Title: From here to eternity: Global warming in geologic time

Date: Dec 06, 2006 07:00 PM

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Abstract: Using results from models of the atmosphere/ocean/sediment carbon cycle, the impacts of fossil-fuel CO2 release will be examined Â-including the effect on climate many thousands of years into the future, rather than for just a few centuries as commonly claimed. Prof. Archer will explain how aspects of the Earth system, such as the growth or melting of the great ice sheets, the thawing of permafrost, and the release of methane from the methane hydrate deposits in the deep ocean, take thousands of years to respond to a change in climate. The duration of our potential climate adventure is comparable to the pacing of climate changes in the past, which enables us to use the geologic record of past climate changes to predict the trajectory of global warming into the deep future. In particular, the record of sea level variations in the past suggests that the ultimate sea level response to fossil fuel CO2 use could be 10 to 100 times higher than the Intergovernmental Panel on Climate Change (IPCC) forecast for the year 2100. <kw>models, greenhouse gas, temperature forecast, medieval warm, little ice age, Greenland, Heinrich Events, fossil fuel, Climber Model Hysteresis, Ganopolski, Buffett, methane hydrates, Palaeocene, Eocene, Thermal Maximum Event </kw>

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# From Here to Eternity

David Archer University of Chicago

Global Warming in Geologic Time

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LONDON, EDINBURGH, AND DUBLIN

#### PHILOSOPHICAL MAGAZINE

AND

#### JOURNAL OF SCIENCE.

[FIFTH SERIES.]

APRIL 1896.

XXXI. On the Influence of Carbonic Acid in the Air upon the Temperature of the Ground. By Prof. SVANTE ARRHENIUS \*.

I. Introduction: Observations of Langley on Atmospherical Absorption.

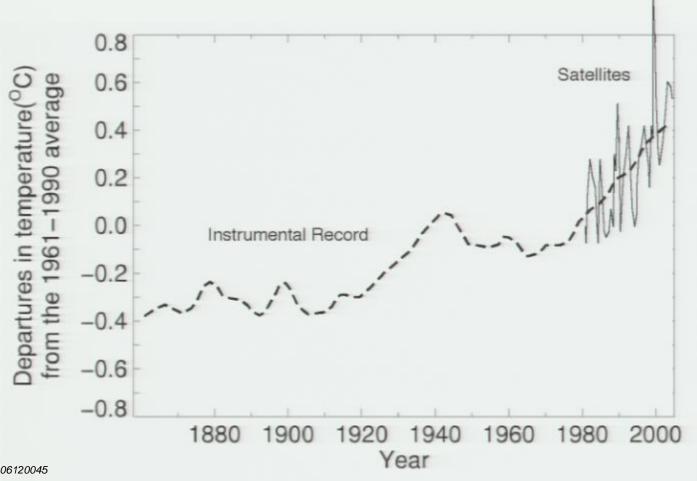
GREAT deal has been written on the influence of A the absorption of the atmosphere upon the climate. Tyndail t in particular has pointed out the enormous importance of this question. To him it was chiefly the diurnal and annual variations of the temperature that were lessened by this circumstance. Another side of the question, that has long attracted the attention of physicists, is this: Is the mean temperature of the ground in any way influenced by the presence of heat-absorbing gases in the atmosphere? Fourier; maintained that the atmosphere acts like the glass of a hothouse, because it lets through the light rays of the sun but retains the dark rays from the ground. This idea was elaborated by Pouillet §; and Langley was by some of his researches led to the view, that "the temperature of the earth under direct sunshine, even though our atmosphere Pirsa: 06120045 were present as now, would probably fall to  $-200^{\circ}$  C., if that atmosphere did not possess the quality of selective

## Arrhenius, 1896

Predicted that doubling CO<sub>2</sub> would raise global temperature by 4-6°C.

Now, of course, we predict 3-5°C (a century of progress!).

## The Earth is Warming

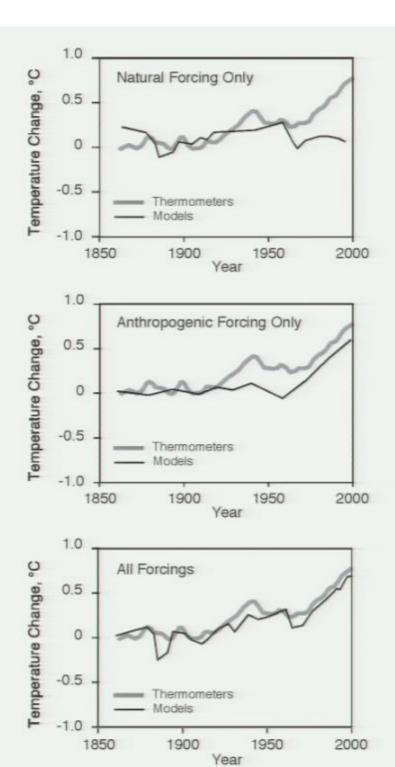


Satellites and weather stations now agree.

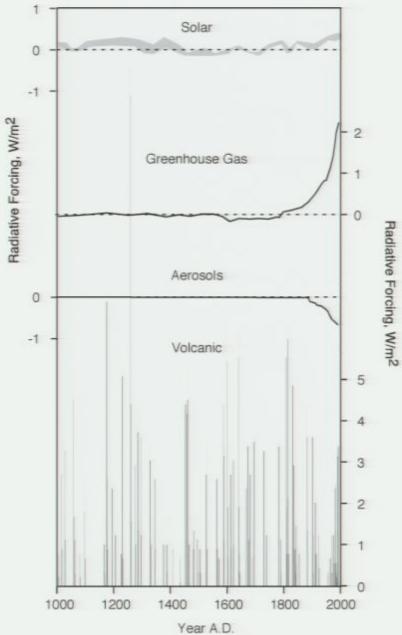


Ice sheets and glaciers are melting around the world

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Climate models
can explain the
recent temperature
changes, but only
by taking into account
both human and
natural climate
forcings

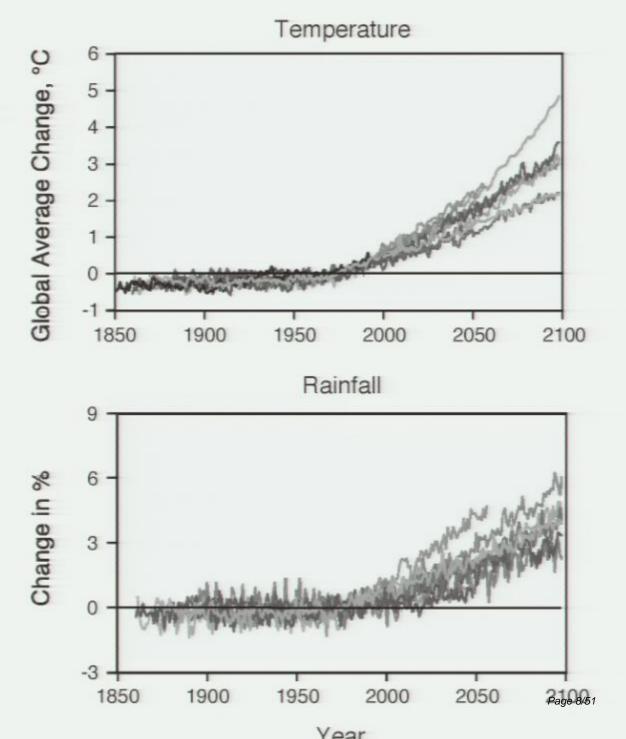


Only greenhouse gas forcing looks like the recent temperature rise.

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Temperature Forecast: 2-4° C warming by 2100

General increase in precipitation in a warmer world.

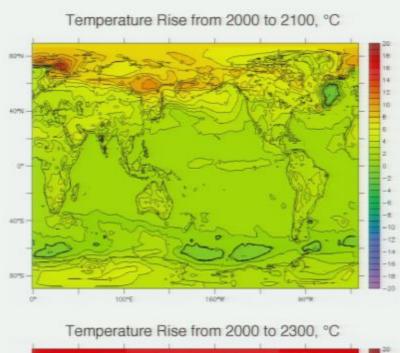


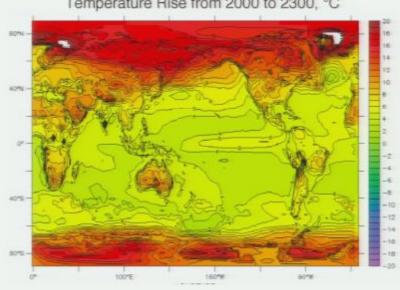
Temperature changes from 1750 (natural)

For the year 2100

And 2300

Warming is most intense in high latitudes





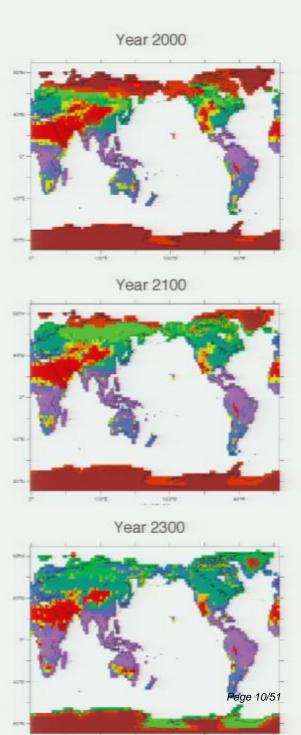
#### Land Biomes

Climate determines the landscape of the Earth.

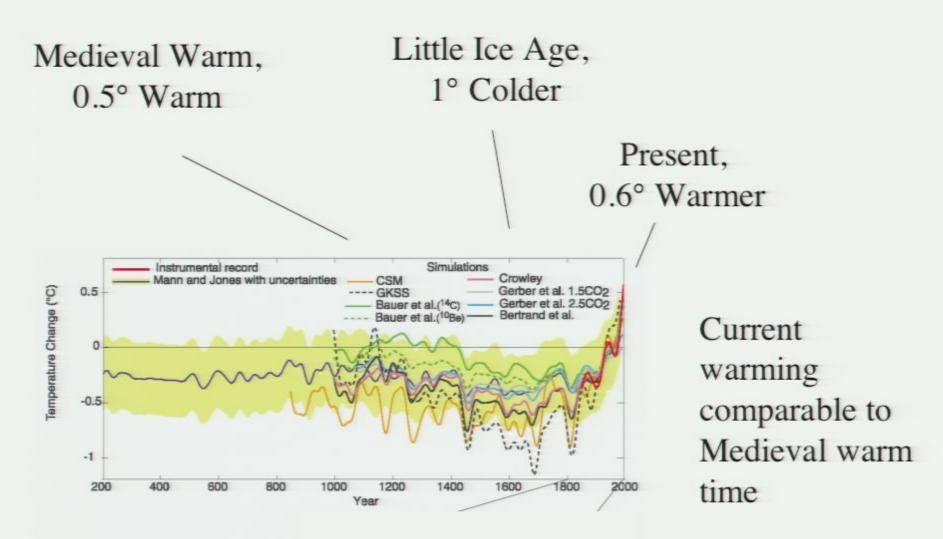
Tundra is lost by 2300.

Tropical Evergreen Tropical Deciduous Temperate Evergreen Broadleaf Gemperate Evergreen Conifer Temperate Deciduous Forest Boreal Evergreen Forest Boreal Deciduous Forest Mixed Forest Savana Grassland Dense Shrubland Open shrubland Tundra Desert

Polar Desert



## The Past

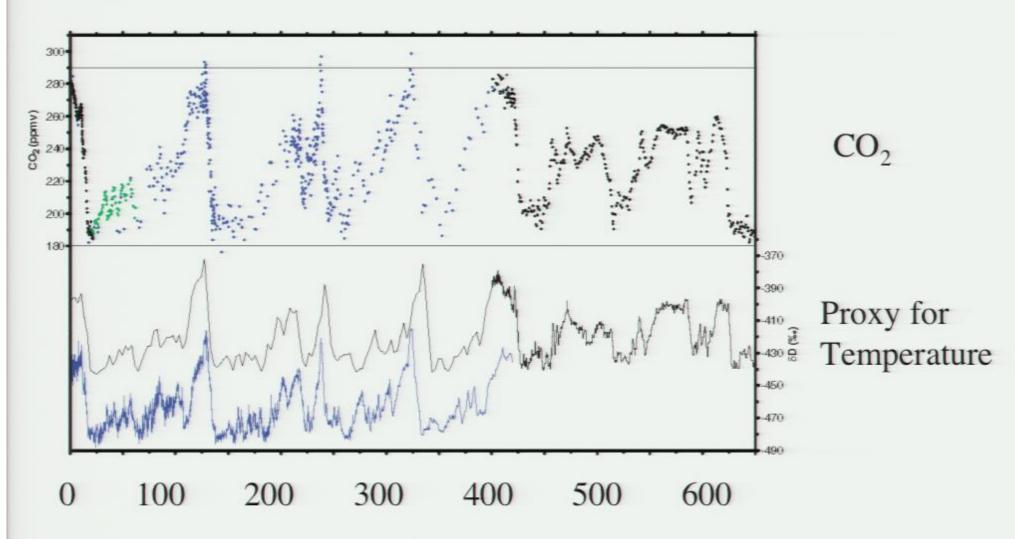


## Lessons from the past

Present-day climate is comparable to medieval warm. Real, noticable, but not globally catastrophic.

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## Glacial cycles in ice sheets and CO<sub>2</sub>



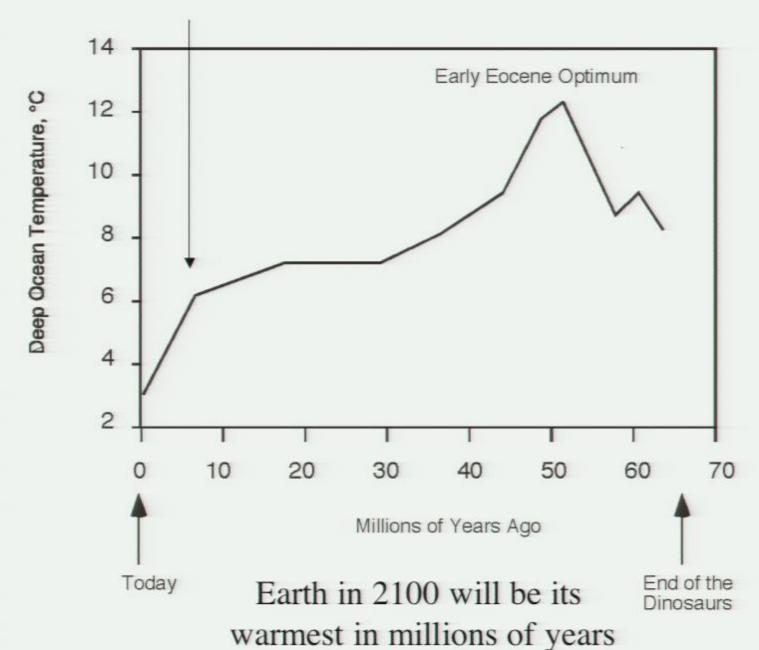
## Lessons from the past

Present-day climate is comparable to medieval warm. Real, noticable, but not globally catastrophic.

Temperature change in the coming century is more comparable to the end of the last glacial time.

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#### 3° Warmer 5 Million Years Ago



## Lessons from the past

Present-day climate is comparable to medieval warm. Real, noticable, but not globally catastrophic.

Temperature change in the coming century is more comparable to the end of the last glacial time, but to a warmth unlike any in millions of years.

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## Lessons from the past

Present-day climate is comparable to medieval warm. Real, noticable, but not globally catastrophic.

Temperature change in the coming century is more comparable to the end of the last glacial time, but to a warmth unlike any in millions of years.

Another note from the past: climate changes can be abrupt and unpredictable, rather than the smooth changes that the models predict.

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## The Future

## Fate of fossil fuel CO<sub>2</sub>

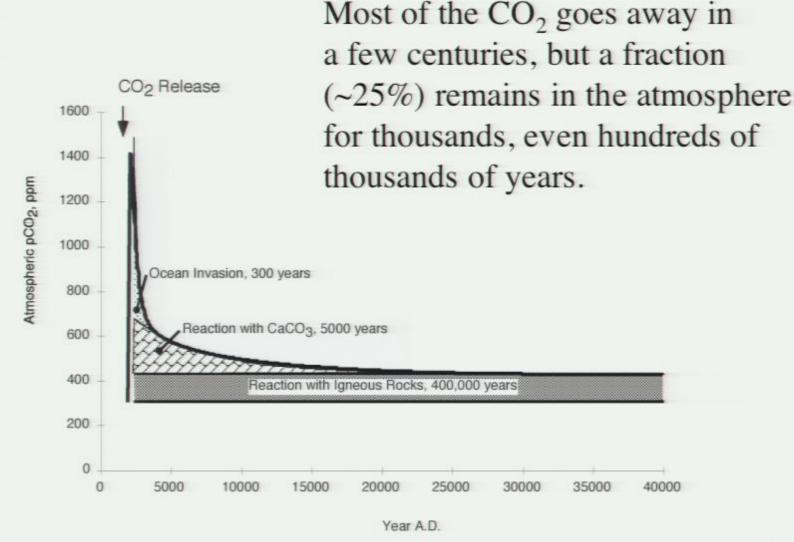
Dissolves in the ocean (centuries)

Uptake / release from terrestrial biosphere (centuries)

Neutralization by CaCO<sub>3</sub> (5-10 kyr)

Lithification by weathering of silicate rocks (400 kyr)

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Archer, 1997 and 2005

## IPCC 2001 got this wrong

Table 1: Examples of greenhouse gases that are affected by human activities. [Based upon Chapter 3 and Table 4.1]

	CO <sub>2</sub> (Carbon Dioxide)	CH <sub>4</sub> (Methane)	N <sub>2</sub> O (Nitrous Oxide)	CFC-11 (Chlorofluoro -carbon-11)	HFC-23 (Hydrofluoro -carbon-23)	CF <sub>4</sub> (Perfluoro- methane)
Pre-industrial concentration	about 280 ppm	about 700 ppb	about 270 ppb	zero	zero	40 ppt
Concentration in 1998	365 ppm	1745 ppb	314 ppb	268 ppt	14 ppt	80 ppt
Rate of concentration change <sup>b</sup>	1.5 ppm/yr <sup>a</sup>	7.0 ppb/yr <sup>a</sup>	0.8 ppb/yr	-1.4 ppt/yr	0.55 ppt/yr	1 ppt/yr
Atmospheric lifetime	5 to 200 yr °	12 yr <sup>d</sup>	114 yr <sup>d</sup>	45 yr	260 yr	>50,000 yr

<sup>\*</sup> Rate has fluctuated between 0.9 ppm/yr and 2.8 ppm/yr for CO<sub>2</sub> and between 0 and 13 ppb/yr for CH<sub>4</sub> over the period 1990 to 1999.

This lifetime has been defined as an "adjustment time" that takes into account the indirect effect of the gas on its own residence time.



#### Since then, it's repeated everywhere

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<sup>&</sup>lt;sup>b</sup> Rate is calculated over the period 1990 to 1999.

<sup>&</sup>lt;sup>c</sup> No single lifetime can be defined for CO<sub>o</sub> because of the different rates of uptake by different removal processes.

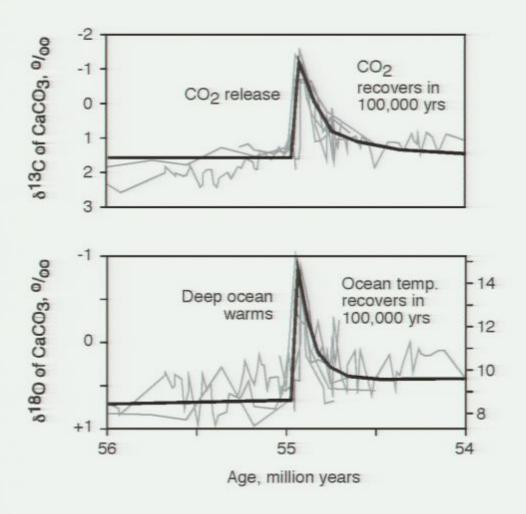
#### Airborne Fraction of a Large CO<sub>2</sub> Release

Archer 2005
Lenton 2006
Brovkin in prep.
Goodwin subm.
Ridgwell subm.
Tyrell subm.

Peak	1 kyr	10 kyr	
60%	33%	15%	
67-75%	14-16%	10-15%	
67%	57%	26%	
50%	40%		
50%	34%	12%	
70%	42%	21%	

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## Paleocene/Eocene Thermal Maximum Event 55 Myr Ago



A natural release of CO<sub>2</sub>, comparable to the potential fossil fuel release.

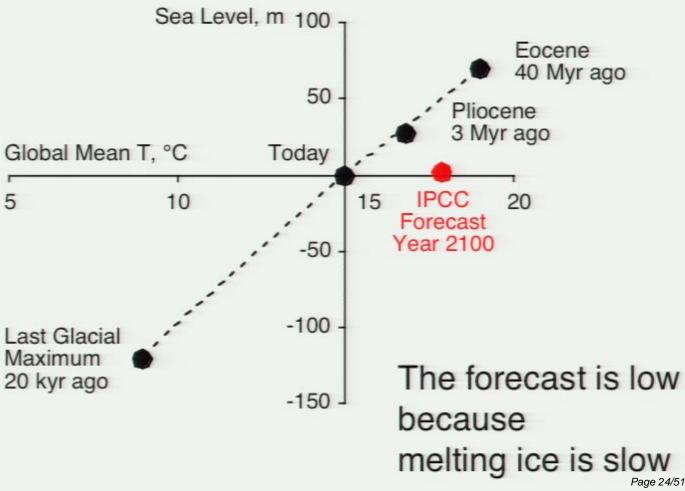
Warming, with a recovery that took 100,000 years.

## Long-Term Temperature Impact

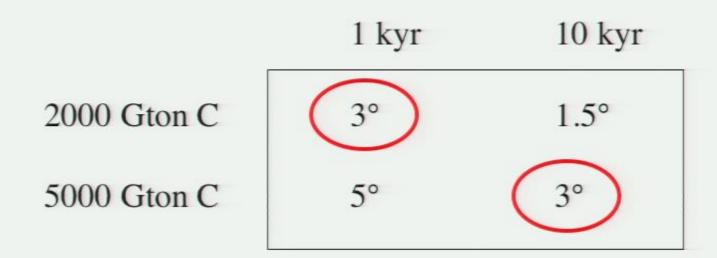
	1 kyr	10 kyr	
2000 Gton C	3°	1.5°	
5000 Gton C (all the coal)	5°	3°	

Assumes 3°C warming for doubling CO<sub>2</sub>

## Sea Level

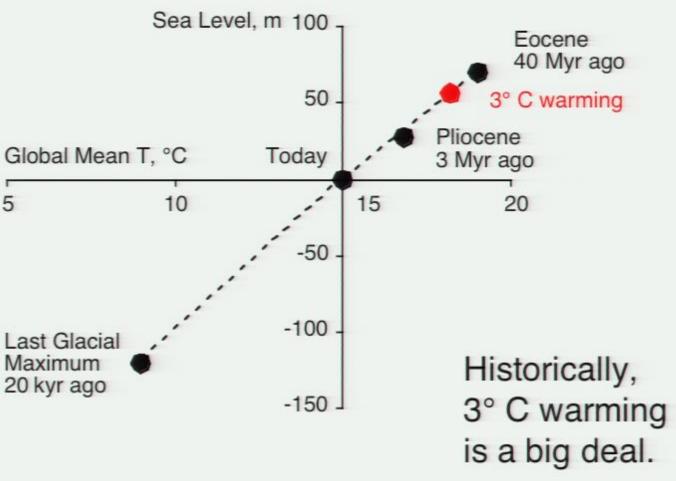


## Long-Term Temperature Impact

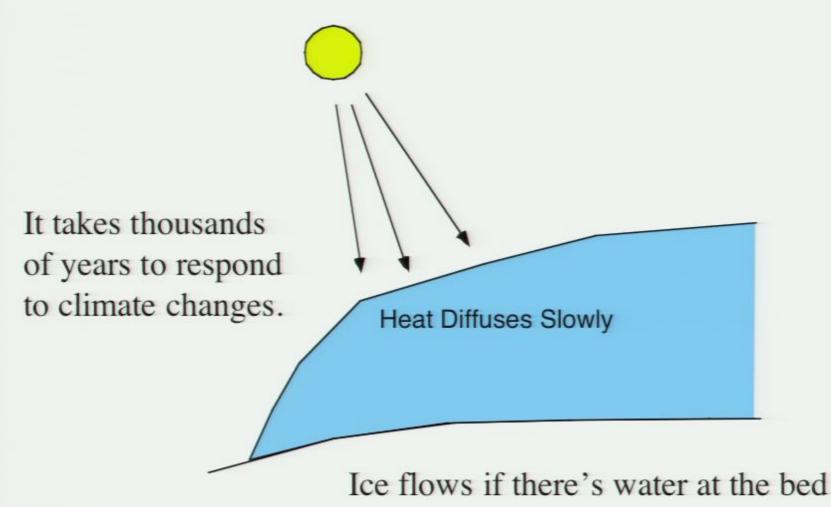


Assumes 3°C warming for doubling CO<sub>2</sub>

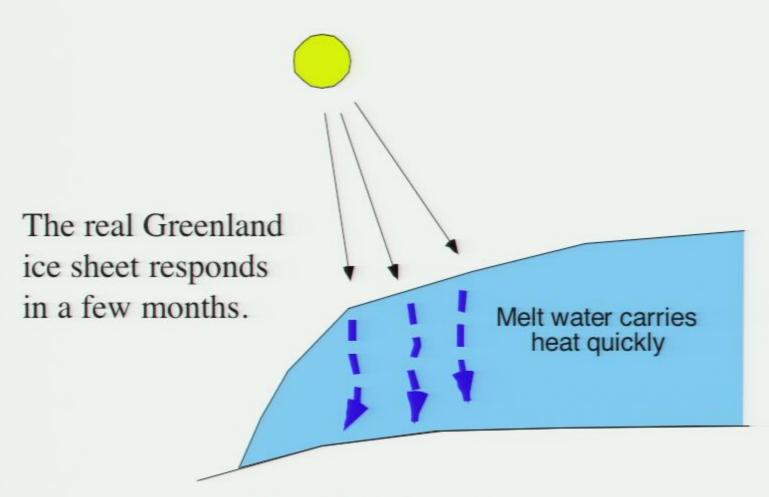
## Sea Level



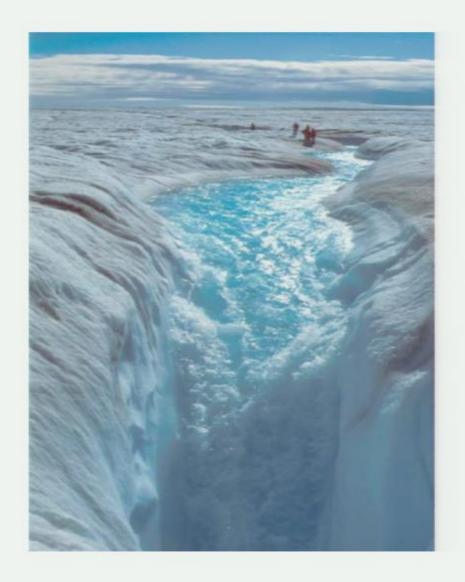
Ice sheet models are probably too sluggish.



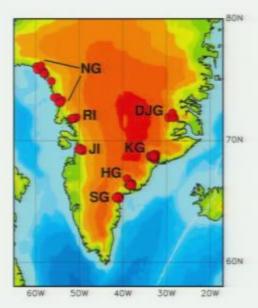
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We don't know how water gets through the ice



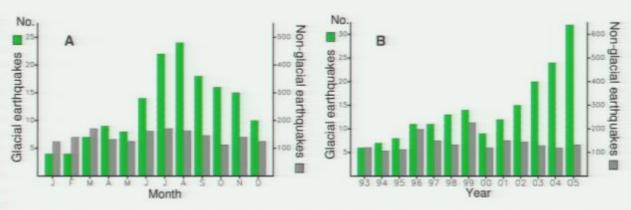
A moulin in Greenland where water submerges into the ice.



# Earthquakes under Greenland ice

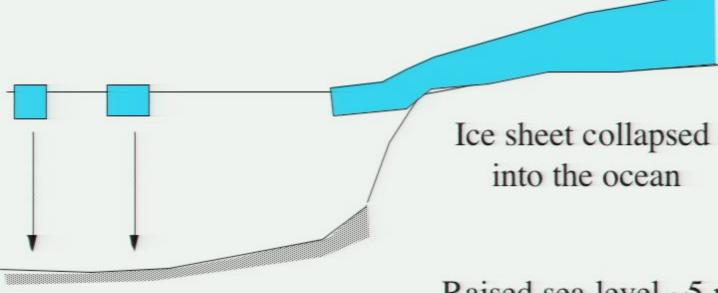
Fig. 1. Topographic map of southern Greenland and vicinity. The locations of 136 glacial earth-quakes defining seven groups are indicated with red circles: DJG, Daugaard Jensen Glacier (5 events); KG, Kangerdlugssuaq Glacier (61); HG, Helheim Glacier (26); SG, southeast Greenland glaciers (6); JJ, Jakobshavn Isbrae (11); RI, Rinks Isbrae (10); NG, northwest Greenland glaciers (17). Owing to the tight dustering of the earth-quakes, many of the individual symbols on the map overlap.

Fig. 2. (A) Histogram showing seasonality of glacial earthquakes on Greenland. Green bars show the number of detected Greenland glacial earthquakes in each month during the period 1993 to 2004. Gray bars show the number of earthquakes of similar magnitude detected elsewhere north of 45°N during



the same period. (B) Histogram showing the increasing number of Greenland glacial earthquakes (green bars) since at least 2002. No general increase in the detection of earthquakes north of 45°N (gray bars) is observed during this time period.

## Heinrich Events 30-70 kyr ago



Ice Rafted Debris (layers of rocks in ocean sediments) Raised sea level ~5 m in a few centuries. Could the Greenland ice sheet start doing this?

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## Meltwater Pulse 1A 19kyr ago

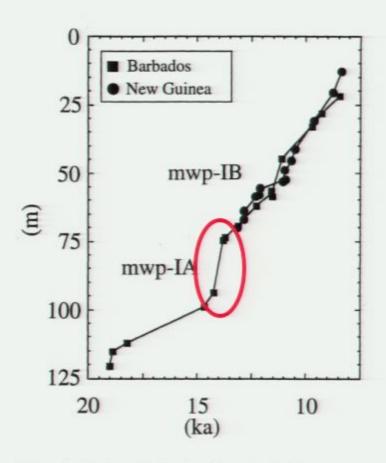
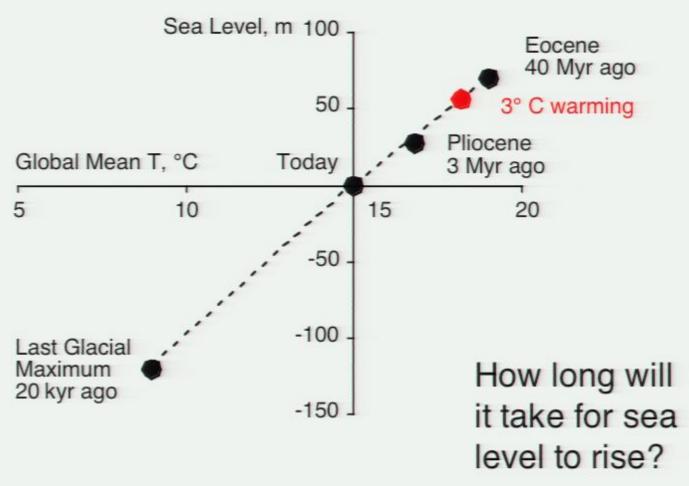


Figure 1. Coral records of sea level dated by U/Th from Barbados [Bard et al., 1990a, 1993] and New Guinea [Edwards et al., 1993]. Two periods of rapid rise of sea level are identified as mwp-IA and mwp-IB.

1.5 to 3 Greenlands in 1-5 centuries.

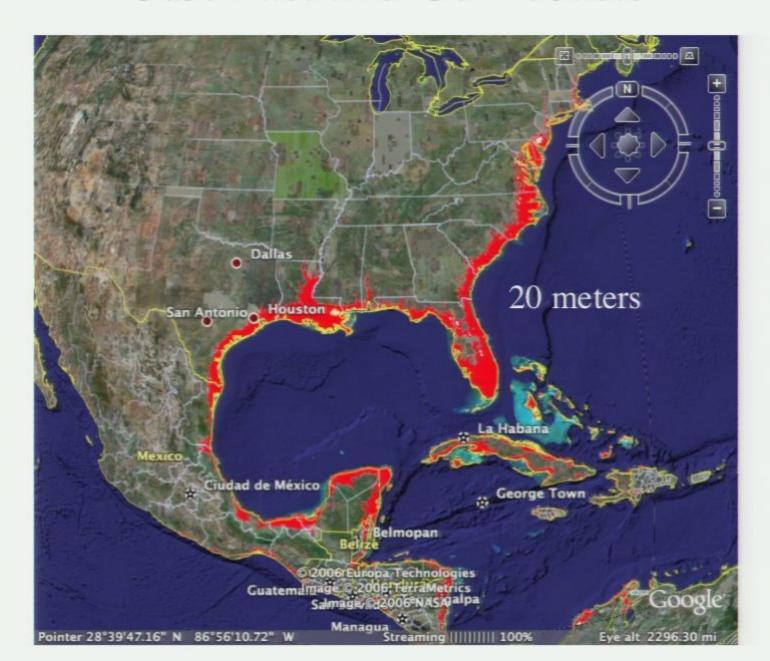
We're not even sure where this water came from.

## Sea Level

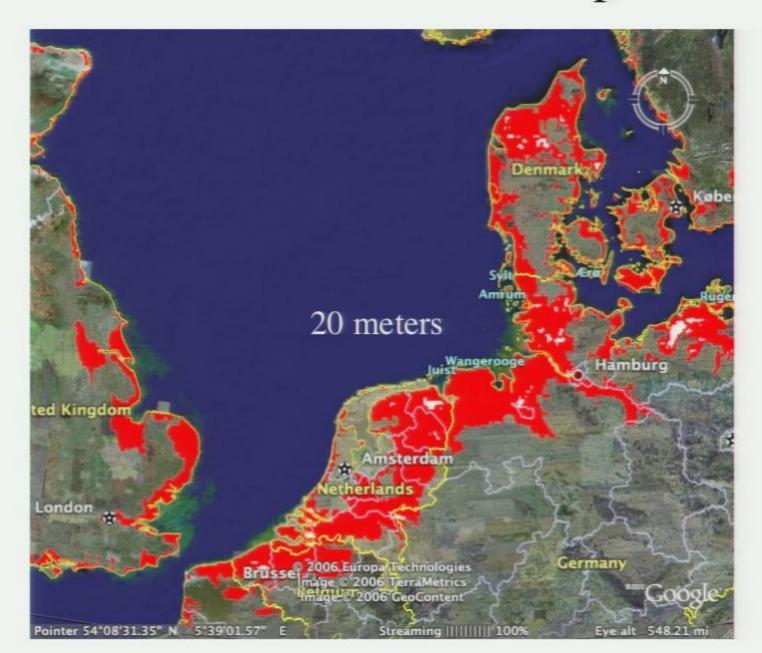


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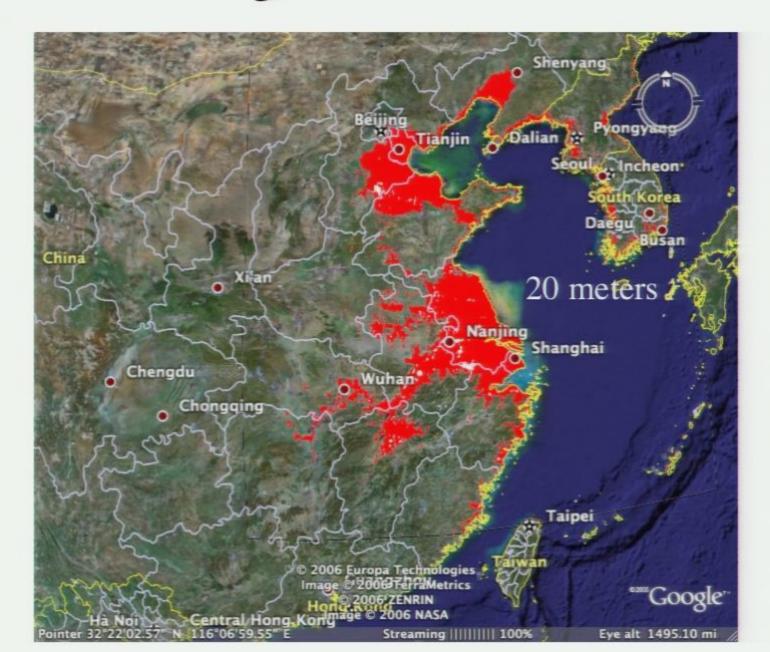
### U.S. East and Gulf coasts



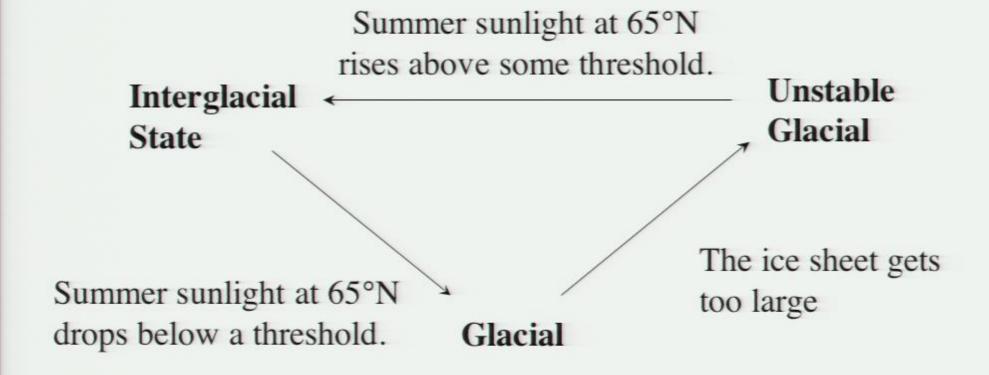
## Low Countries of Europe



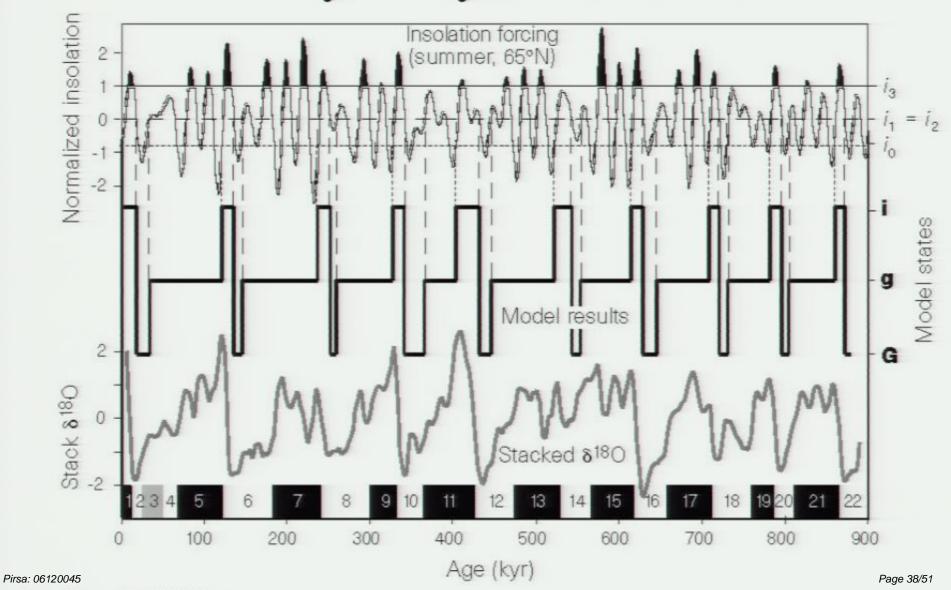
## Yangtze Delta, China



# Fossil Fuel Carbon and Ice Ages

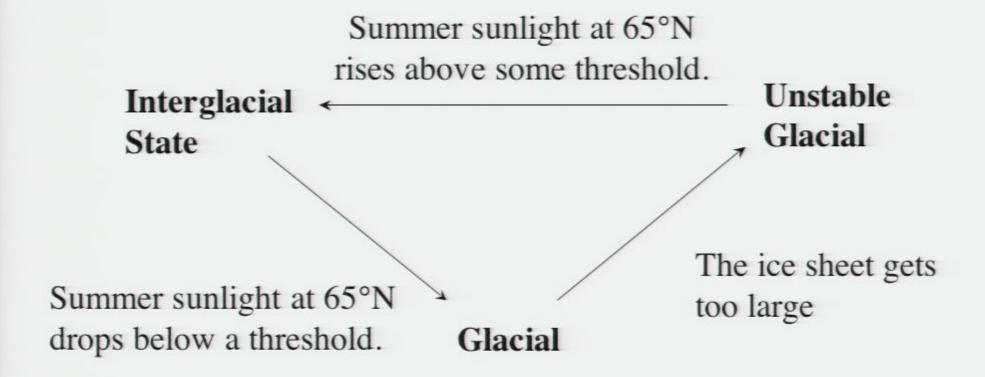


### 35 kyr sticky switch model



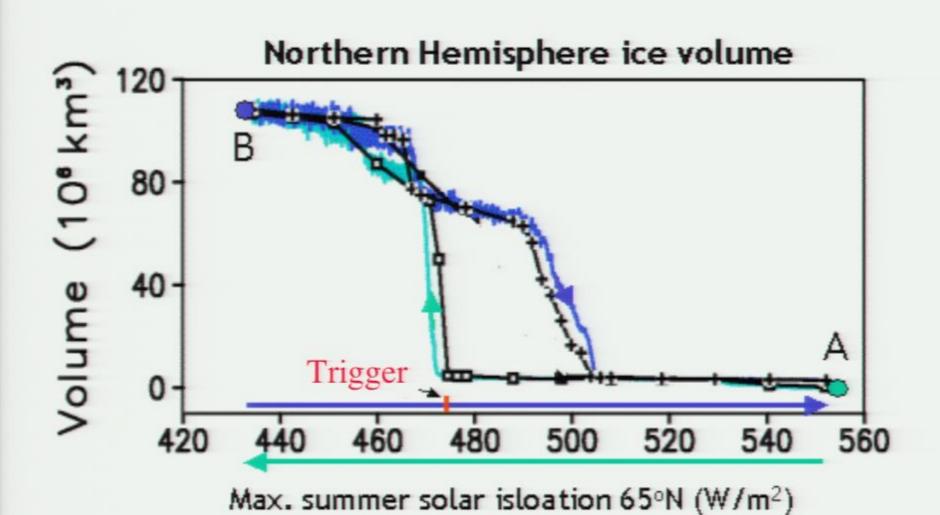
Paillard, 1998

# Fossil Fuel Carbon and Ice Ages



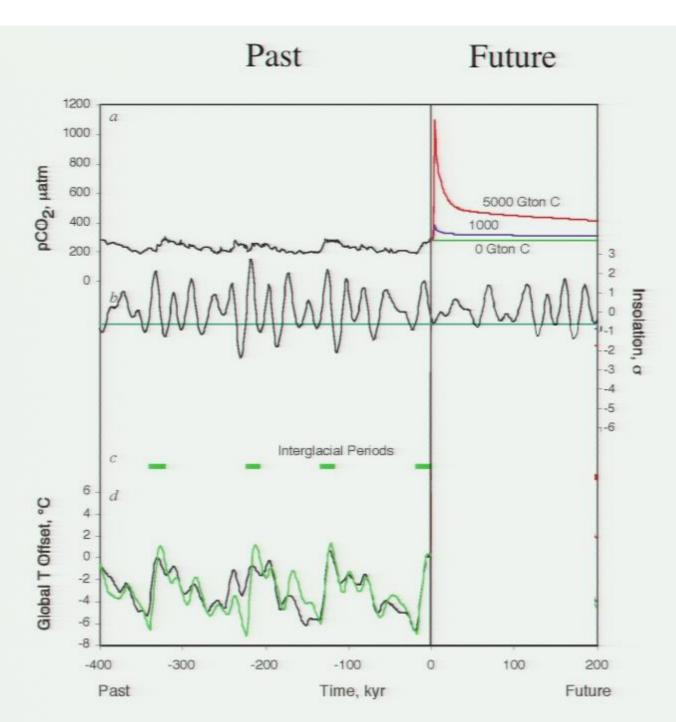
The Trigger

#### CLIMBER Model Hysteresis



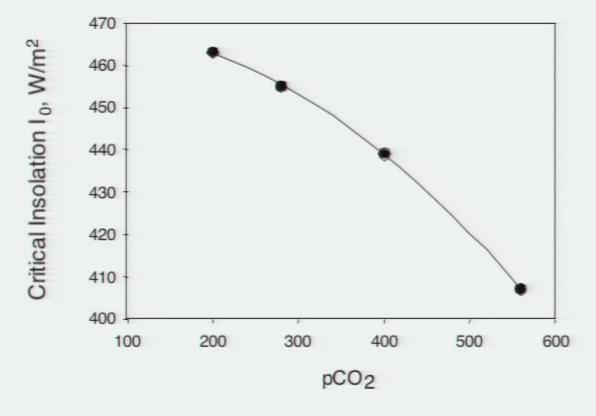
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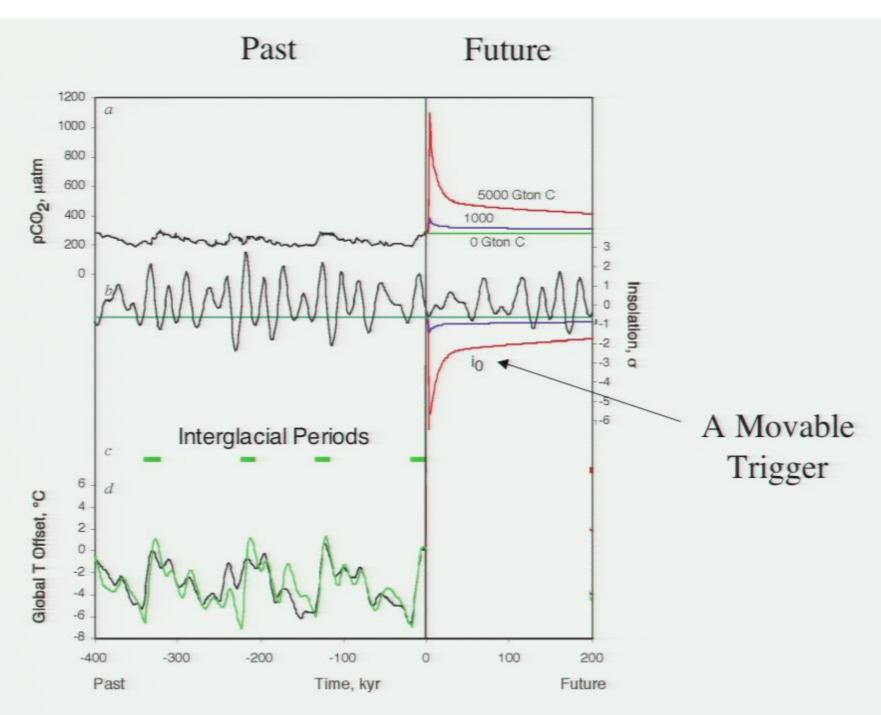


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Archer and Ganopolski, 2005

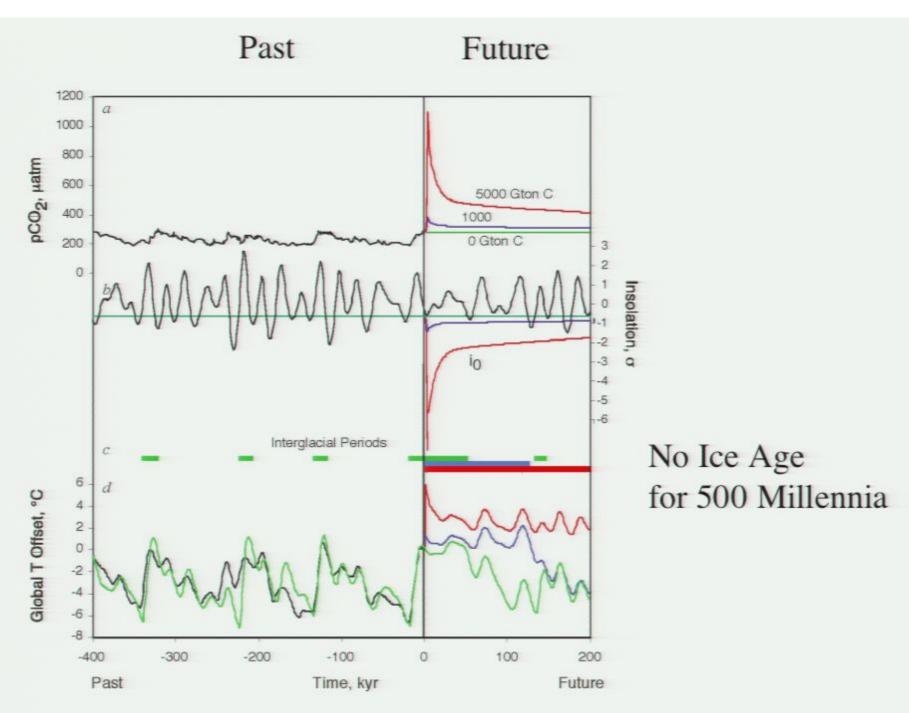


If CO<sub>2</sub> is higher, it takes a stronger sunlight change to trigger formation of an ice sheet.



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Archer and Ganopolski, 2005



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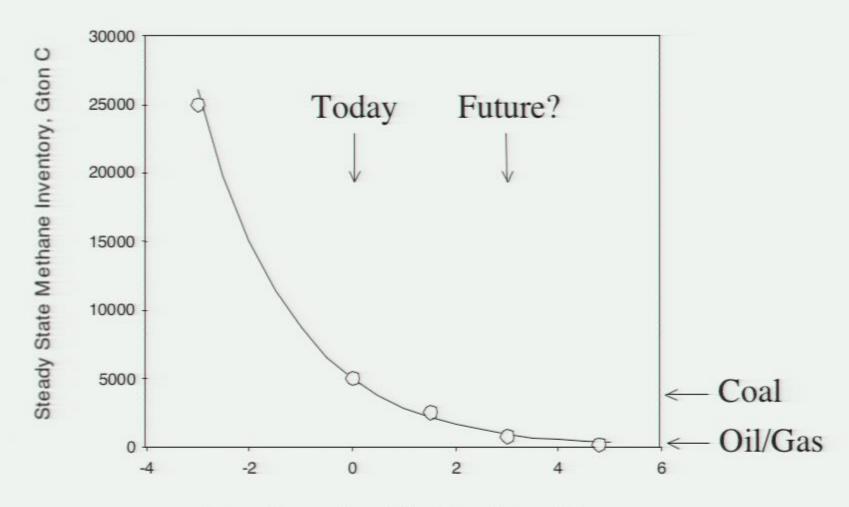
Archer and Ganopolski, 2005

# Methane Hydrates



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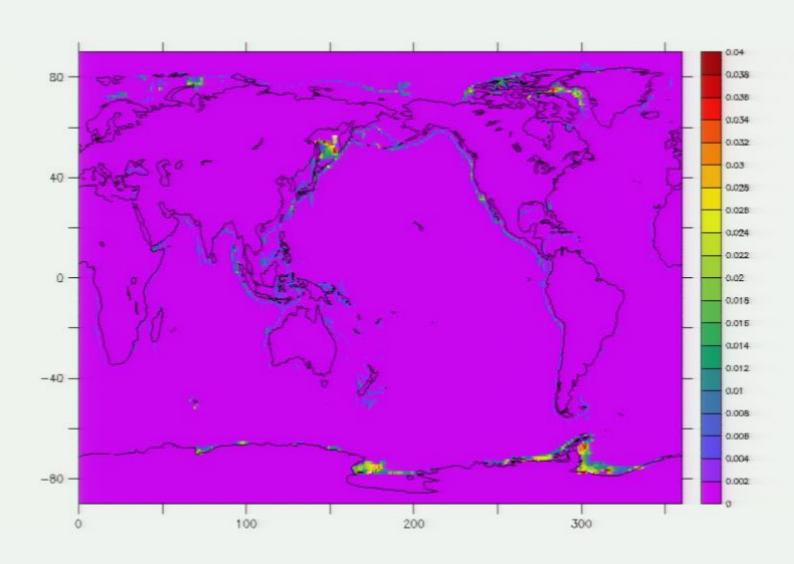
#### Methane Hydrates



Ocean Temperature, Offset from Present Day

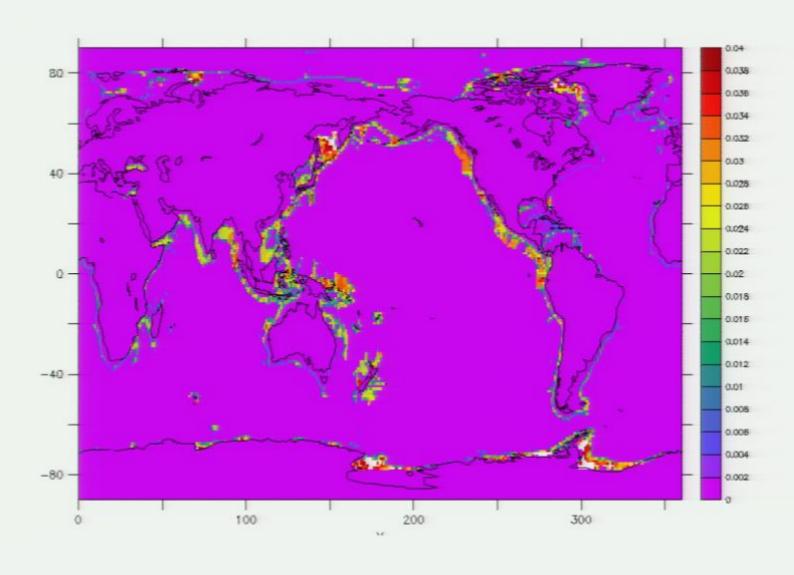
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### Bubble Volume, %



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## Bubble volume upon melting, %



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## Hydrates bottom line

Probably won't catastrophically blow out in the coming century, because it takes a long time to warm the deep ocean and into the sediment.

On time scales of millennia and longer, the hydrates could release carbon to match our fossil fuel carbon release.

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## From Here to Eternity

The long lifetime of nuclear waste matters to people.

Why would the long lifetime of global warming be any different?

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#### What to do?

We have already emitted about 300 billion metric tons of carbon from fossil fuels and deforestation.

We could ultimately emit about 700 billion metric tons and just avoid a "dangerous climate change of 2° C.

The 400 billion metrics tons we haven't released yet is about equivalent to the remaining oil and gas reservoirs. Just stop burning coal.

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