

Title: Modelling Surface Driven Flows in the Ocean

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Abstract: Buoyancy driven flows at the top of the ocean or bottom of the atmosphere are inherently different from their interior dynamics. One idealized model that has recently become very popular to idealize these surface flows with strong rotation is Surface Quasi-Geostrophic (SQG) dynamics. This model is appropriate for large-scale dynamics and assumes the motion is in near geostrophic and hydrostatic balance. Many of the numerical simulations of SQG have shown that vortices are frequently generated at very small scales that are well beyond the SQG limits. In this talk we examine the dynamics of a rotating three-dimensional elliptic vortex in both the SQG model and a more general and much more complicated primitive equation model. In order to compute high resolution solutions to the three dimensional primitive equations we make use of Sharcnet resources. We find that in the case of strong rotation (small Rossby number) we confirm the predictions from SQG. With weaker rotation (moderate Rossby number) we see the non-SQG effects that arise and find that the regime where SQG can be appropriate can be very limited. We conclude that some of the predictions that arise from the SQG model might not be very accurate in idealizing geophysical flows at the surface.

## Introduction

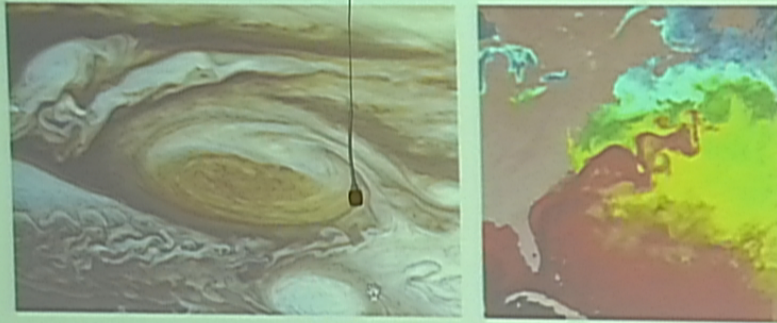


Figure : Left: Jupiter's Red Spot (Image from universetoday.com). Right: Sea surface Temperature of Gulf Stream (Image from the NC State Climate Office)



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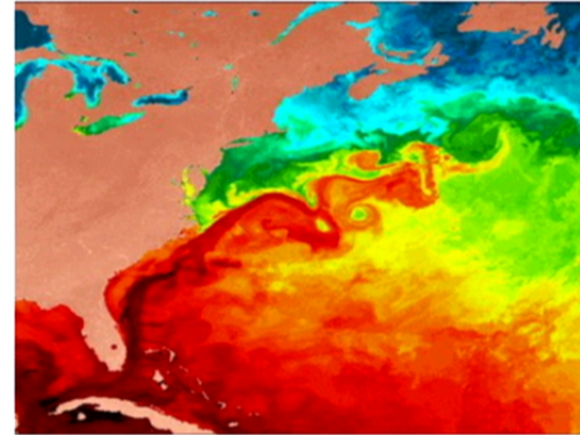


Figure : Left: Jupiter's Red Spot (Image from universetoday.com). Right: Sea Surface Temperature of Gulf Stream (Image from the NC State Climate Office)

- Large scale oceanic flows are rather slow
- Rotation of the Earth is important (Coriolis force)
- Quasi-Geostrophy (QG) can describe large scale motions very well
- Surface Quasi-Geostrophy (SQG) considers a scenario where all motion is driven by surface temperature

## Mathematical Models for the Oceans

- Non-dimensionalized Primitive equations (PE):

$$Ro \left( \frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} \right) + \mathbf{f} \times \mathbf{u} = -\frac{1}{\rho_0} \nabla p' + b' \mathbf{e}_z,$$

$$\nabla \cdot \mathbf{u} = 0,$$

$$Ro \left( \frac{\partial b'}{\partial t} + \mathbf{u} \cdot \nabla b' \right) + \left( \frac{L}{L_d} \right)^2 w = 0,$$

where  $\mathbf{u} = (u, v, w)$ ,  $Ro = \frac{U}{fL}$ ,  $L_d = \frac{NH}{f}$ ,  $N = \text{const.}$



## Mathematical Models for the Oceans

- Assumptions:

- $Ro = \frac{\text{advection}}{\text{rotation}} = \frac{U}{fL} \ll 1$
- $L \sim L_d \approx 100 \text{ km}$
- $L \gg H$

- Asymptotic expansion yields, at first order,

$$\frac{\partial q}{\partial t} + \mathbf{u} \cdot \nabla q = 0, \text{ for } -H < z < 0$$

$$\frac{\partial b'}{\partial t} + \mathbf{u} \cdot \nabla b' = 0 \text{ at } z = 0, -H$$

where  $q = \nabla^2 \psi + \left(\frac{f}{N}\right)^2 \frac{\partial^2 \psi}{\partial z^2}$  is the QG potential vorticity

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- Streamfunction,  $\psi$ , gives

$$u = -\frac{\partial \psi}{\partial y}, \quad v = \frac{\partial \psi}{\partial x}, \quad b' = f \frac{\partial \psi}{\partial z}$$



## Mathematical Models for the Oceans

- For SQG, set  $q = \nabla^2 \psi + \left(\frac{f}{N}\right)^2 \frac{\partial^2 \psi}{\partial z^2} = 0$
- At  $z = 0$ , take  $b' = b^t$  as given
- Find analytic solution in Fourier space,

$$\hat{\psi} = \frac{1}{NK} \frac{\cosh\left(\frac{NK}{f}(z + H)\right)}{\sinh\left(\frac{NKH}{f}\right)} \hat{b}^t,$$

where  $K = \sqrt{k^2 + l^2}$ ,  $(k, l)$  are horizontal wavenumbers



# Numerical Methods

- Two numerical models
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  - QG3 (G.R. Flierl, priv. comm.) for SQG dynamics
  - SPINS (Spectral Parallel Incompressible Navier-Stokes Solver) for the primitive equations (Subich et al 2013)



## Numerical Methods

	QG3	SPINS
Grid Size	$512^2$	$512^3$
Filter	Radially in spectral space	Applied in each direction
Parallel?	No	Yes (64 CPUs on Sharcnet)
Runtime	~ 10 minutes	~ 3 – 5 days!

