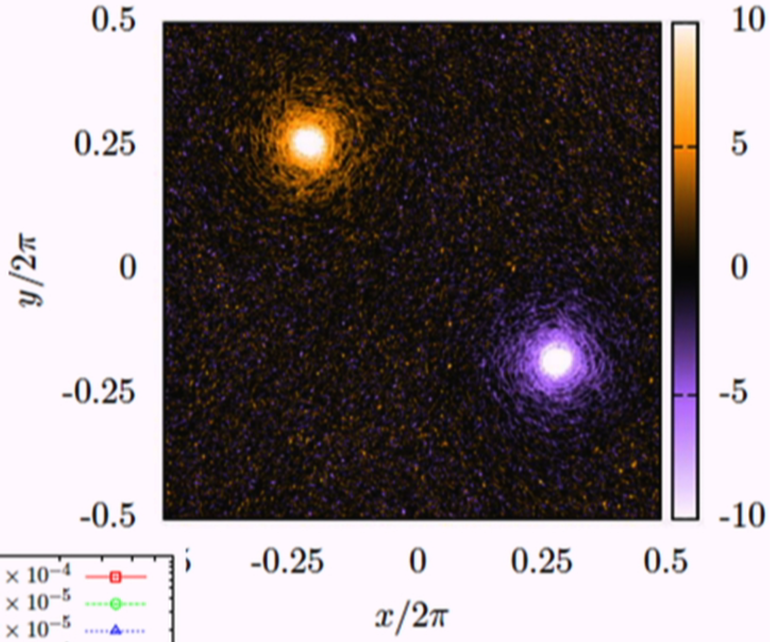
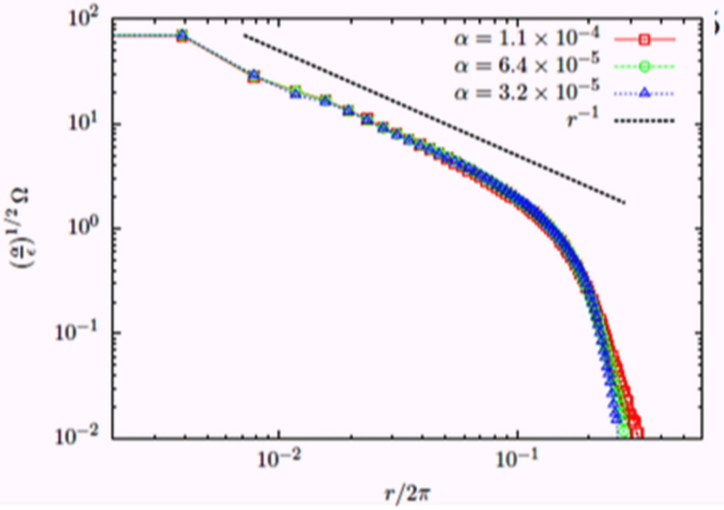
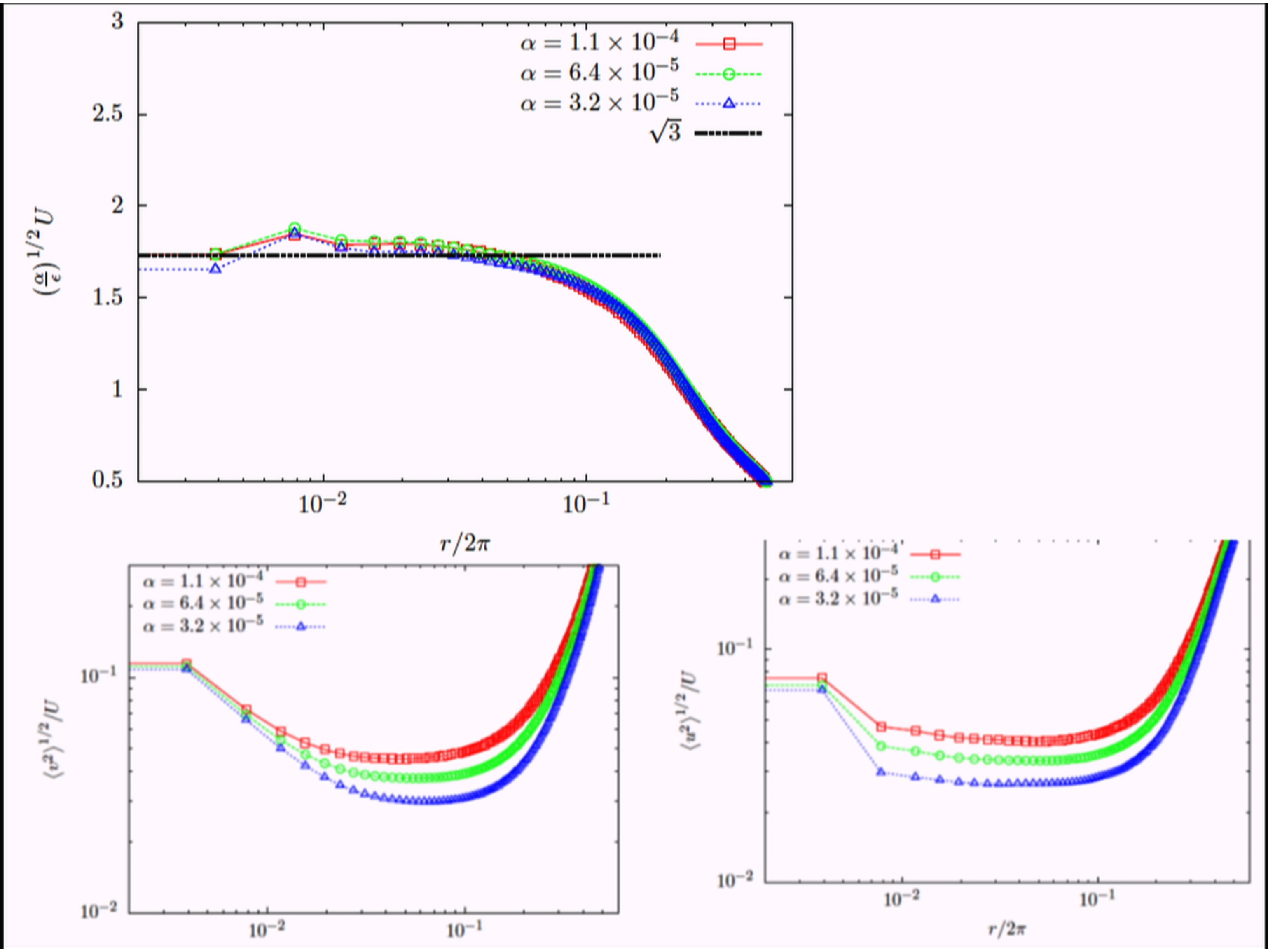
Universal profile of a coherent vortex



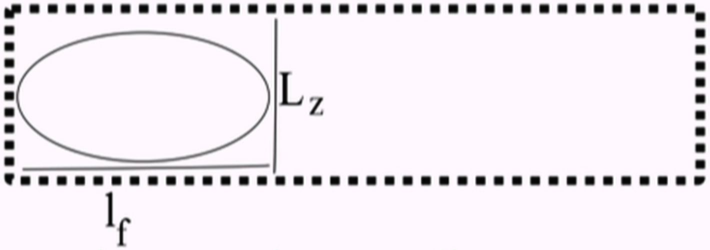


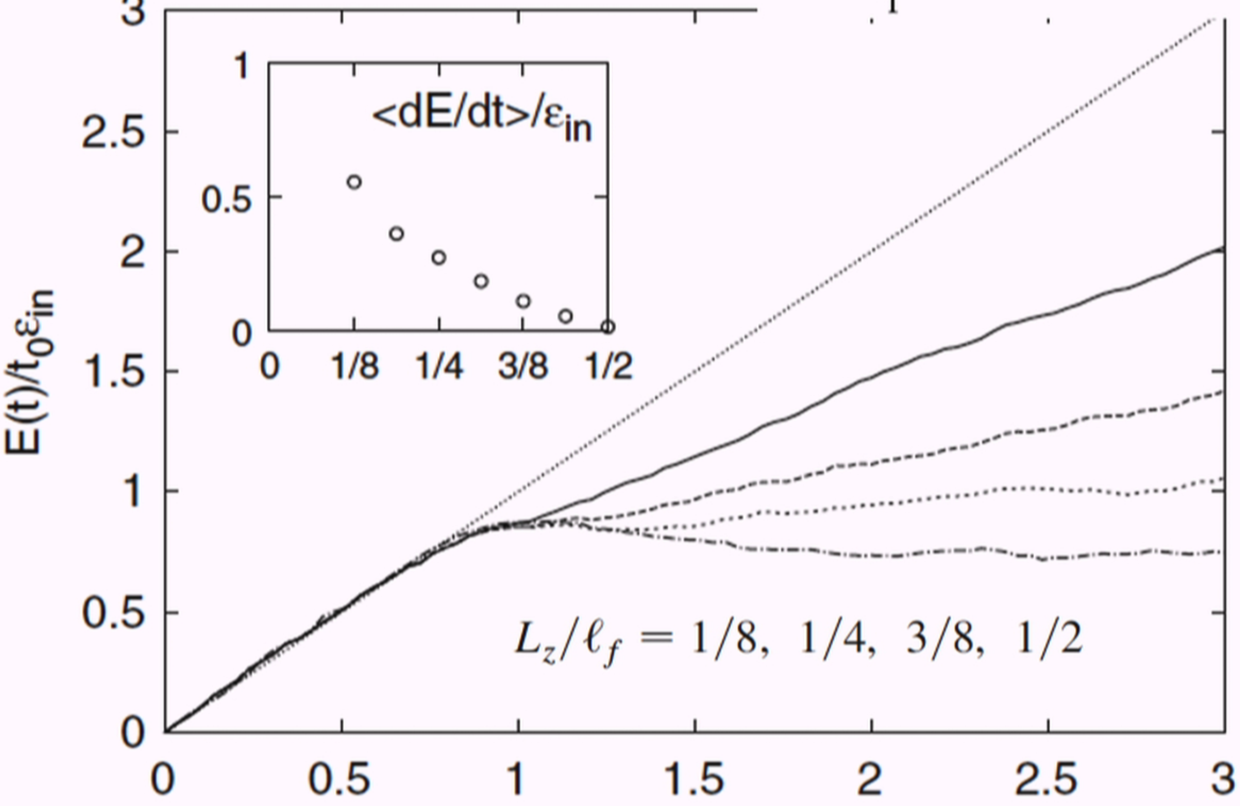


G. Boffetta, J. Laurie, I. Kolokolov,
V. Lebedev, GF, 2014

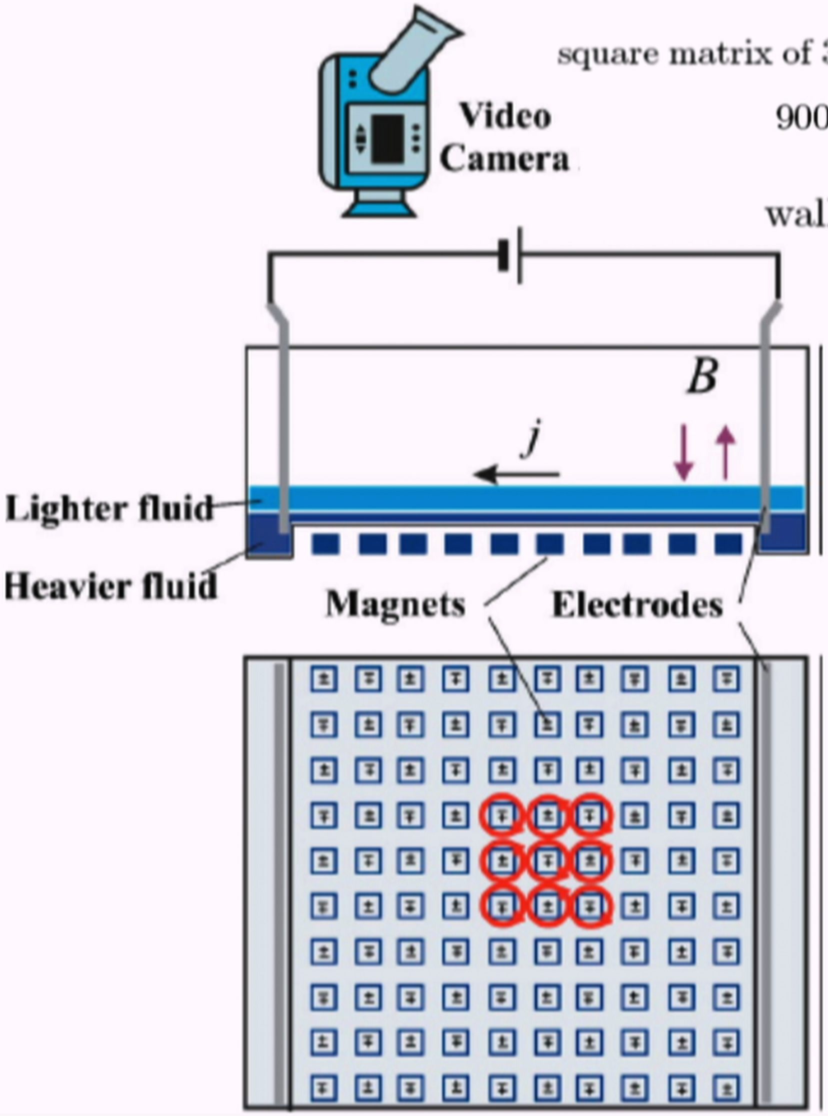

$\alpha = 1.1 \times 10^{-4}$
$\alpha = 6.4 \times 10^{-5}$
$\alpha = 3.2 \times 10^{-5}$
$\sqrt{3}$
$\left(\frac{\alpha}{\epsilon}\right)^{1/2} U$
3
2.5
2
1.5
1
0.5
10^{-2}
10^{-1}
$r/2\pi$
$\alpha = 1.1 \times 10^{-4}$
$\alpha = 6.4 \times 10^{-5}$
$\alpha = 3.2 \times 10^{-5}$
$\langle v^2 \rangle^{1/2}/U$
10^{-1}
10^{-2}
10^{-2}
10^{-1}
$\alpha = 1.1 \times 10^{-4}$
$\alpha = 6.4 \times 10^{-5}$
$\alpha = 3.2 \times 10^{-5}$
$\langle u^2 \rangle^{1/2}/U$
10^{-1}
10^{-2}
10^{-2}
10^{-1}
$r/2\pi$

To understand atmosphere
one needs to move
from thin to **thick layers**





Antonio Celani,[1] Stefano Musacchio,[2,3] and Dario Vincenzi[3] PRL **104,** 184506 (2010)



square matrix of 30×30 permanent magnets (10 mm apart)

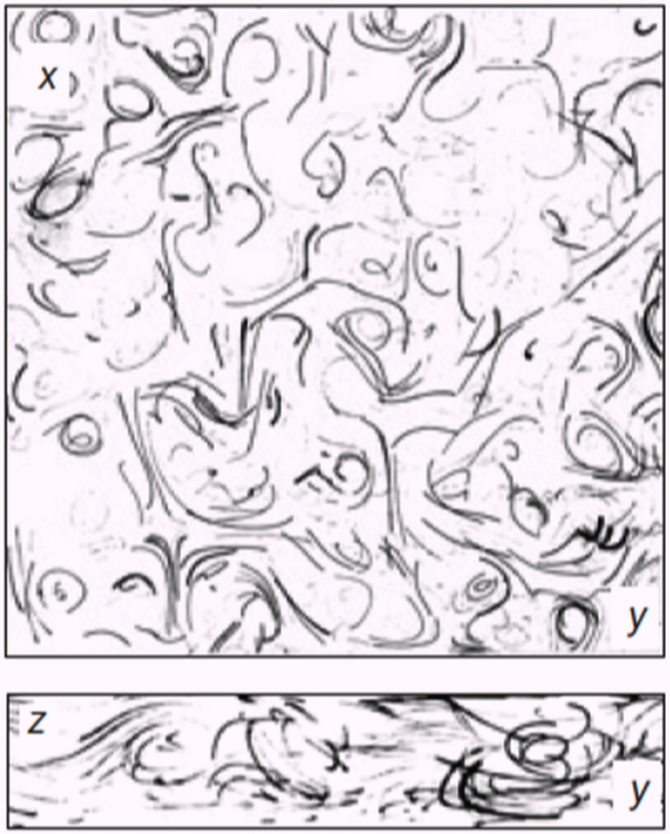
900 vortices (9 mm in diameter)

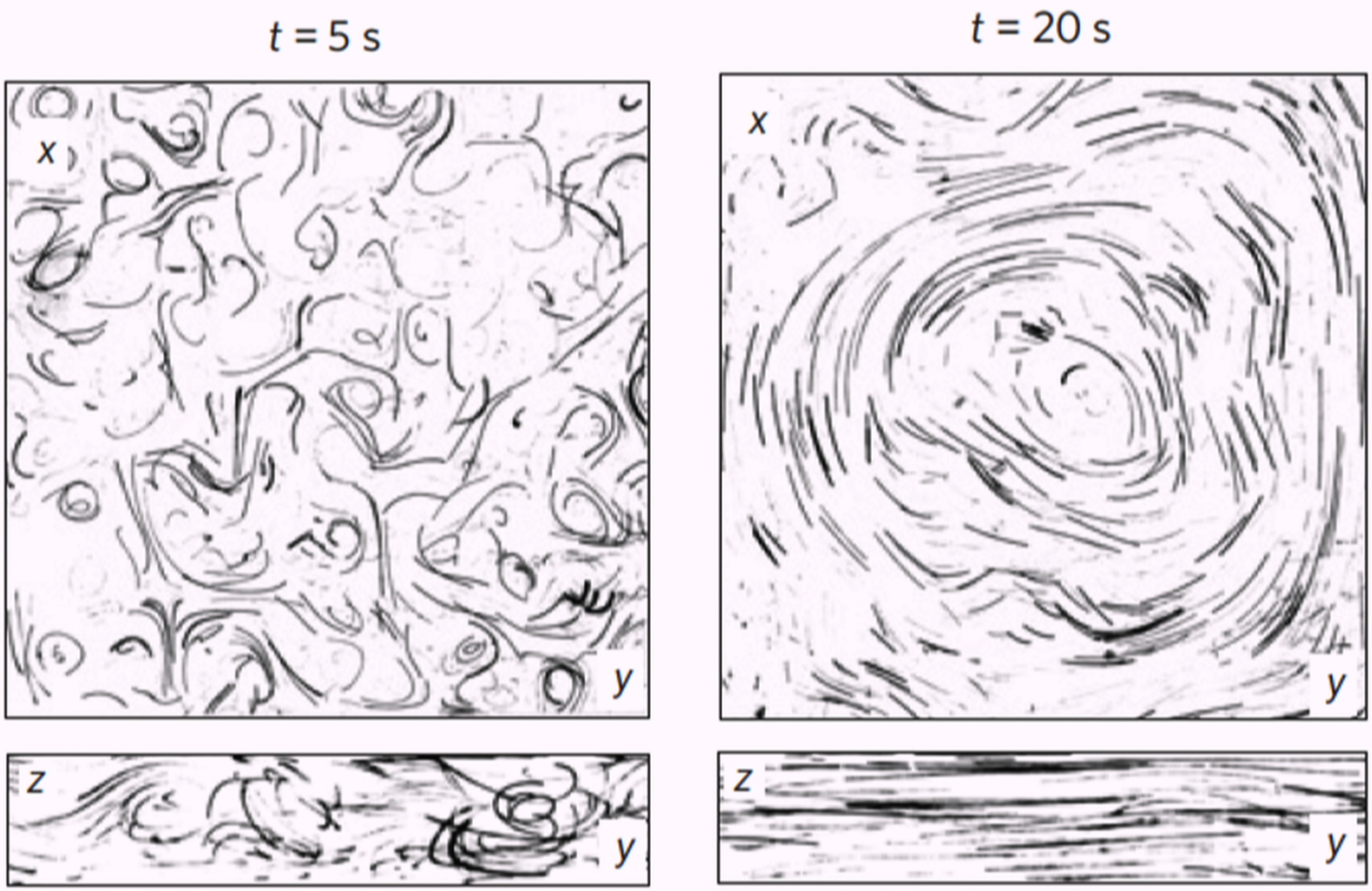
walls of the fluid cell (0.3×0.3 m^2)

insertable square boundaries

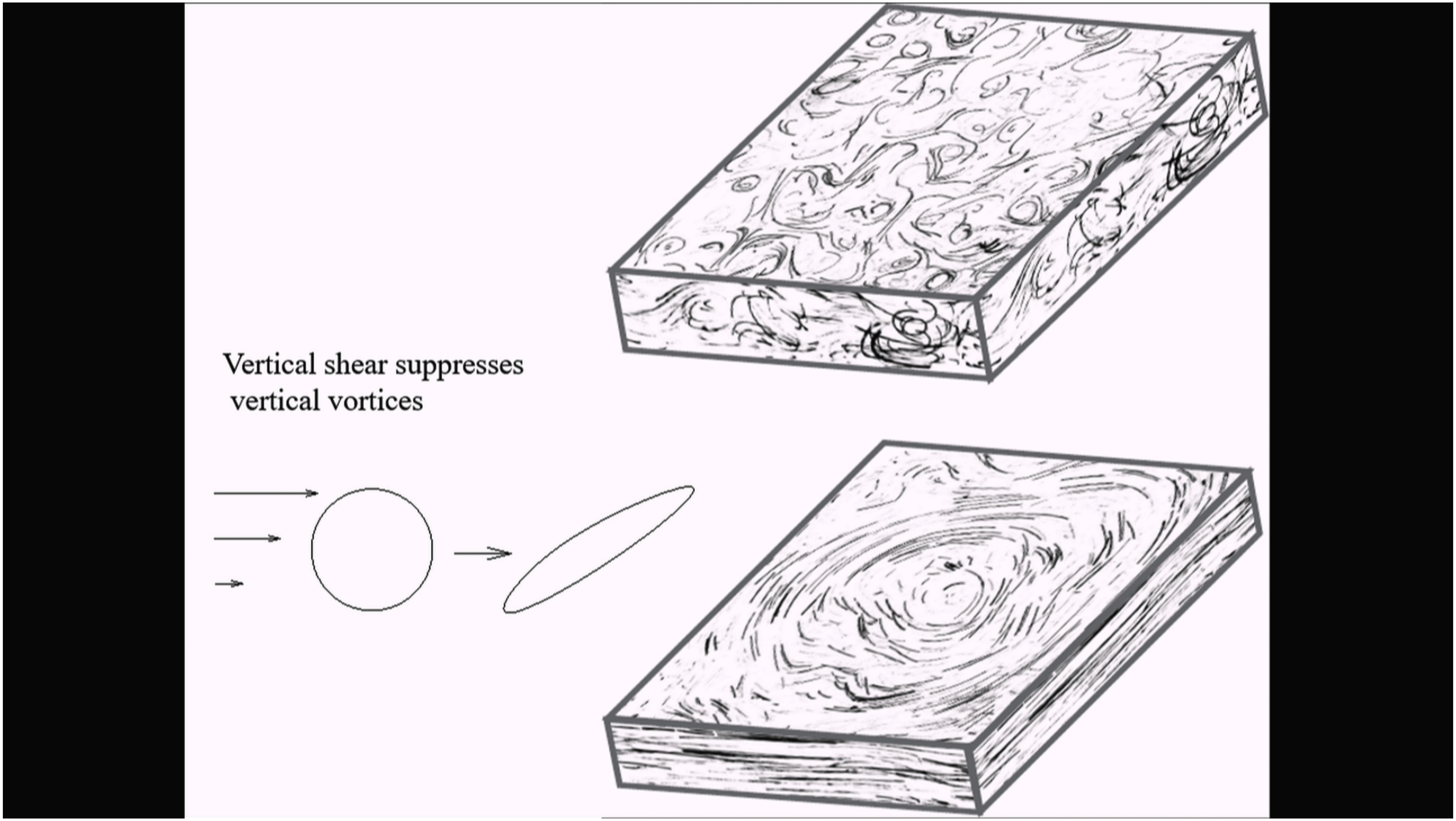
$$L = (0.09 - 0.24)\ \mathrm{m}$$

$t = 5$ s

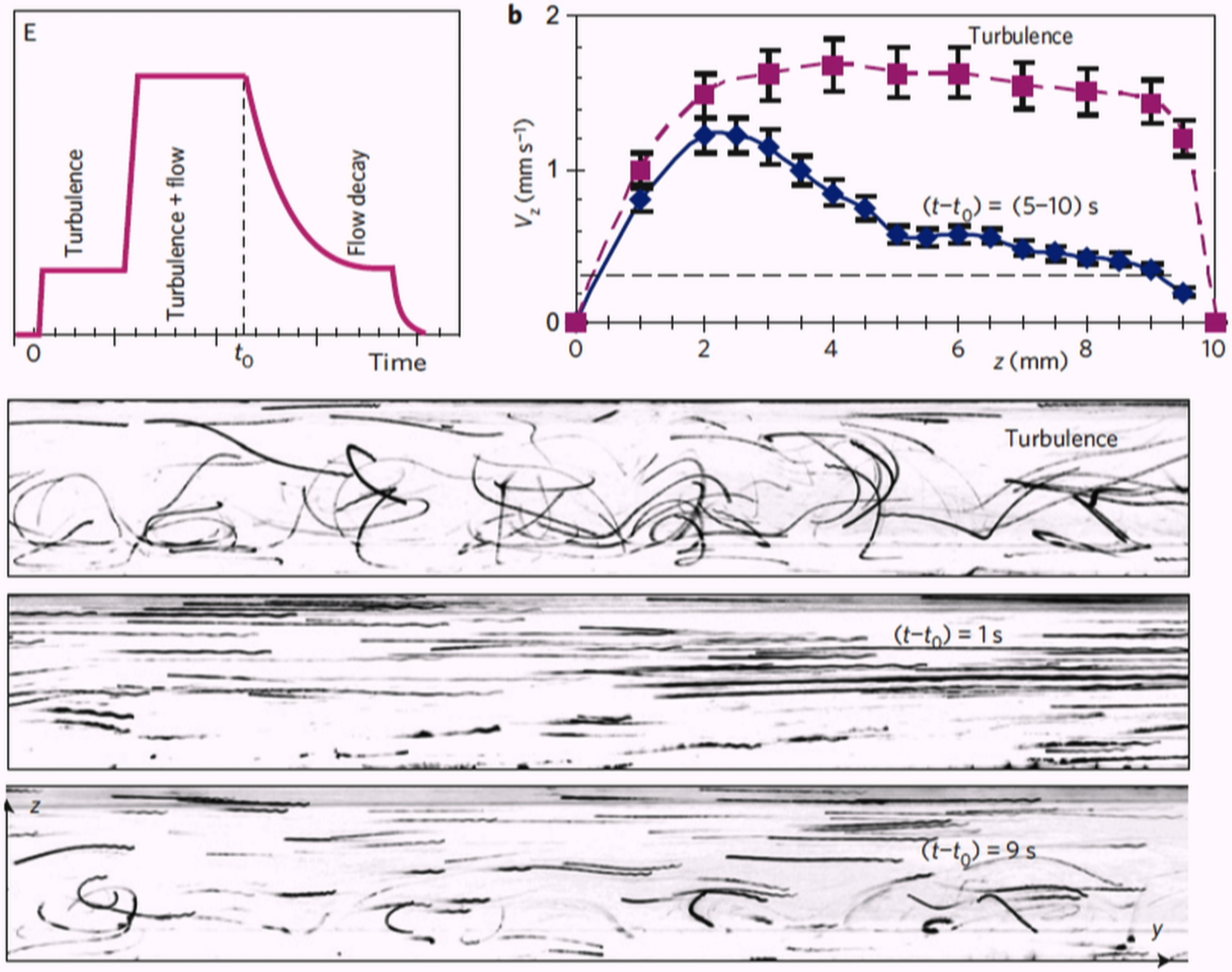


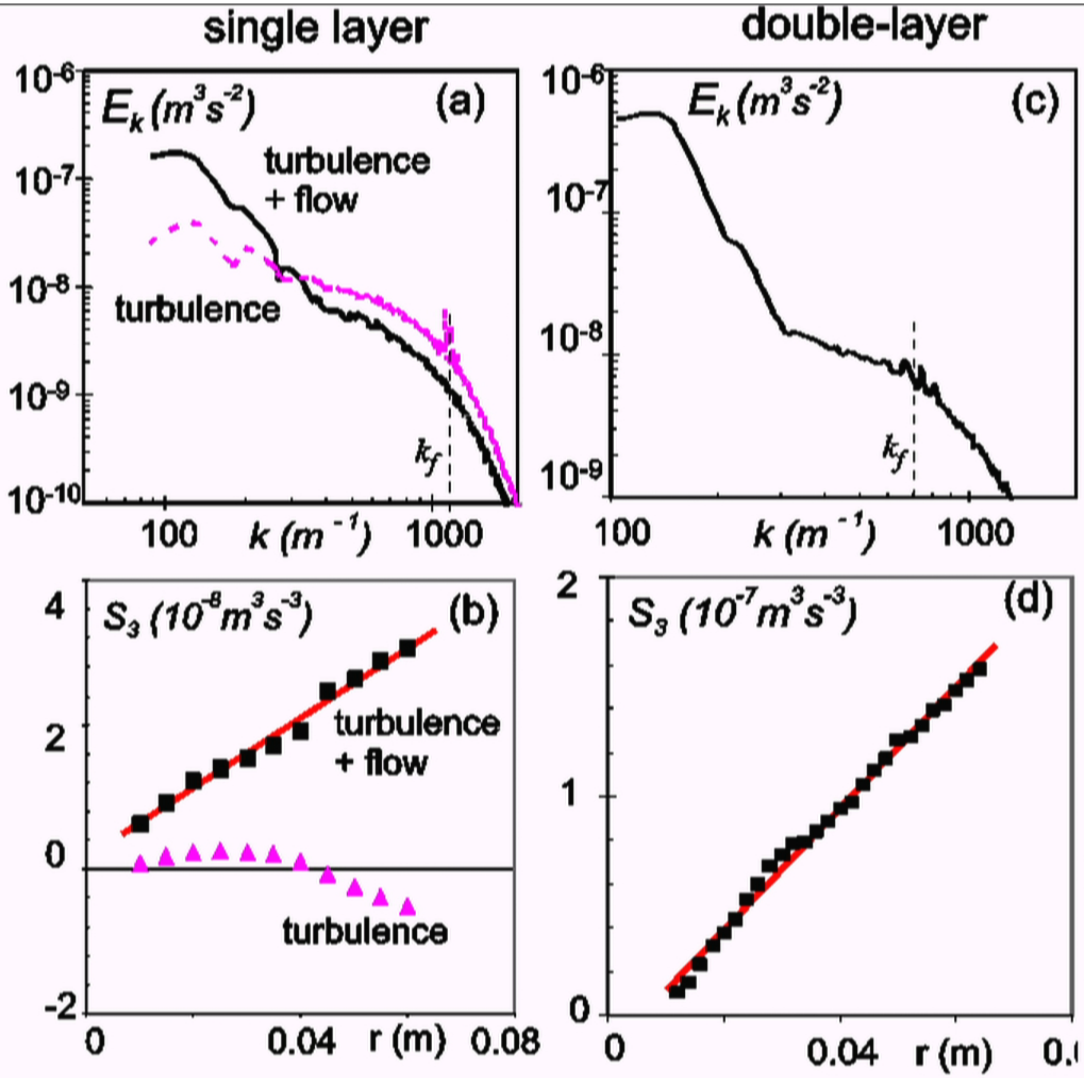


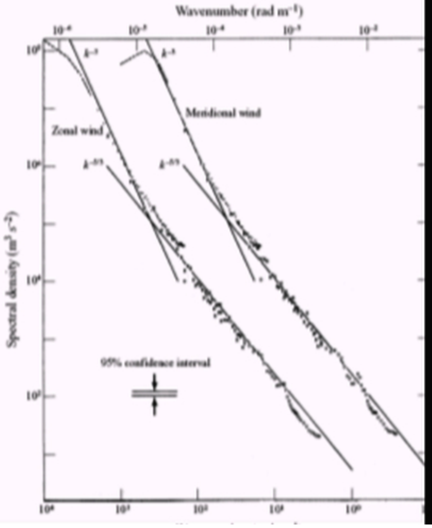
H. Xia[1], D. Byrne[1], G. Falkovich[2] and M. Shats[1] NATURE PHYSICS, April 1, 2011



Vertical shear suppresses vertical vortices


E
Turbulence
Turbulence + flow
Flow decay
0
t_0
Time
b
2
1
0
V_z (mm s^{-1})
Turbulence
$(t-t_0)$ = (5–10) s
0
2
4
6
8
10
z (mm)
Turbulence
$(t-t_0)$ = 1 s
z
$(t-t_0)$ = 9 s
y


single layer
double-layer
$E_k (m^3 s^{-2})$
(a)
turbulence + flow
turbulence
k_f
10^{-6}
10^{-7}
10^{-8}
10^{-9}
10^{-10}
100
$k\ (m^{-1})$
1000
$E_k (m^3 s^{-2})$
(c)
k_f
10^{-6}
10^{-7}
10^{-8}
10^{-9}
100
$k\ (m^{-1})$
1000
$S_3\ (10^{-8} m^3 s^{-3})$
(b)
turbulence + flow
turbulence
4
2
0
-2
0
0.04
r (m)
0.08
$S_3\ (10^{-7} m^3 s^{-3})$
(d)
2
1
0
0
0.04
r (m)
0.(


Wavenumber (rad m^{-1})
10^{-6}
10^{-5}
10^{-4}
10^{-3}
10^{-2}
k^{-3}
k^{-3}
Meridional wind
Zonal wind
$k^{-5/3}$
$k^{-5/3}$
Spectral density ($m^3 s^{-2}$)
95% confidence interval

Moral

- A strong large-scale flow effectively suppresses fluctuations in the vertical velocity.
- The resulting flow is planar even at small scales yet it is three-dimensional as it depends strongly on the vertical coordinate.
- Turbulence in such flows transfers energy towards large scales.

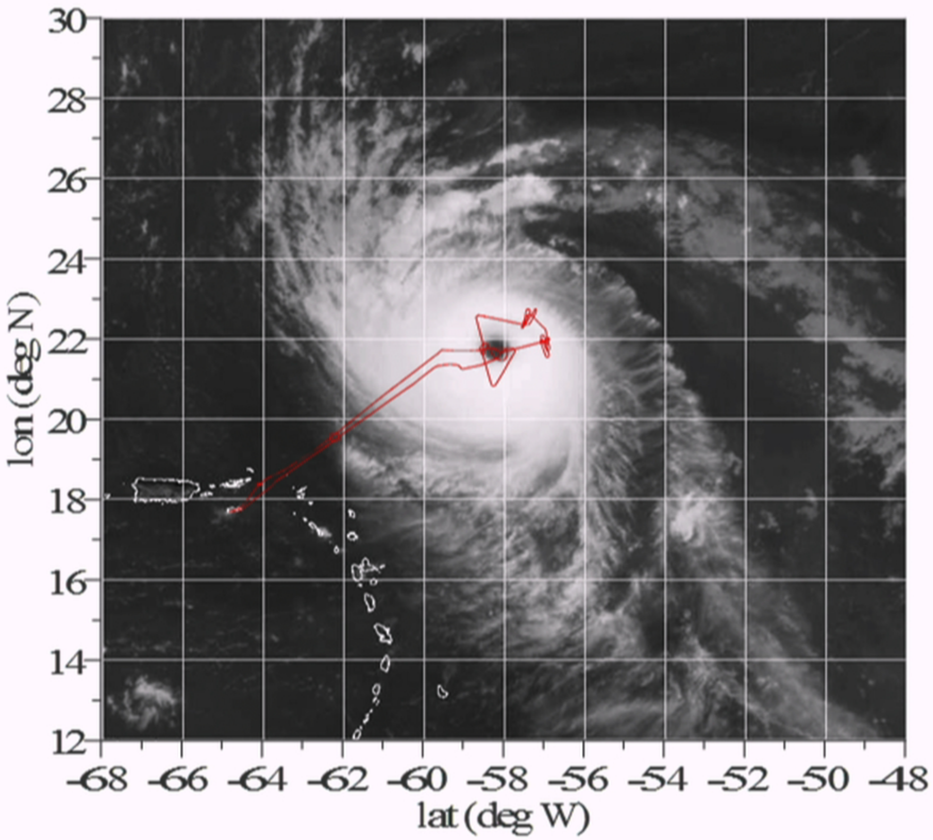
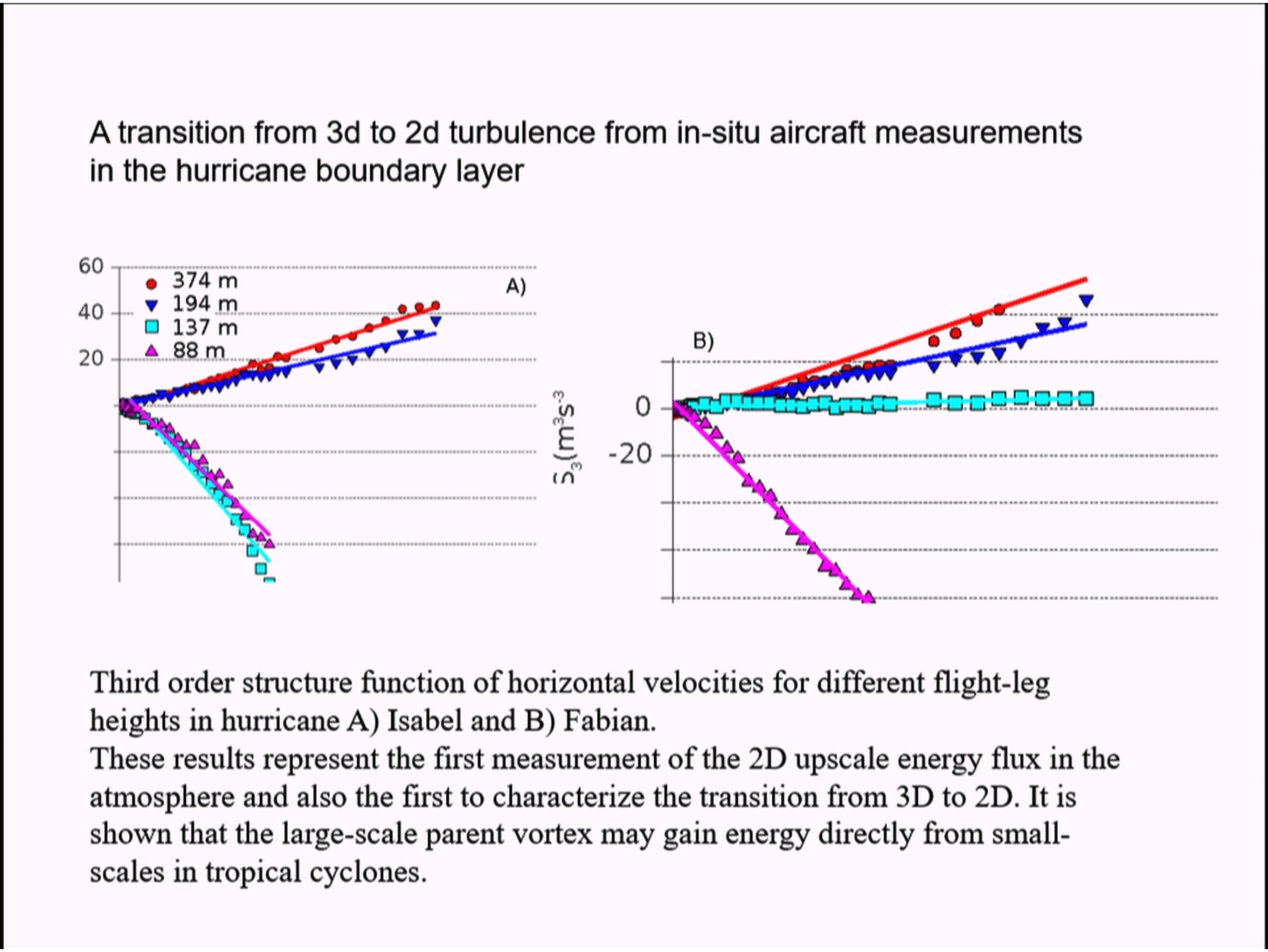


Figure 1. Satellite image of hurricane Isabel, 12th September 2003 with the flight track of the WP-3D Orion aircraft (N43RF) overlaid. Stepped decent measurements of the boundary layer performed between outer rain bands.

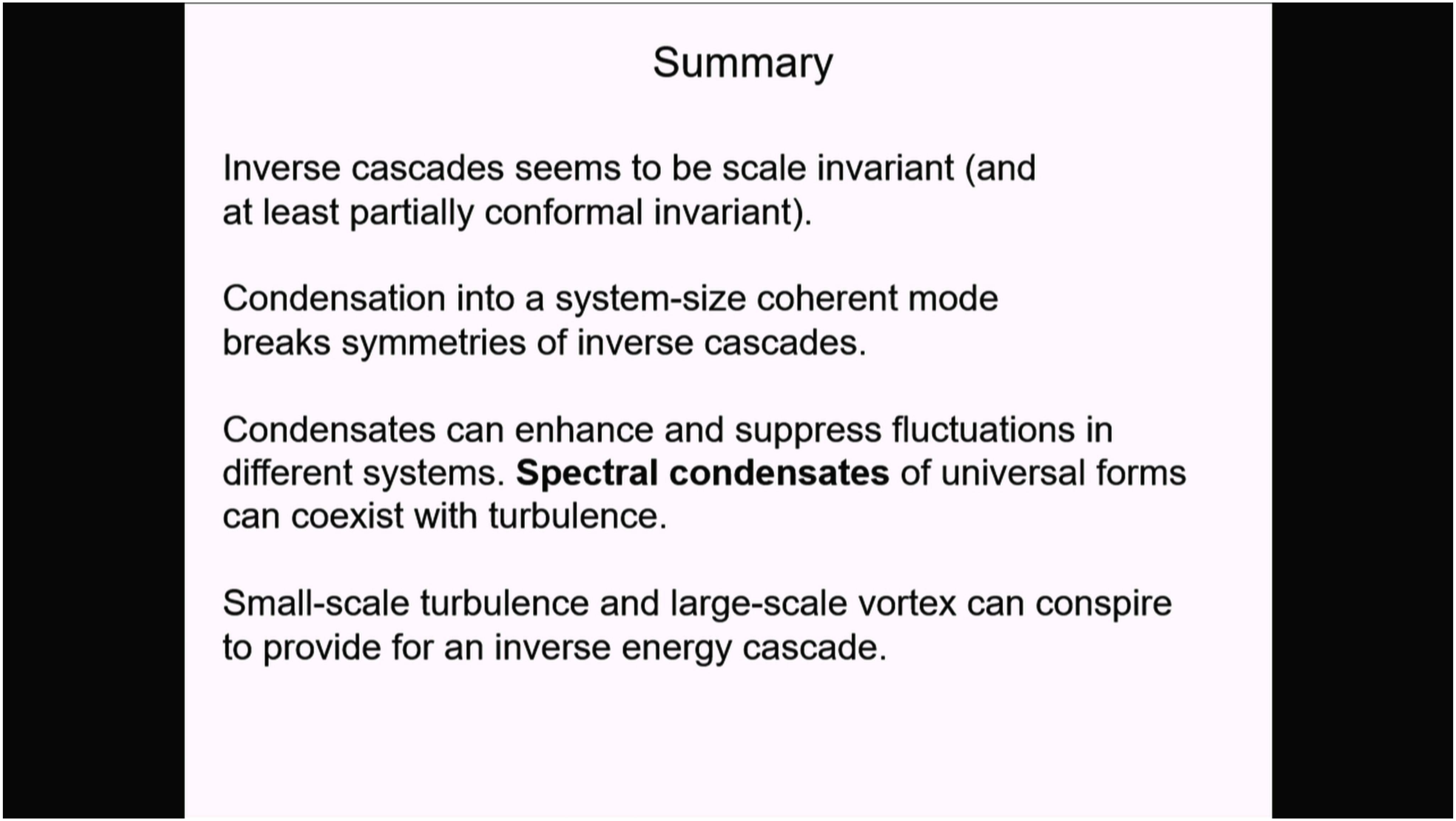
Three- to two-dimensional turbulence transition in the hurricane boundary layer
D. Byrne and A. Zhang, 2014



A transition from 3d to 2d turbulence from in-situ aircraft measurements in the hurricane boundary layer

Third order structure function of horizontal velocities for different flight-leg heights in hurricane A) Isabel and B) Fabian.

These results represent the first measurement of the 2D upscale energy flux in the atmosphere and also the first to characterize the transition from 3D to 2D. It is shown that the large-scale parent vortex may gain energy directly from small-scales in tropical cyclones.



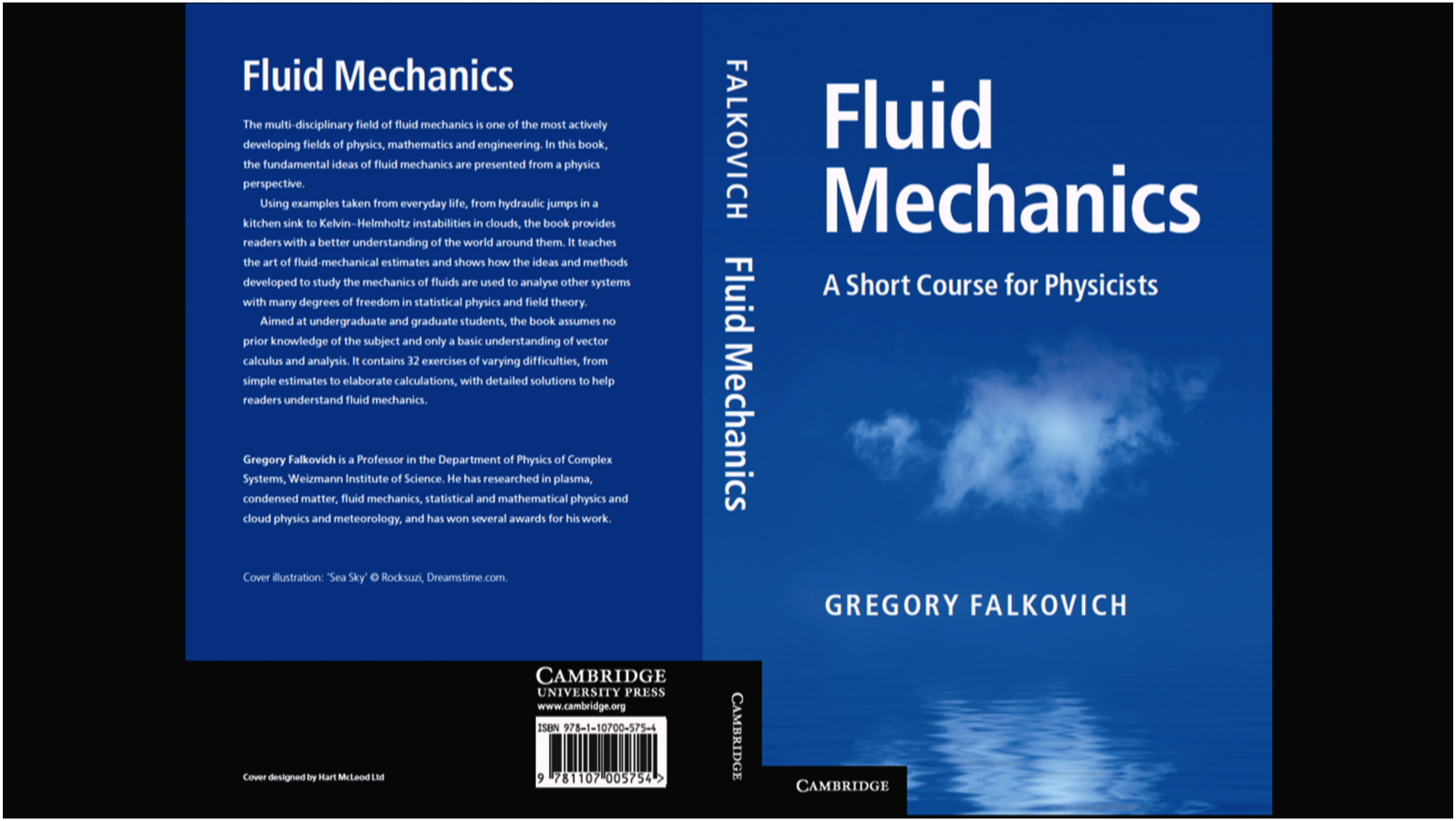
Summary

Inverse cascades seems to be scale invariant (and at least partially conformal invariant).

Condensation into a system-size coherent mode breaks symmetries of inverse cascades.

Condensates can enhance and suppress fluctuations in different systems. **Spectral condensates** of universal forms can coexist with turbulence.

Small-scale turbulence and large-scale vortex can conspire to provide for an inverse energy cascade.



Fluid Mechanics

The multi-disciplinary field of fluid mechanics is one of the most actively developing fields of physics, mathematics and engineering. In this book, the fundamental ideas of fluid mechanics are presented from a physics perspective.

Using examples taken from everyday life, from hydraulic jumps in a kitchen sink to Kelvin–Helmholtz instabilities in clouds, the book provides readers with a better understanding of the world around them. It teaches the art of fluid-mechanical estimates and shows how the ideas and methods developed to study the mechanics of fluids are used to analyse other systems with many degrees of freedom in statistical physics and field theory.

Aimed at undergraduate and graduate students, the book assumes no prior knowledge of the subject and only a basic understanding of vector calculus and analysis. It contains 32 exercises of varying difficulties, from simple estimates to elaborate calculations, with detailed solutions to help readers understand fluid mechanics.

Gregory Falkovich is a Professor in the Department of Physics of Complex Systems, Weizmann Institute of Science. He has researched in plasma, condensed matter, fluid mechanics, statistical and mathematical physics and cloud physics and meteorology, and has won several awards for his work.

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Fluid Mechanics

A Short Course for Physicists

GREGORY FALKOVICH

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