

Title: A Movable Trigger: Fossil fuel CO2 and the Next Glaciation

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Abstract:

A Movable Trigger

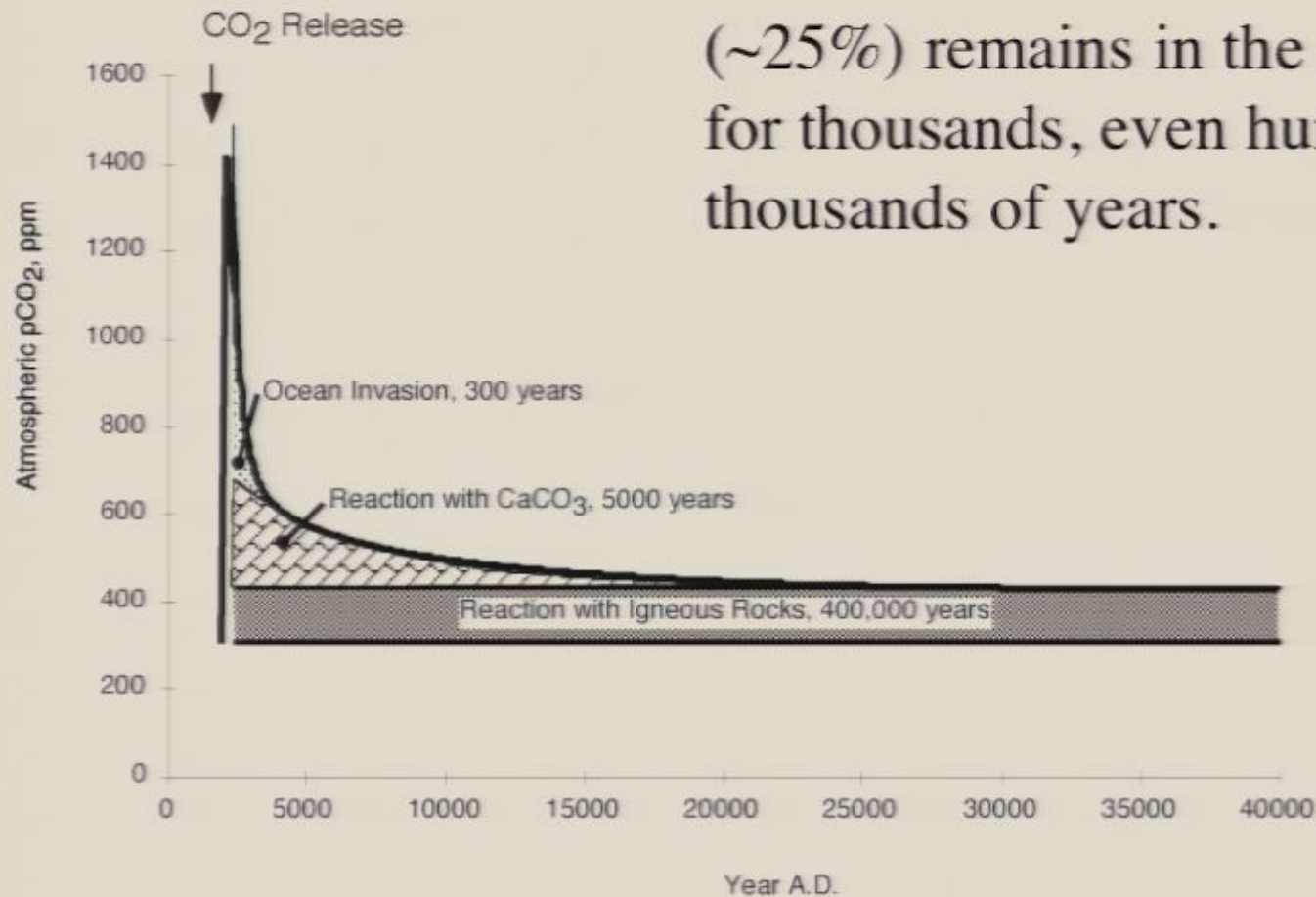
David Archer
University of Chicago

Fossil fuel CO₂
and the
Next Glaciation

I. Fate of fossil fuel CO₂

Dissolves in the ocean	(centuries)
Uptake / release from terrestrial biosphere	(centuries)
Neutralization by CaCO ₃	(5-10 kyr)
Lithification by weathering of silicate rocks	(400 kyr)

Most of the CO_2 goes away in a few centuries, but a fraction (~25%) remains in the atmosphere for thousands, even hundreds of thousands of years.



IPCC 2001 got this wrong

Table 1. Examples of greenhouse gases that are affected by human activities. [Based upon Chapter 3 and Table 4.1]

	CO ₂ (Carbon Dioxide)	CH ₄ (Methane)	N ₂ O (Nitrous Oxide)	CFC-11 (Chlorofluoro -carbon-11)	HFC-23 (Hydrofluoro -carbon-23)	CF ₄ (Perfluoro- methane)
Pre-industrial concentration	about 280 ppm	about 700 ppb	about 270 ppb	zero	zero	40 ppt
Concentration in 1998	365 ppm	1745 ppb	314 ppb	268 ppt	14 ppt	80 ppt
Rate of concentration change ^b	1.5 ppm/yr ^a	7.0 ppb/yr ^a	0.8 ppb/yr	-1.4 ppt/yr	0.55 ppt/yr	1 ppt/yr
Atmospheric lifetime	5 to 200 yr ^c	12 yr ^d	114 yr ^d	45 yr	260 yr	>50,000 yr

^a Rate has fluctuated between 0.9 ppm/yr and 2.8 ppm/yr for CO₂ and between 0 and 13 ppb/yr for CH₄ over the period 1990 to 1999.

^b Rate is calculated over the period 1990 to 1999.

^c No single lifetime can be defined for CO₂ because of the different rates of uptake by different removal processes.

^d This lifetime has been defined as an "adjustment time" that takes into account the indirect effect of the gas on its own residence time.

Since then, it's repeated everywhere

Airborne Fraction of a large CO₂ release

	Peak	1 kyr	10 kyr
Archer 2005	60%	33%	15%
Lenton 2006	67-75%	14-16%	10-15%
Brovkin in prep.	67%	57%	26%
Goodwin subm.	50%	40%	
Ridgwell subm.	50%	34%	12%
Tyrell subm.	70%	42%	21%

II. Orbits and Climate

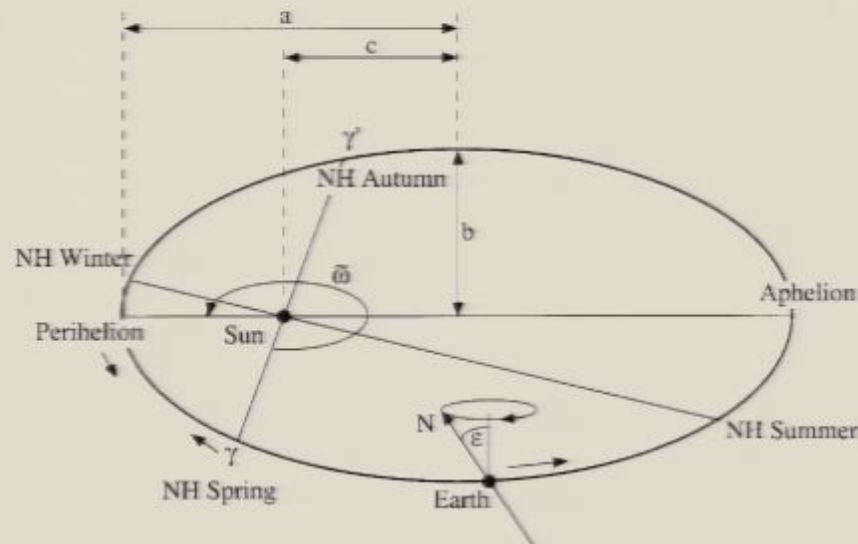
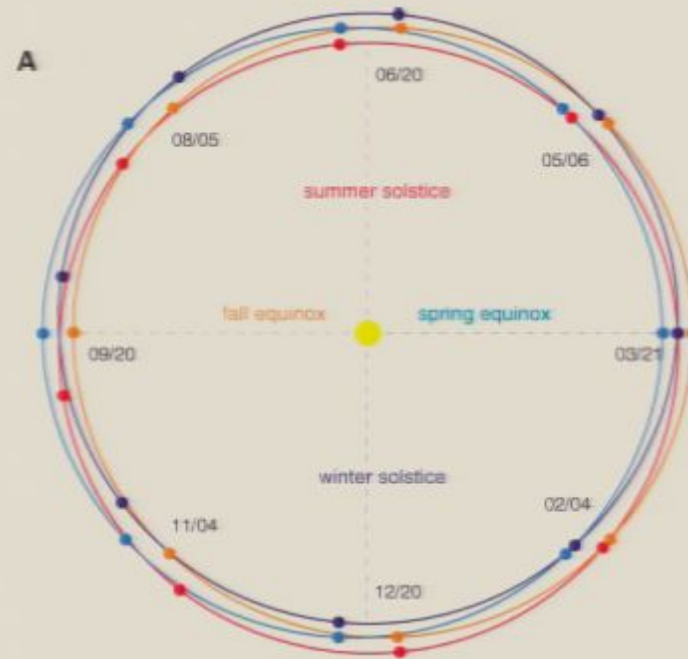
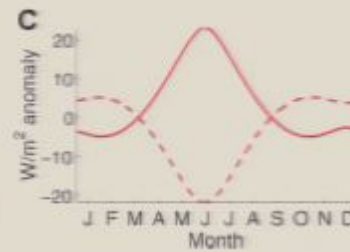
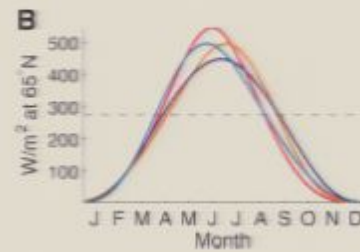


Figure 2. The orbital parameters of the Earth. Eccentricity e is defined as $e = c/a$, where a is the semimajor axis and c is the distance between the focus and the center of the ellipse. The semiminor axis b is then given by Pythagoras's theorem ($a^2 = b^2 + c^2$, which gives $b = a\sqrt{1 - e^2}$). The current eccentricity value is $e = 0.0167$, which means that the Earth's orbit is very close to a circle. The tilt of the Earth's axis with respect to the orbital plane is the obliquity ϵ (current value is $\epsilon = 23.44^\circ$). This tilt implies that the Earth equatorial plane intersects with its orbital plane, the intersection defining the $\gamma\gamma'$ line and the position of equinoxes and solstices. In the current configuration the Earth is closest to the Sun (perihelion) around January 3, just a few weeks after the Northern Hemisphere winter. This position, relative to the vernal equinox γ , is measured by the $\bar{\omega}$ angle.

Annual cycle of insolation at 65° N



Drawn to scale



Eccentricity

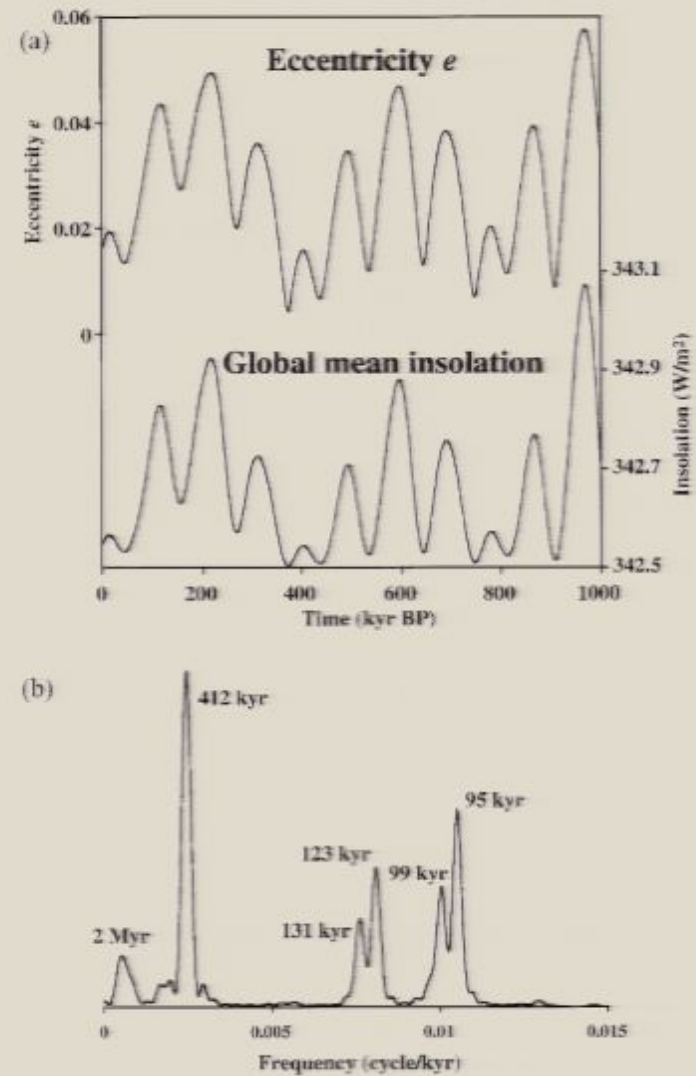


Figure 3. (a) Changes in eccentricity e for the last million years and its small effect on the global annual mean insolation received by the Earth (assuming a constant solar activity). (b) Spectral analysis of the eccentricity changes, revealing major periodicities at ~ 400 kyr and in the 100-kyr band (arbitrary vertical linear scale).

Obliquity

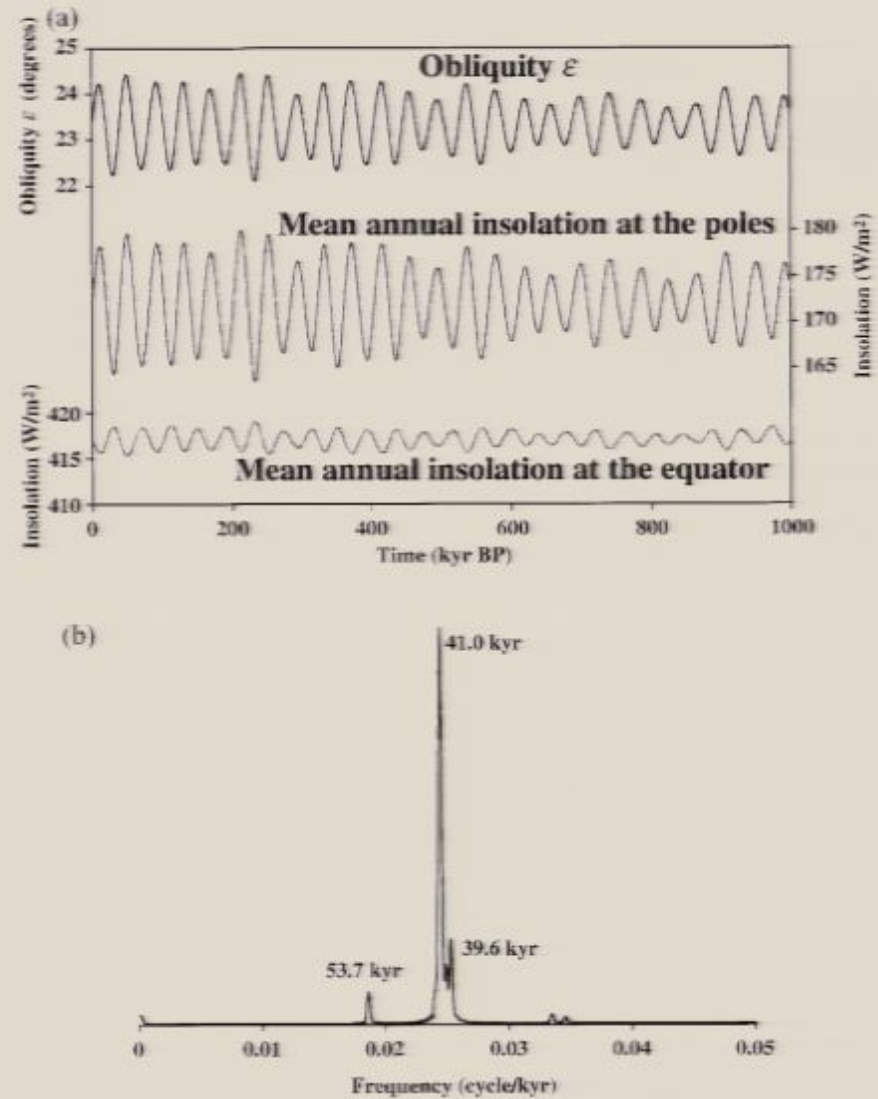


Figure 4. (a) Changes in obliquity ϵ for the last million years and the annual mean insolation received at the poles and at the equator. (b) Spectral analysis of the obliquity changes, revealing one major periodicity at 41 kyr (arbitrary vertical linear scale).

Obliquity stabilized by the moon

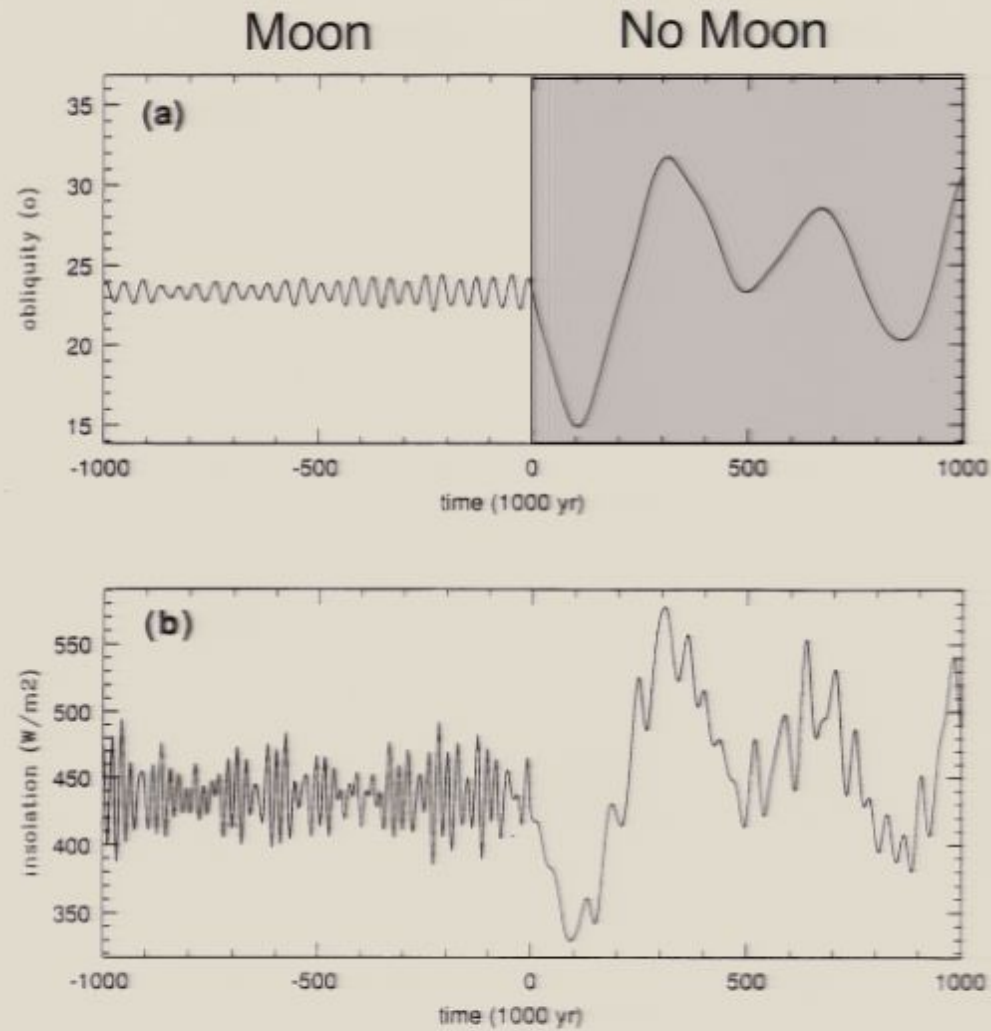


Fig. 11. Changes in obliquity (a) and insolation at 65N ($\lambda_d = 120$ deg) (b) resulting from the suppression at $t = 0$ of the Moon. The Moon is present from -1Myr to 0, and absent from 0 to +1Myr

Precession

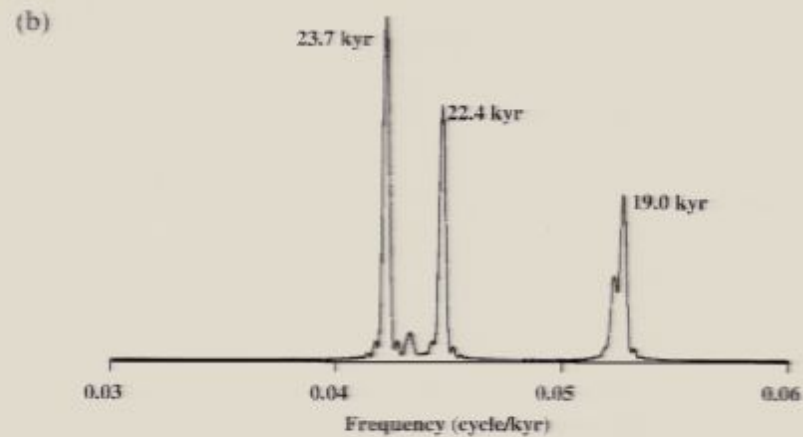
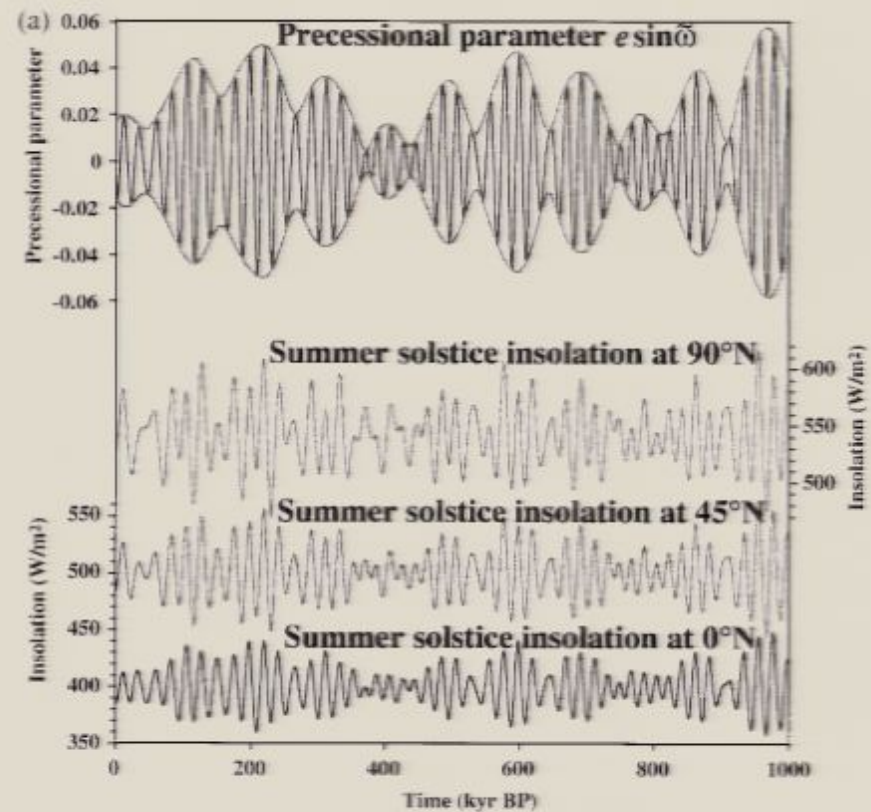
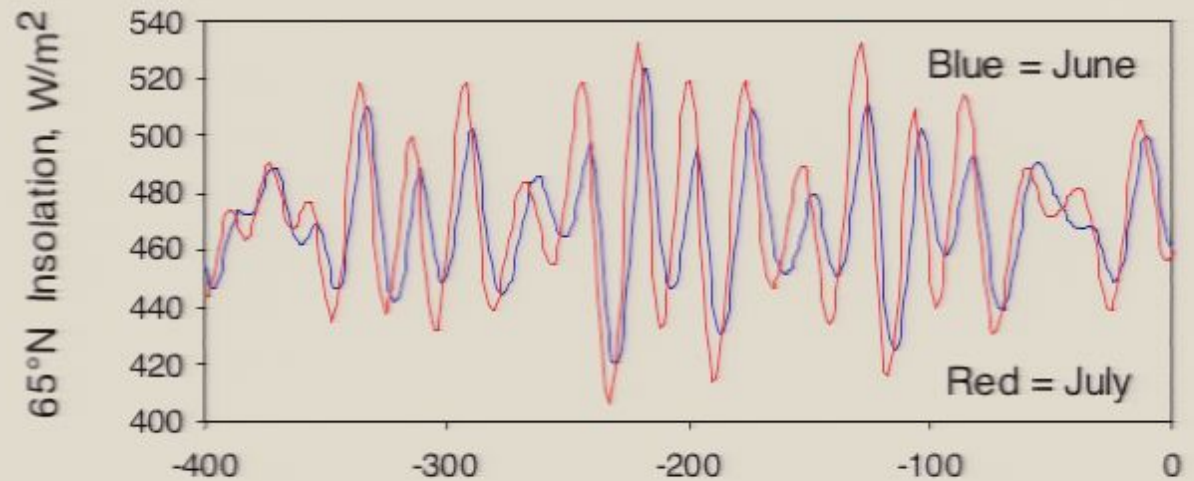


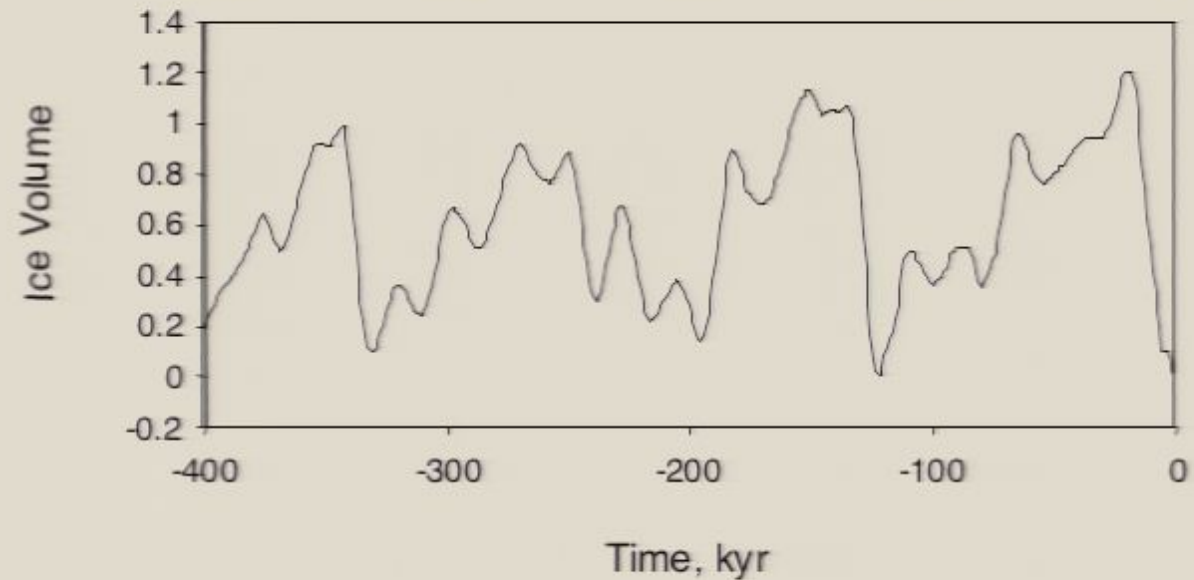
Figure 5. (a) Changes in the precessional parameter $e \sin \bar{\omega}$ for the last million years and the seasonal insolation received at different northern latitudes. (b) Spectral analysis of the precessional parameter changes, revealing two groups of periodicities, around 23 and 19 kyr (arbitrary vertical linear scale).

Orbital forcing and ice volume

Mostly precession.



Ice volume is chunkier than insolation



The spectrum of $\delta^{18}\text{O}$

$\delta^{18}\text{O}$ records a combination of ice volume and deep ocean T

Paillard, 2001

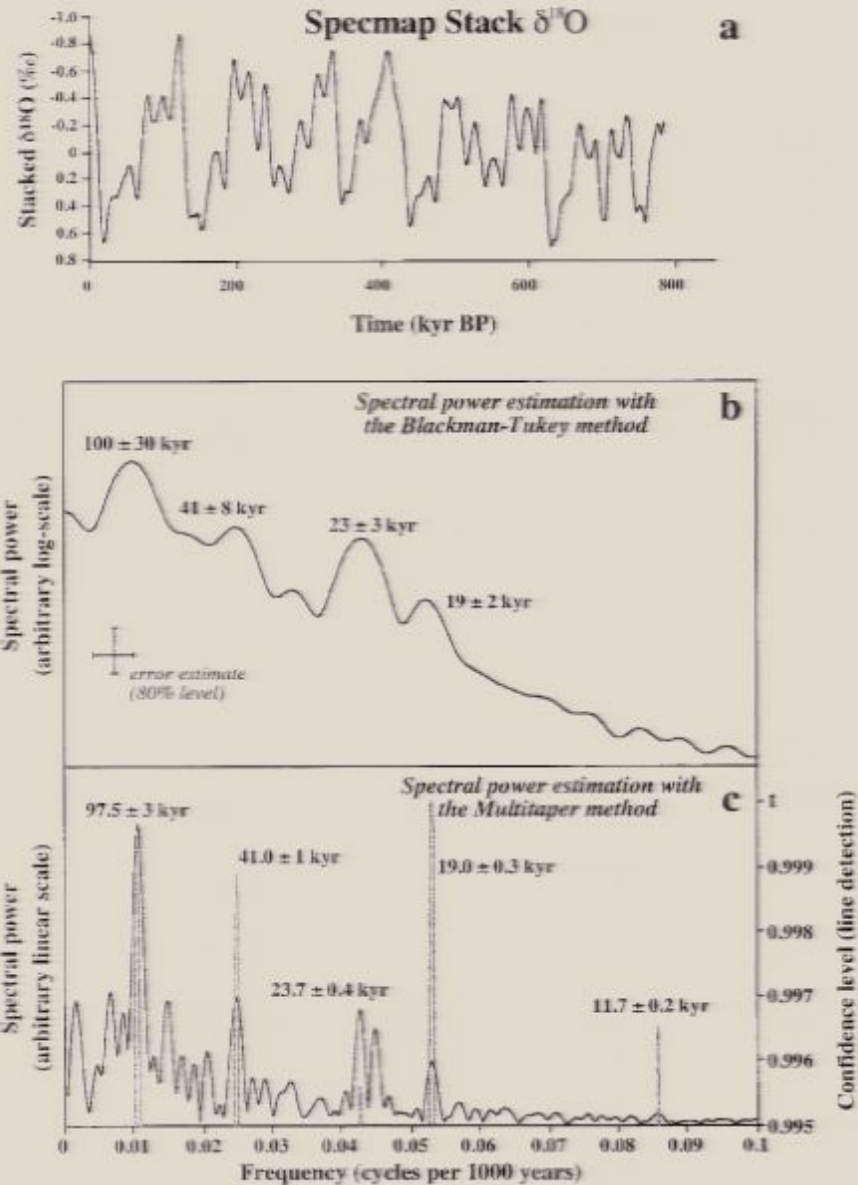


Figure 6. (a) The spectral mapping and prediction (SPECMAP) record [Imbrie et al., 1984]. (b) Spectral analysis of SPECMAP using the standard Blackman-Tukey method. (c) The same analysis with the multitaper method. In Figures 6b and 6c the astronomical frequencies are clearly visible. The first harmonic of the precessional frequency is also detected by the multitaper method.

Precession and obliquity seem to work OK

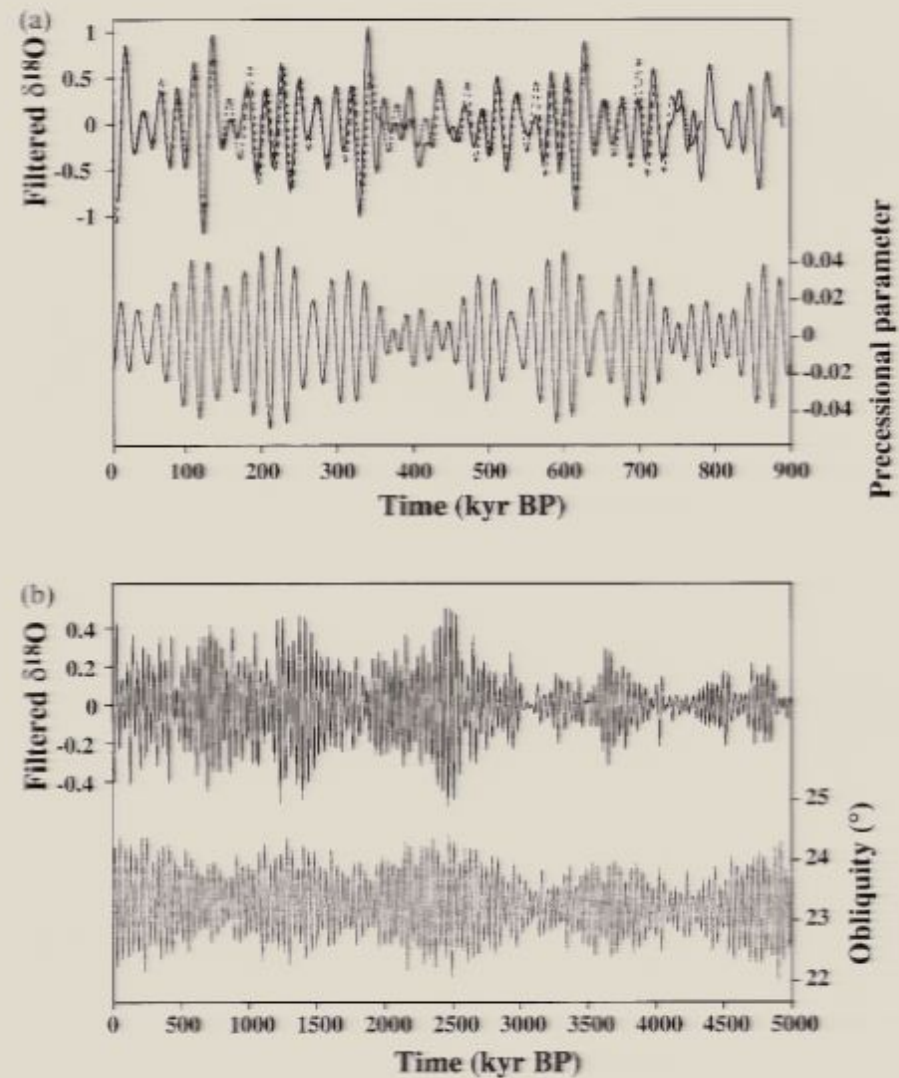
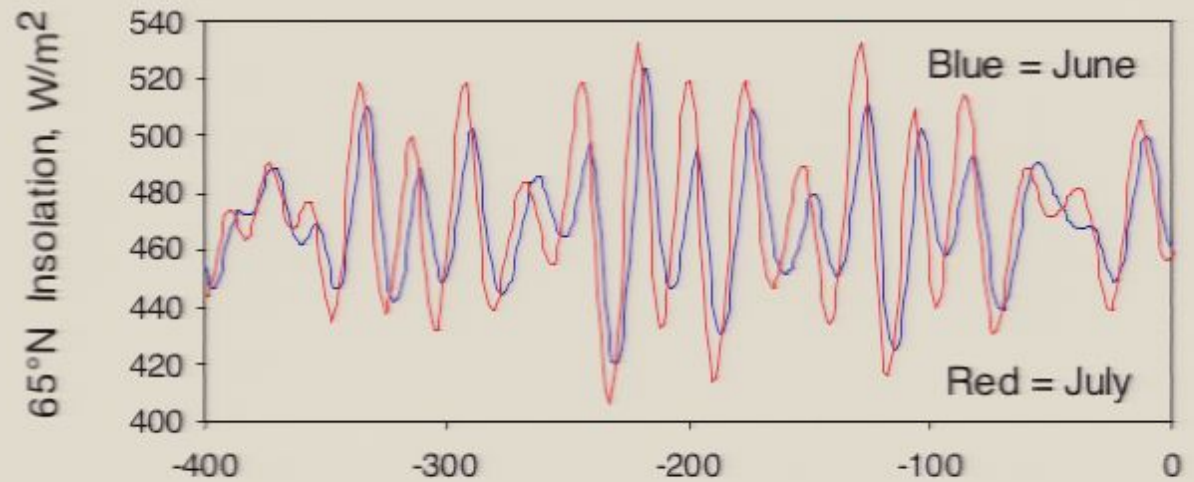


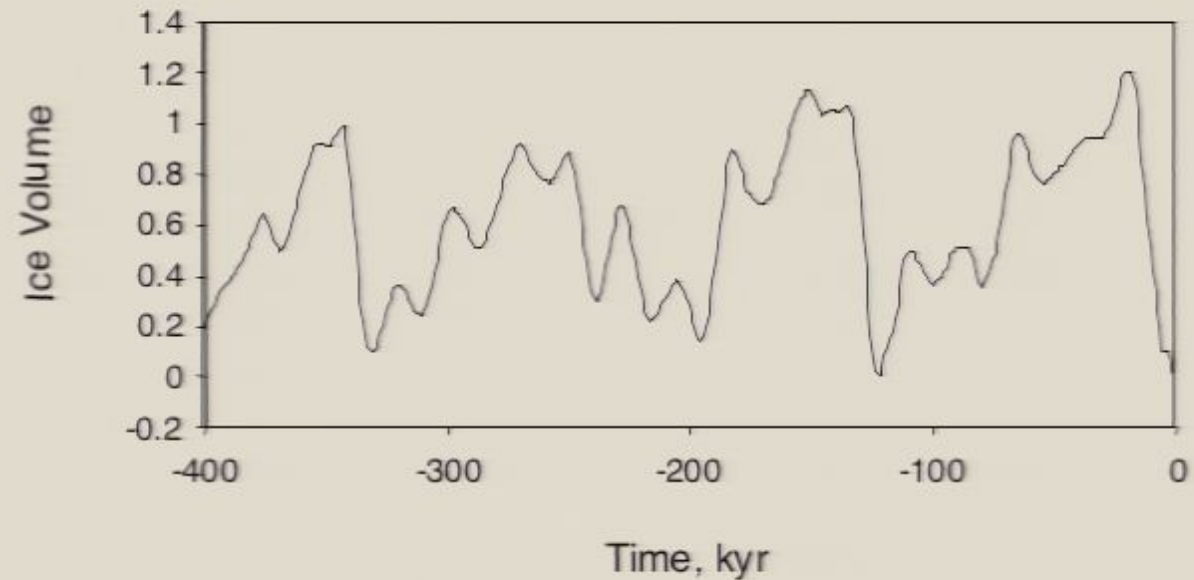
Figure 7. (a) The SPECMAP record [Imbrie *et al.*, 1984] (dashed curve) and the Bassinot *et al.* [1994b] record (bold curve) are filtered in the 23-kyr band and compared with the precessional parameter. (b) The 5-Myr long Ocean Drilling Program (ODP) 659 $\delta^{18}\text{O}$ record [Tiedemann *et al.*, 1994] is filtered in the 41-kyr band and compared with obliquity. The amplitude modulation of both the 23-kyr and the 41-kyr cyclicity appears very similar in the astronomical forcing and in the paleoclimatic record. This is probably the strongest argument in favor of a simple quasi-linear relationship between the climatic system and insolation forcing in these two frequency bands.

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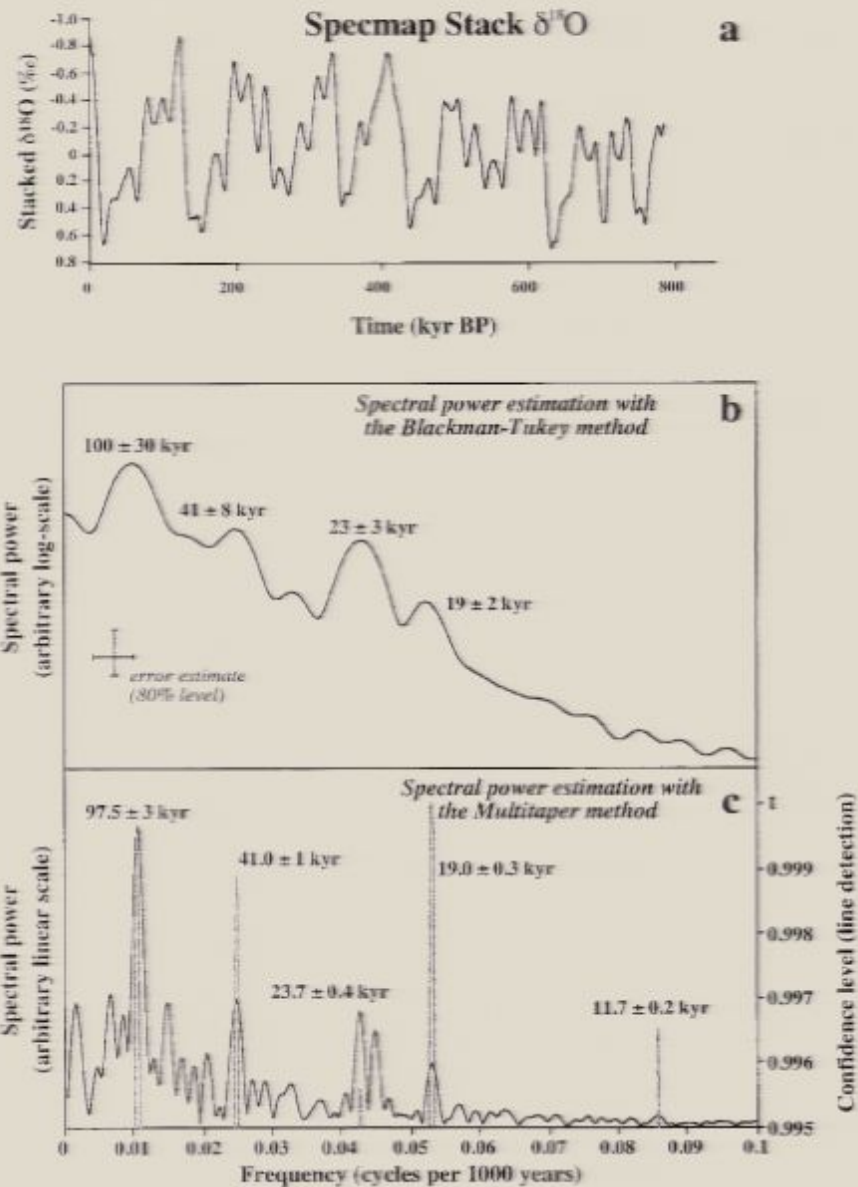


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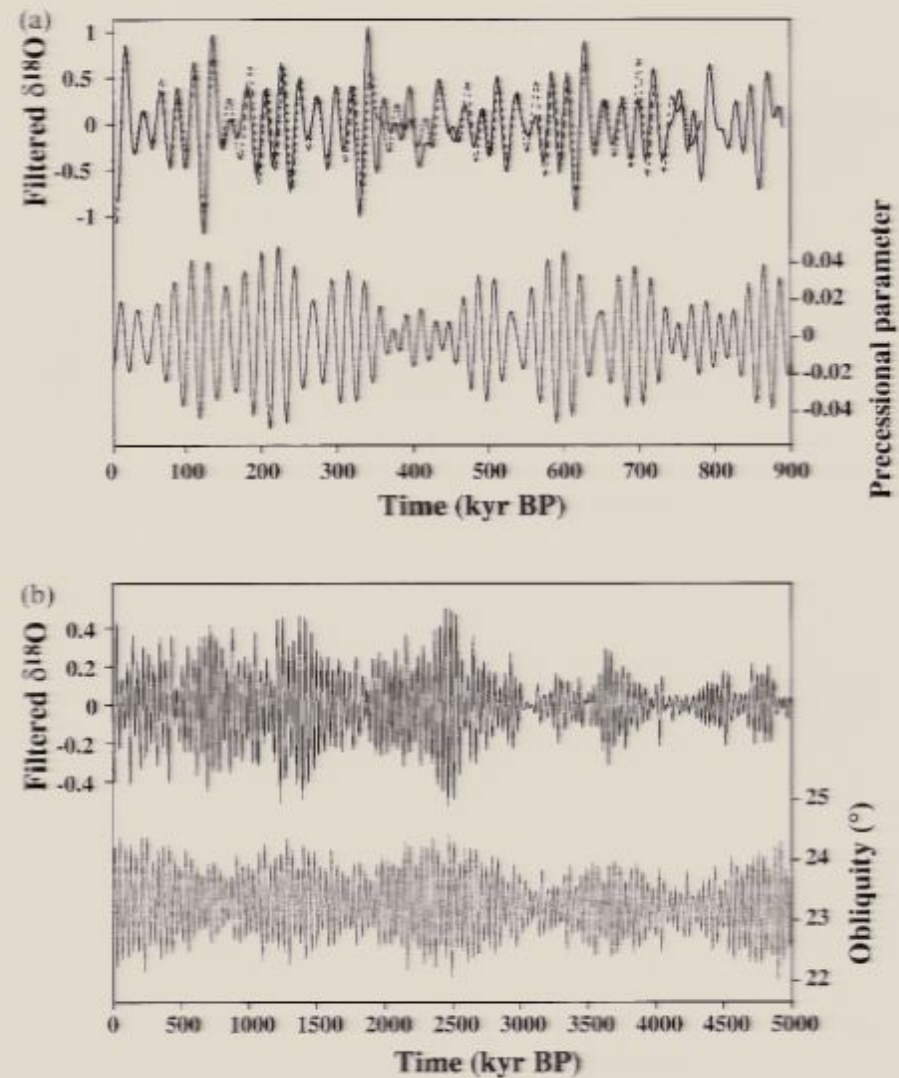


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Paillard, 2001

Eccentricity doesn't look so terrible, either. (On the face of it).

Needs suppression of the 400 kyr part

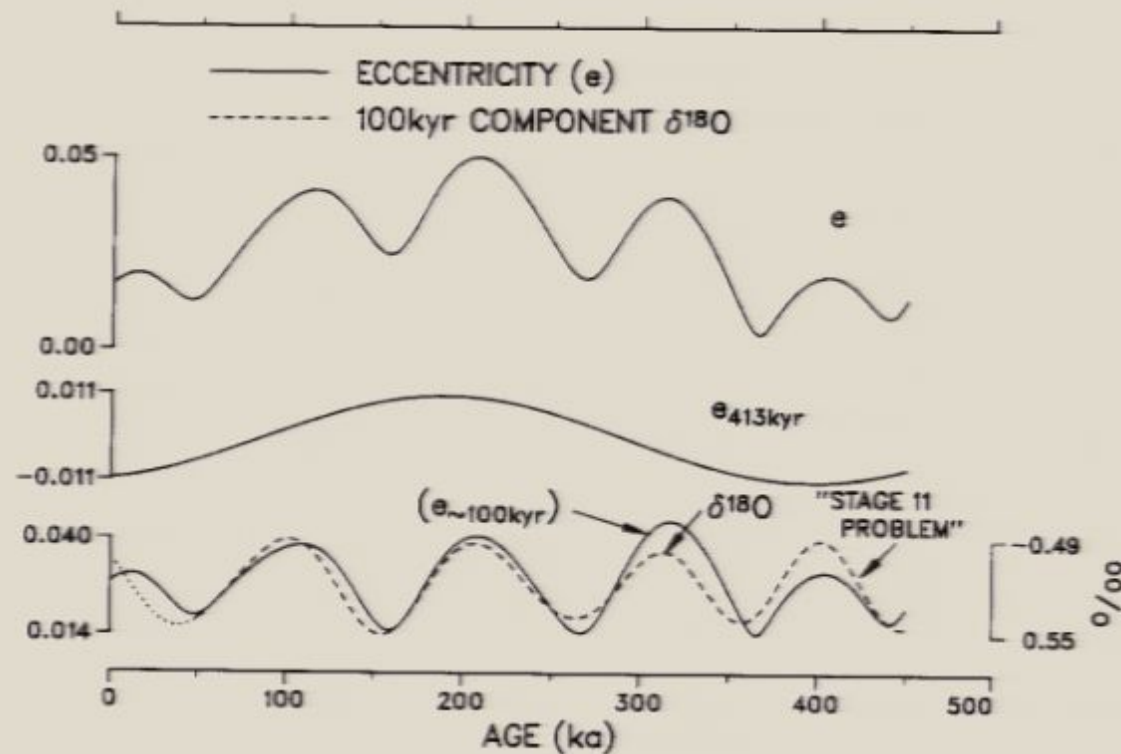


Fig. 2. Eccentricity and the 100-kyr $\delta^{18}\text{O}$ cycle. With its dominant 413-kyr component ($e_{413\text{ kyr}}$) removed from eccentricity e , the residual signal $e_{\sim 100\text{ kyr}}$, calculated as $e - e_{413\text{ kyr}}$, is dominated by variance over a moderately broad band of periods near 100 kyr [Berger, 1978a, b]. Averaged over the entire interval, this part of the eccentricity signal is coherent with the 100-kyr $\delta^{18}\text{O}$ cycle and leads it systematically by $\sim 13^\circ$. But the $\delta^{18}\text{O}$ response is not proportional to eccentricity in Stage 11.

Imbrie, 1993

Eccentricity

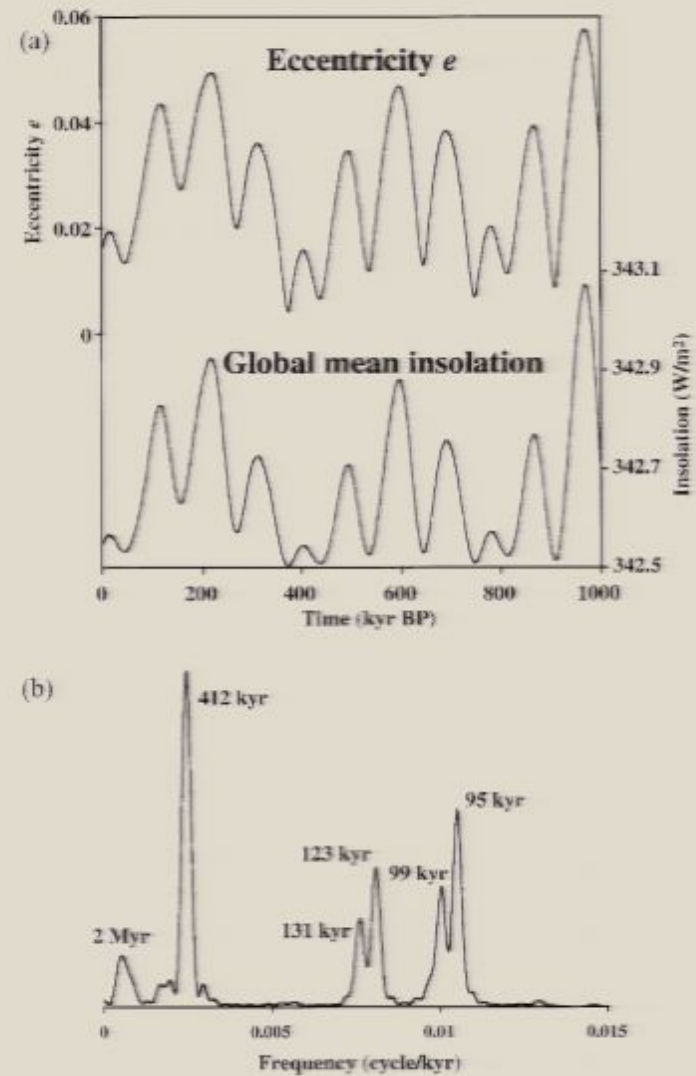


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Eccentricity needs a 20x amplifier

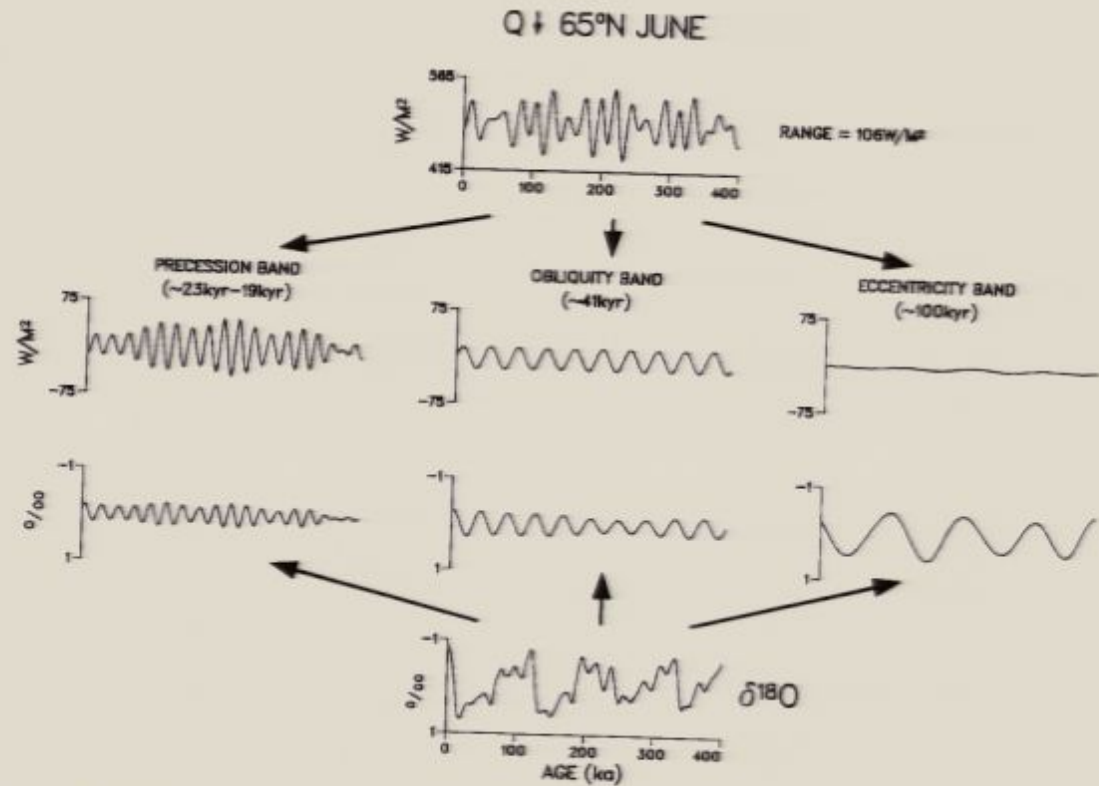


Fig. 1. The 100-kyr cycle problem as seen by partitioning radiation and climate time series into their dominant periodic components (in the precession, obliquity, and 100-kyr eccentricity bands). Radiation time series are from Berger [1978a]; $\delta^{18}\text{O}$ data are from Imbrie et al. [1984]. Partitioning is done using Hamming band-pass filters with a bandwidth of 0.019 kyr^{-1} for the 41- and 100-kyr bands and 0.036 kyr^{-1} for the 23-kyr band [Jenkins and Watts, 1968]. The $\delta^{18}\text{O}$ cycles at periods near 23, 41, and 100 kyr are so strongly correlated with astronomically driven radiation cycles as to suggest a causal linkage in all three bands. But these correlations for the 23-, 41-, and 100-kyr bands (coherencies of 0.95, 0.90, and 0.91, respectively, in Table 2) hide an intriguing physical problem. Why is the system's response so strong in the 100-kyr band? There the amplitude of the radiation signal (2 W m^{-2}) is 1 order of magnitude smaller than in the other two bands.

The three cycles evolved independently

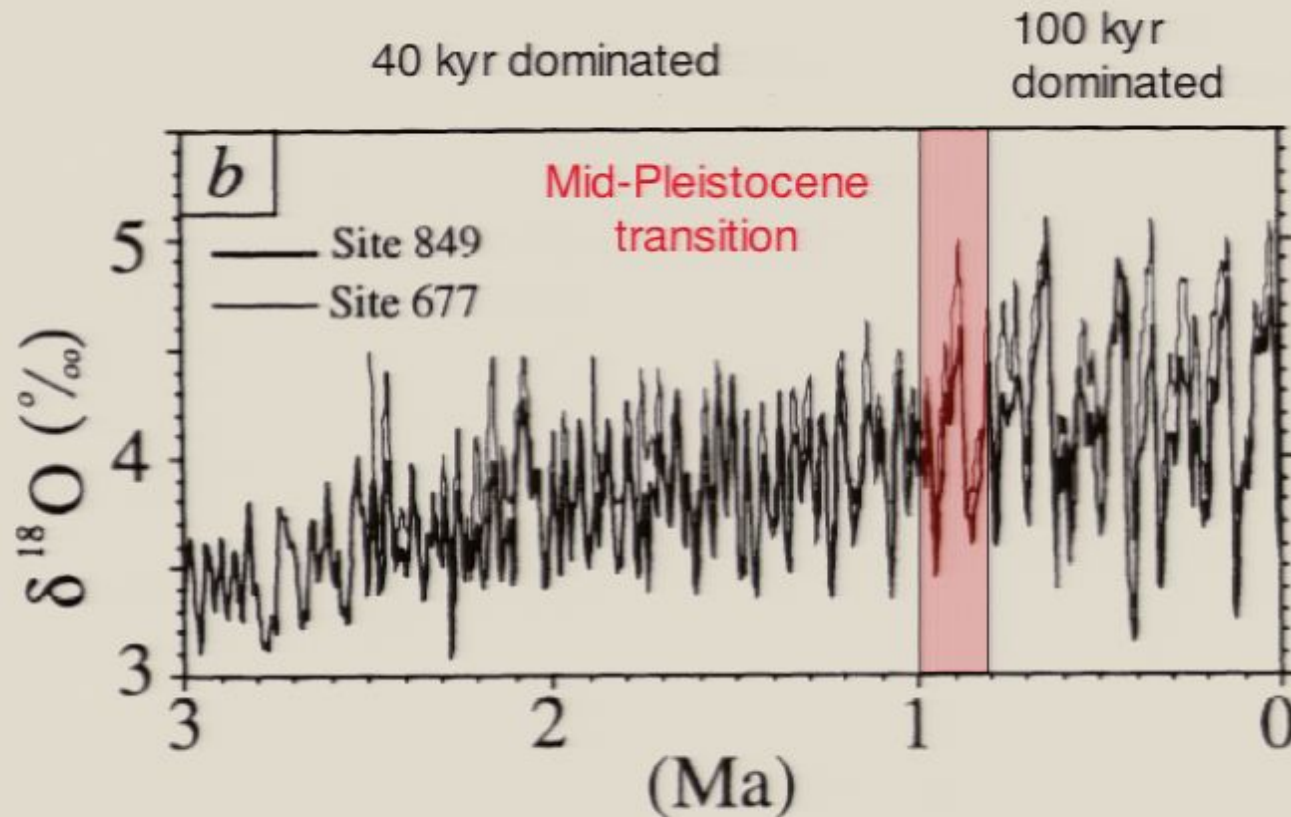


Figure 1. (a) Summer half-year insolation (in W m^{-2}) at 55°N for the last 3 Myr computed using orbital elements from *Berger and Loutre [1991]*. (b) Oxygen isotope records for Ocean Drilling Program Sites 677 and 849 over the last 3 Myr (data and age models are from *Mix et al. [1995]*).

Calder's model

$$\frac{dV}{dt} = -k (i - i_0)$$

insol



ice



Grows if $i < i_0$, melts otherwise

k is different for growing or melting. This nonlinearity generates some 100kyr power in the output.

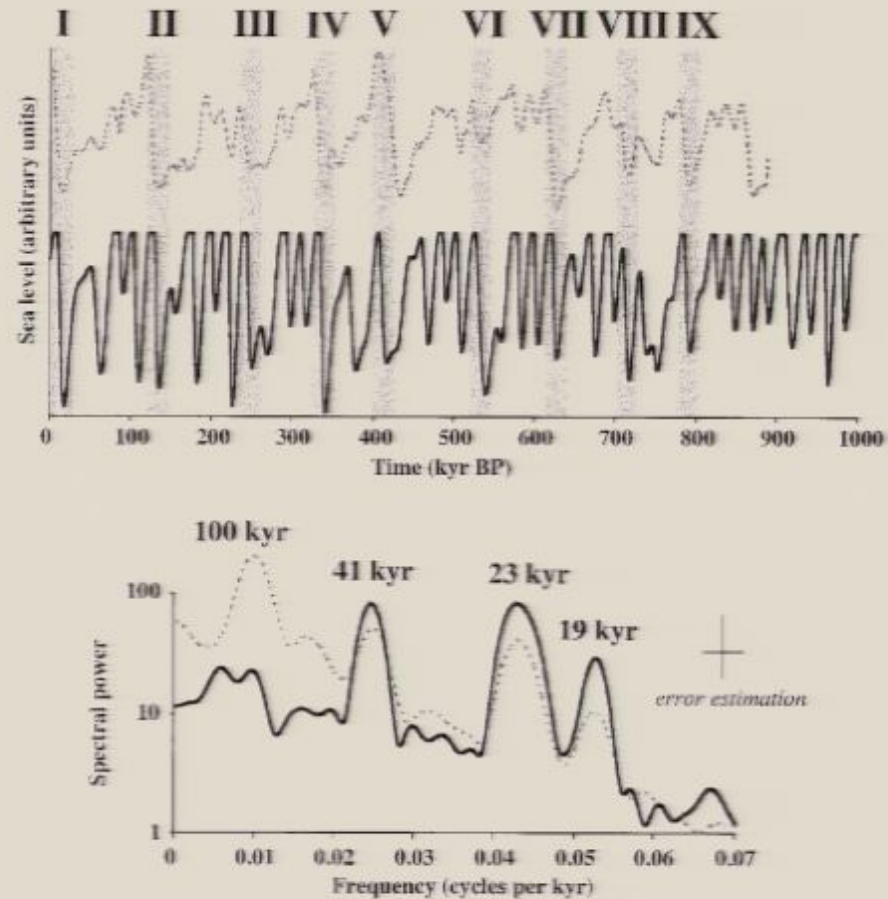
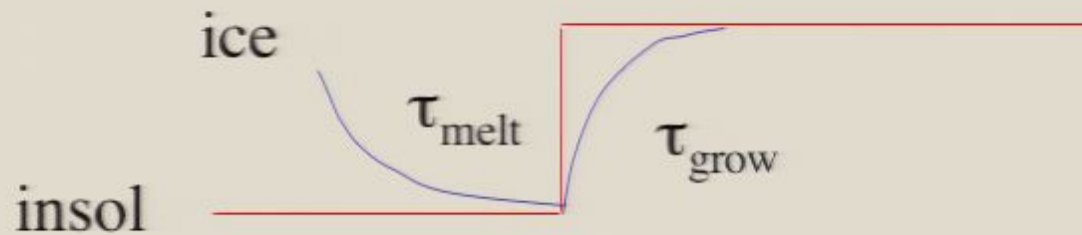


Figure 9. Results from the *Calder* [1974] model. The threshold i_0 is equal to 502 W m^{-2} , and the ratio k_A/k_M is chosen equal to 0.22. The forcing i is the summer solstice insolation at 65°N [Laskar, 1990]. The result is very sensitive to these choices. The agreement with the record is quite poor, but this crude model still predicts the major transitions at the right time, a feature that many, more sophisticated models do not reproduce well. An isotopic record is given here for comparison [Bassinot *et al.*, 1994b].

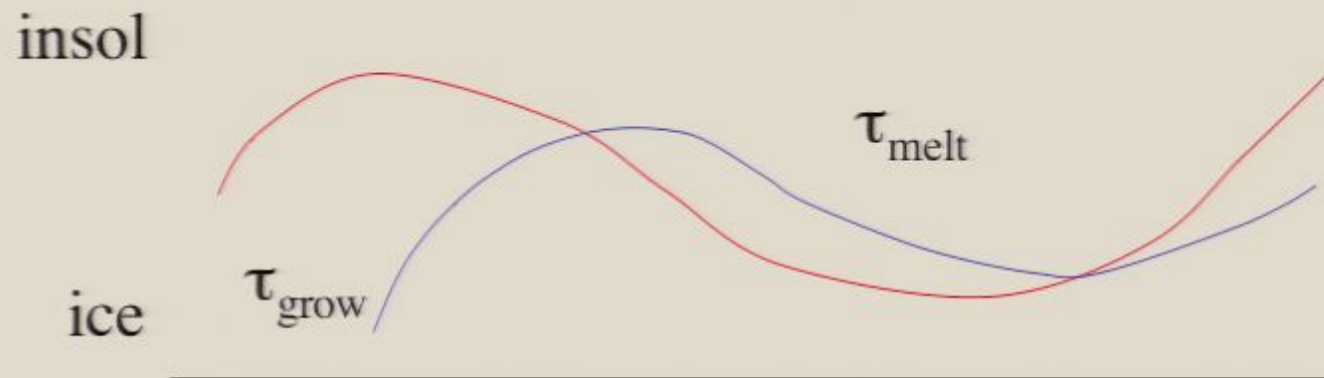
Imbrie and Imbrie model

$$\frac{dV}{dt} = \frac{(i - V)}{\tau}$$



Imbrie and Imbrie model

$$\frac{dV}{dt} = \frac{(i - V)}{\tau}$$



Imbrie and Imbrie model

Gets 100k,
but also 400k

Not there
in the data.

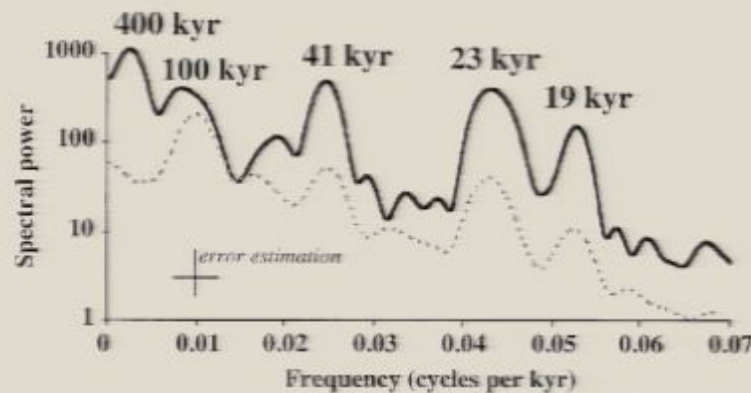
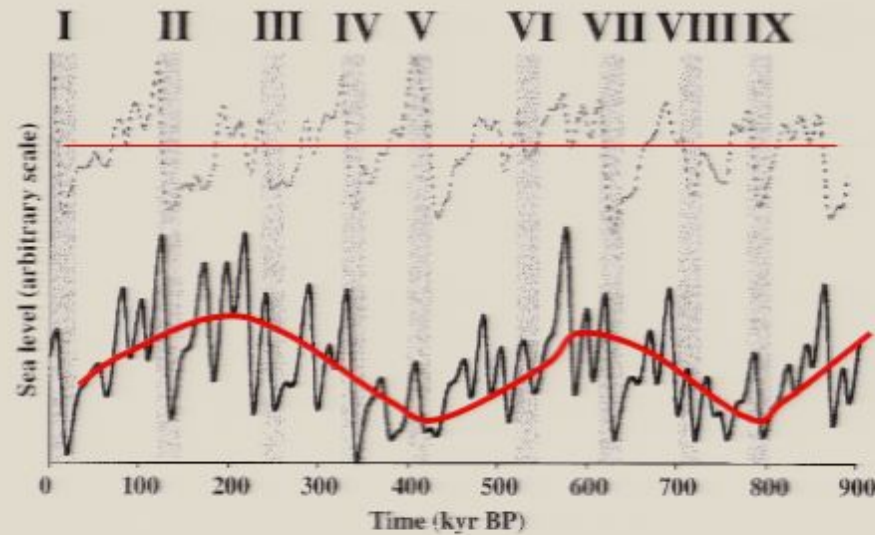


Figure 10. Same as Figure 9, but for the Imbrie model [Imbrie and Imbrie, 1980]. The forcing i is the summer solstice insolation at 65°N. The time constants are $\tau_M = 42$ kyr and $\tau_A = 10$ kyr.

Paillard threshold model

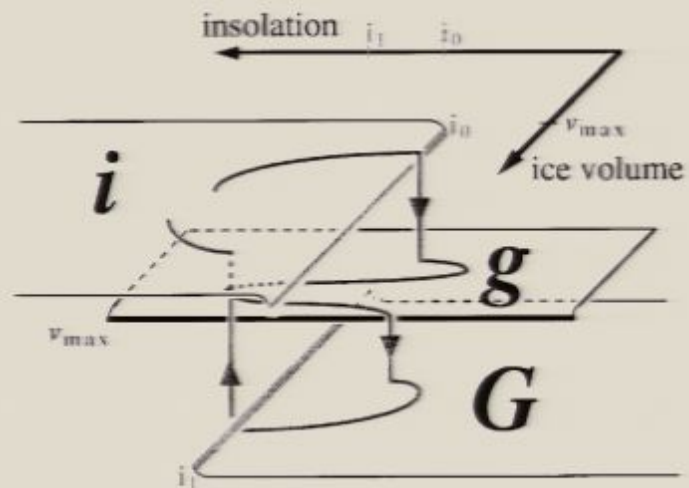
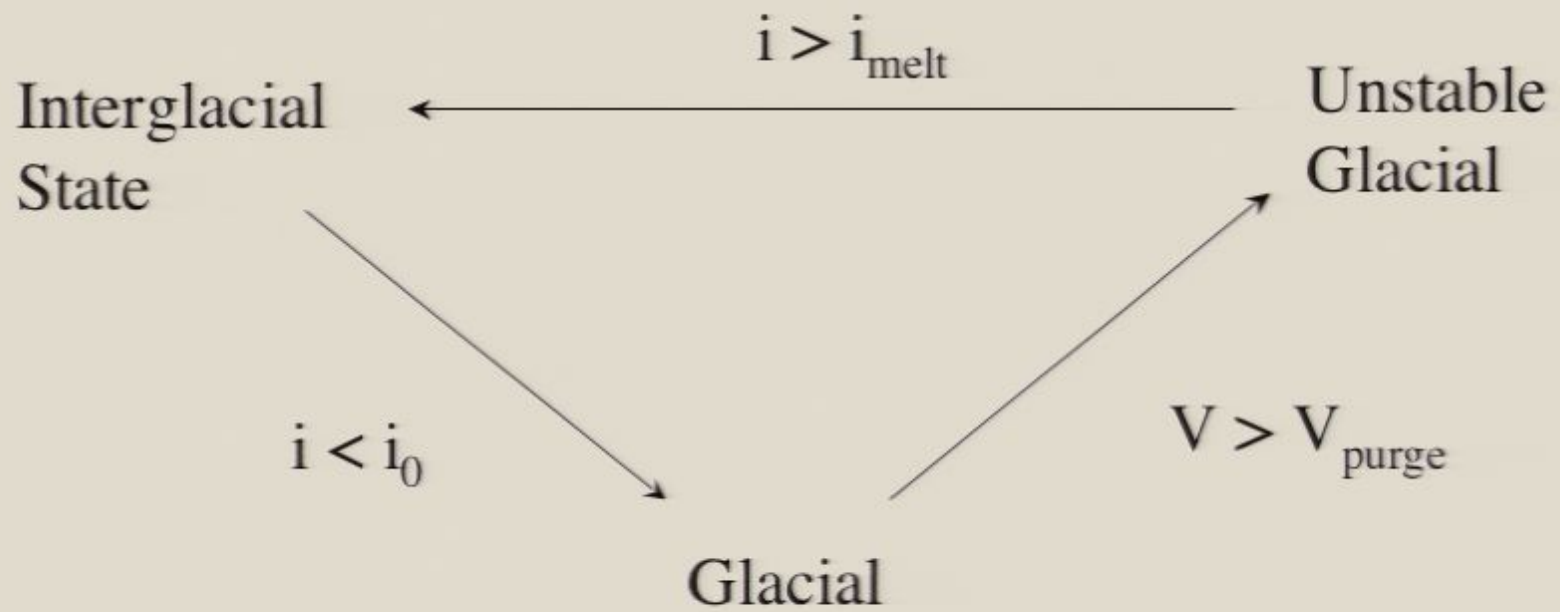
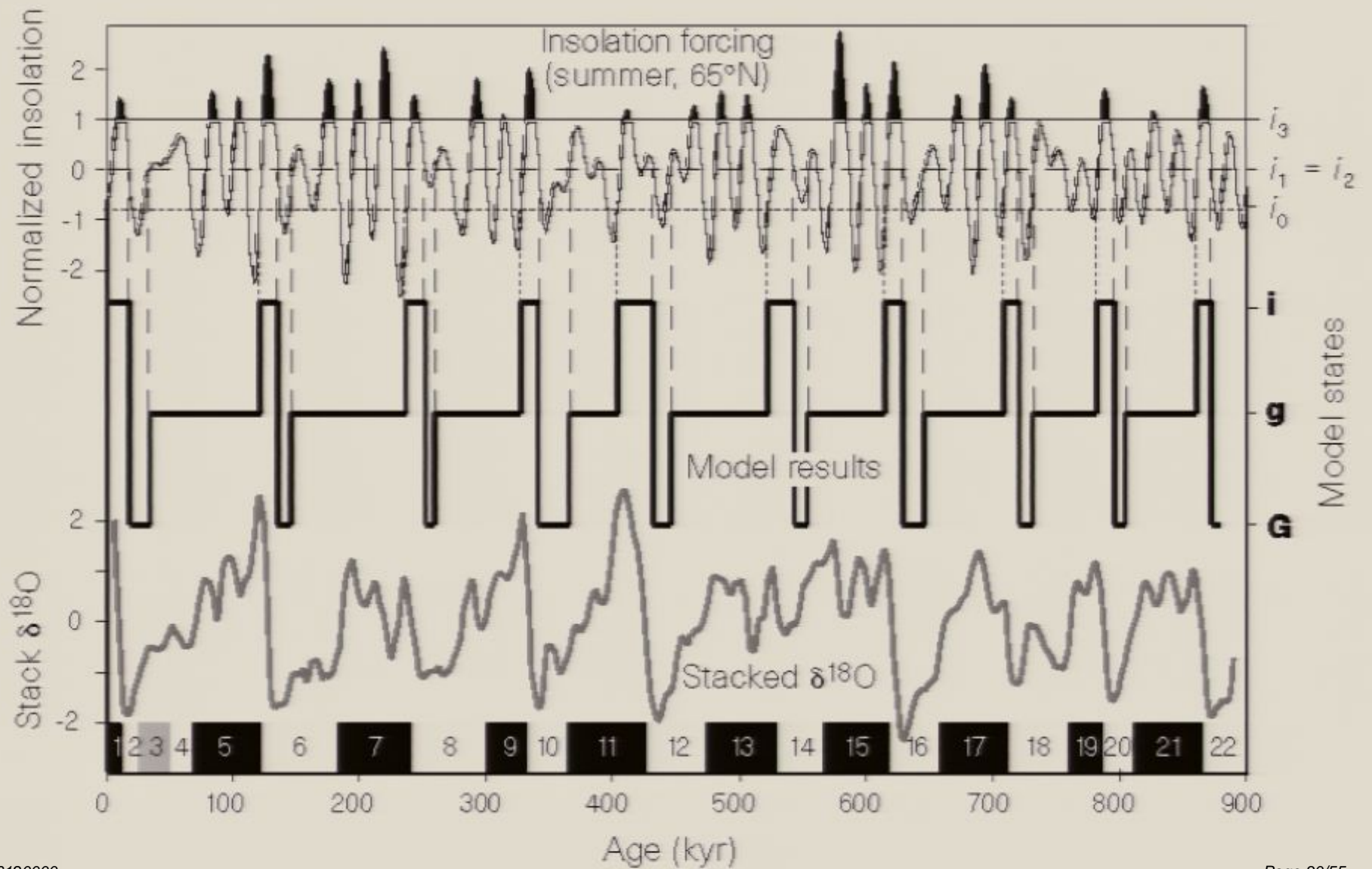


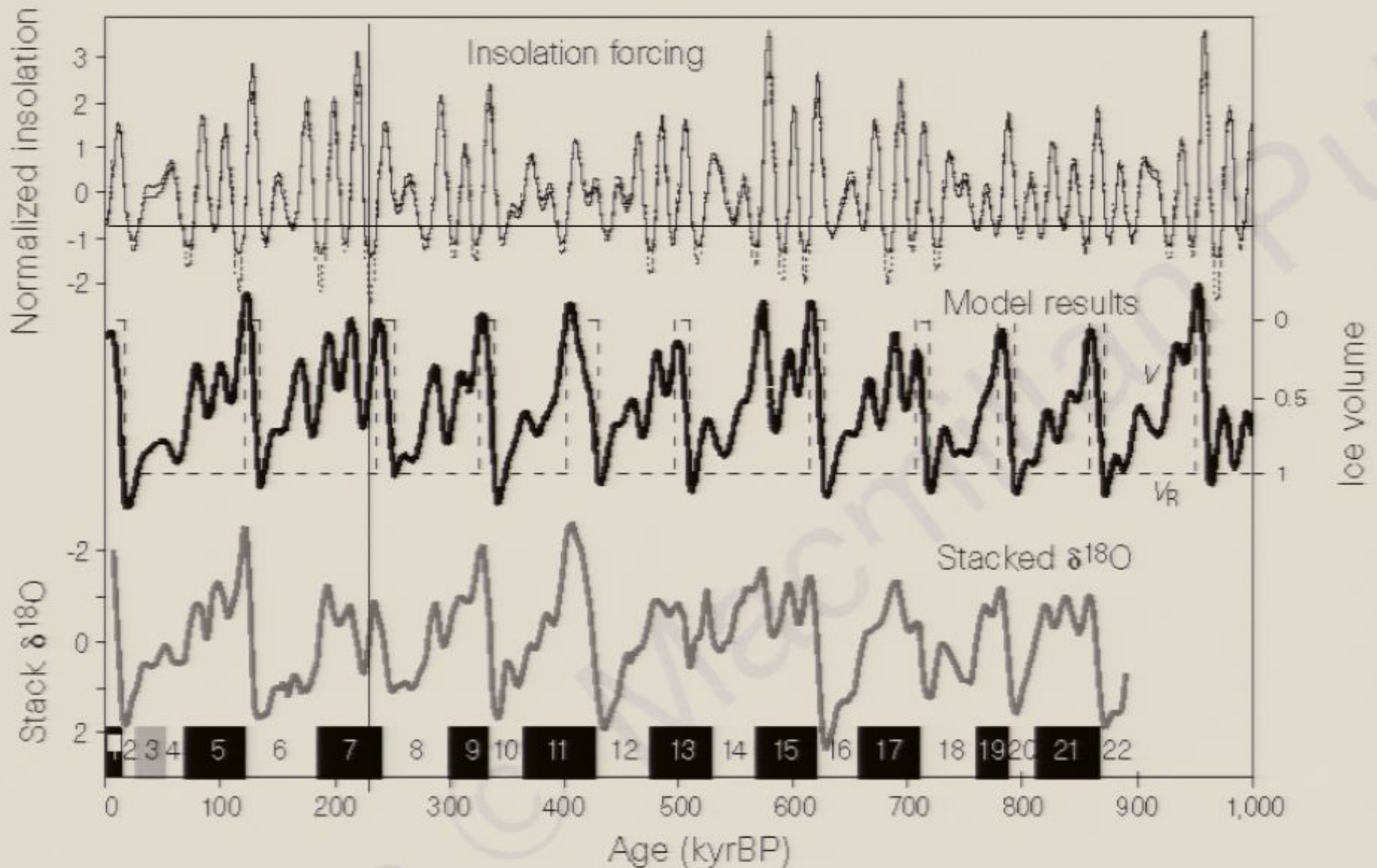
Figure 12. The threshold model. Climate is assumed to have three different regimes: i (interglacial), g (mild glacial), and G (full glacial). Transition between the regimes occurs when the insolation forcing crosses a given threshold i_0 or i_1 , or when the ice volume exceeds the value v_{MAX} .

Paillard model





Gussy it up with ice growth/melting kinetics



Paillard model
gets the spectrum
better

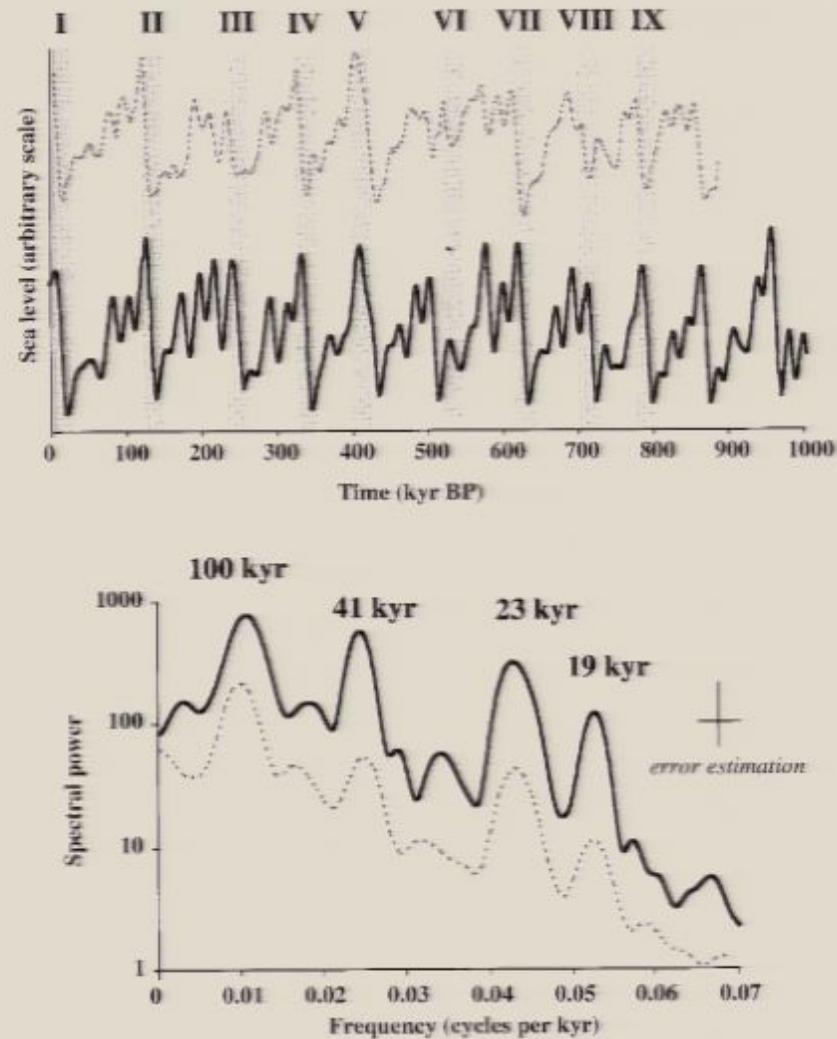
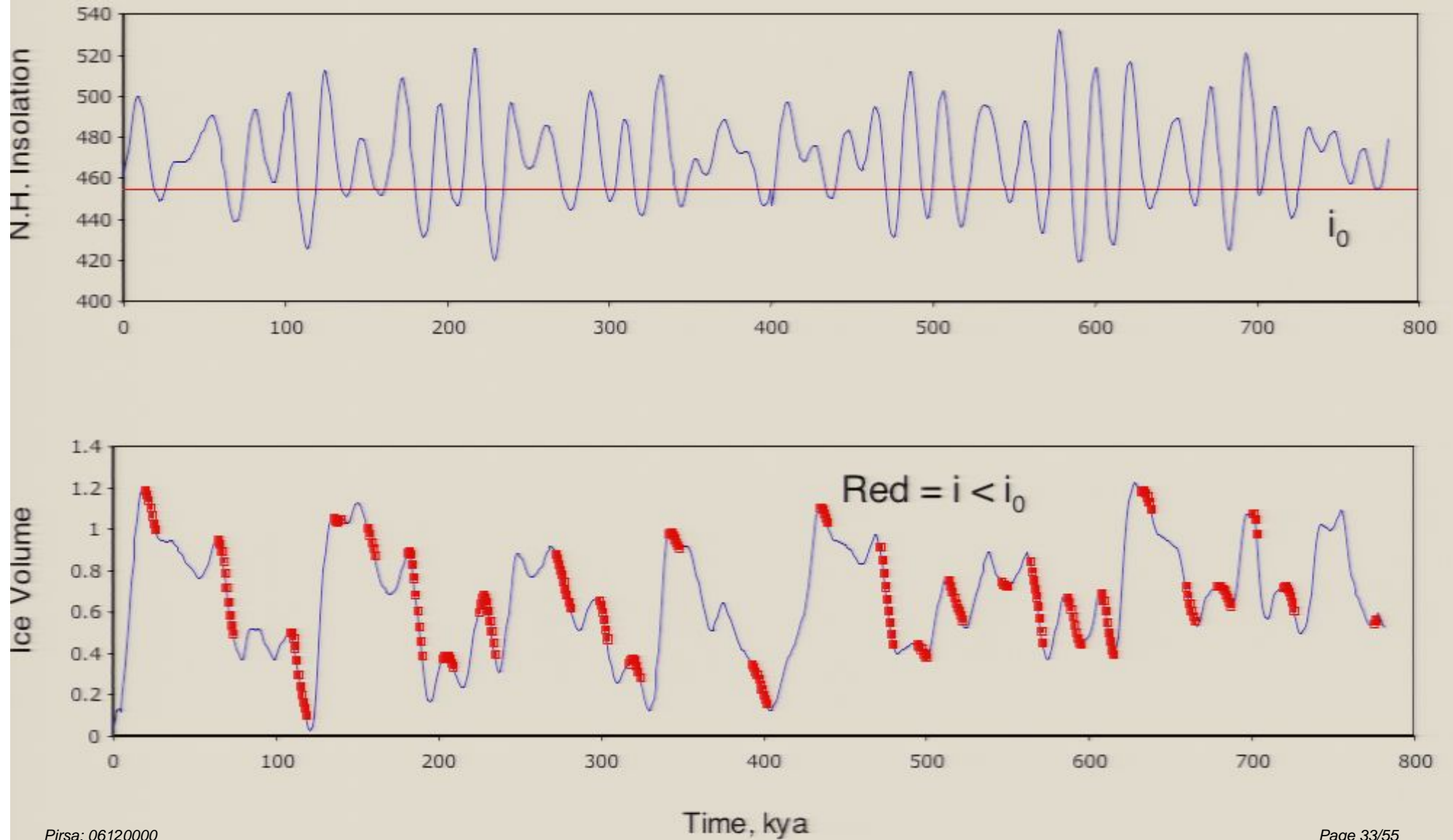


Figure 13. Same as Figure 9, but for Paillard's [1998] model. Threshold values are $i_0 = -0.75$ and $i_1 = 0$. Time constants are $\tau_i = 10$ kyr, $\tau_G = \tau_F = 50$ kyr, and $\tau_F = 25$ kyr.

According to Paillard, 2001

Ice volume responds to N.H. insolation



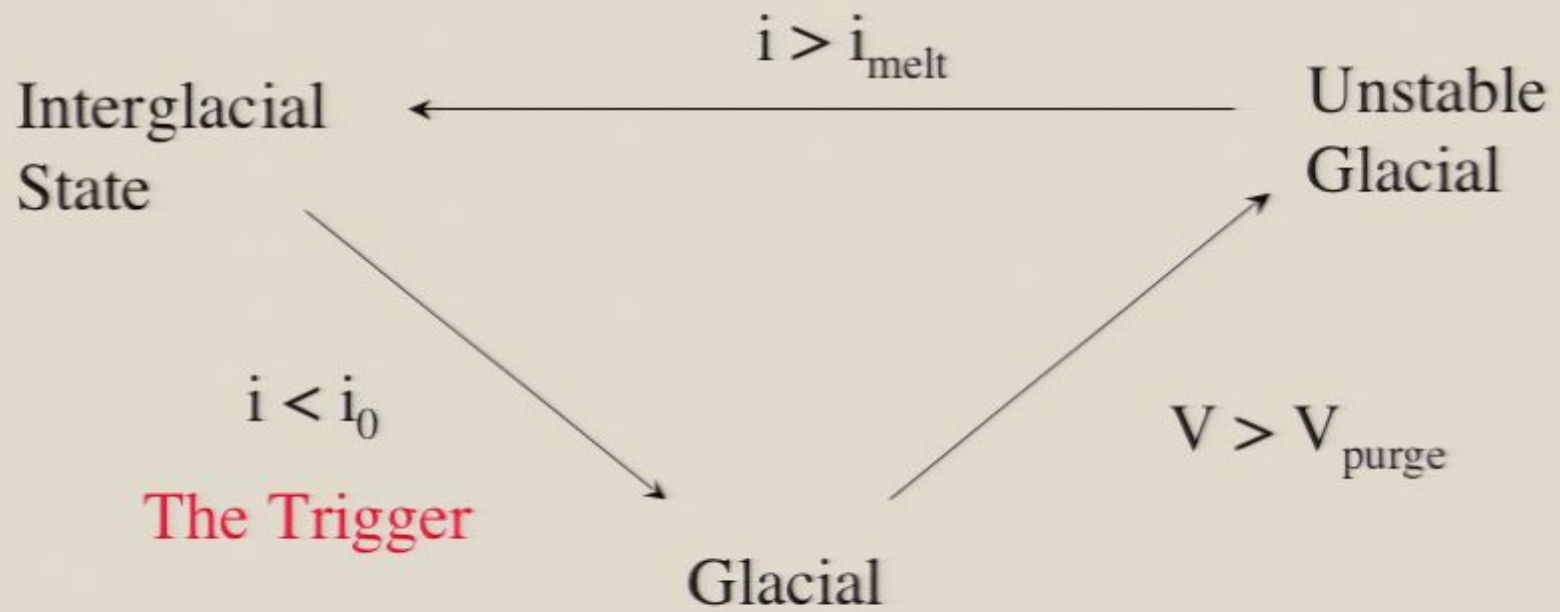
The World According to Paillard

100 kyr cycle arises from the one-way street of the ice sheet life cycle.

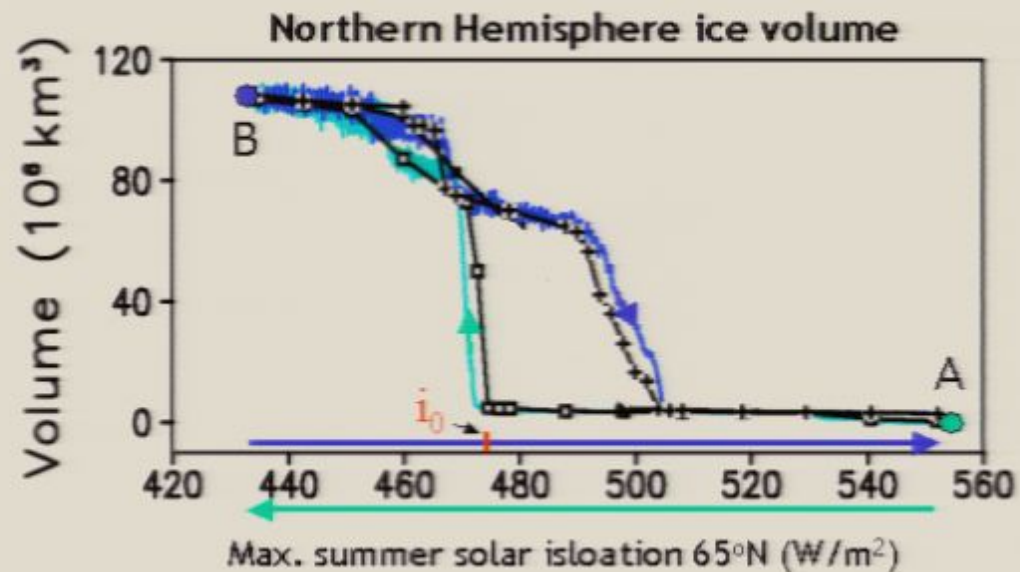
Ice definitely responds to $i < i_0$ cold bursts.

The definite link with nucleation is not as clear.
The $\delta^{18}\text{O}$ impact is similar whether in warm climate or cold.

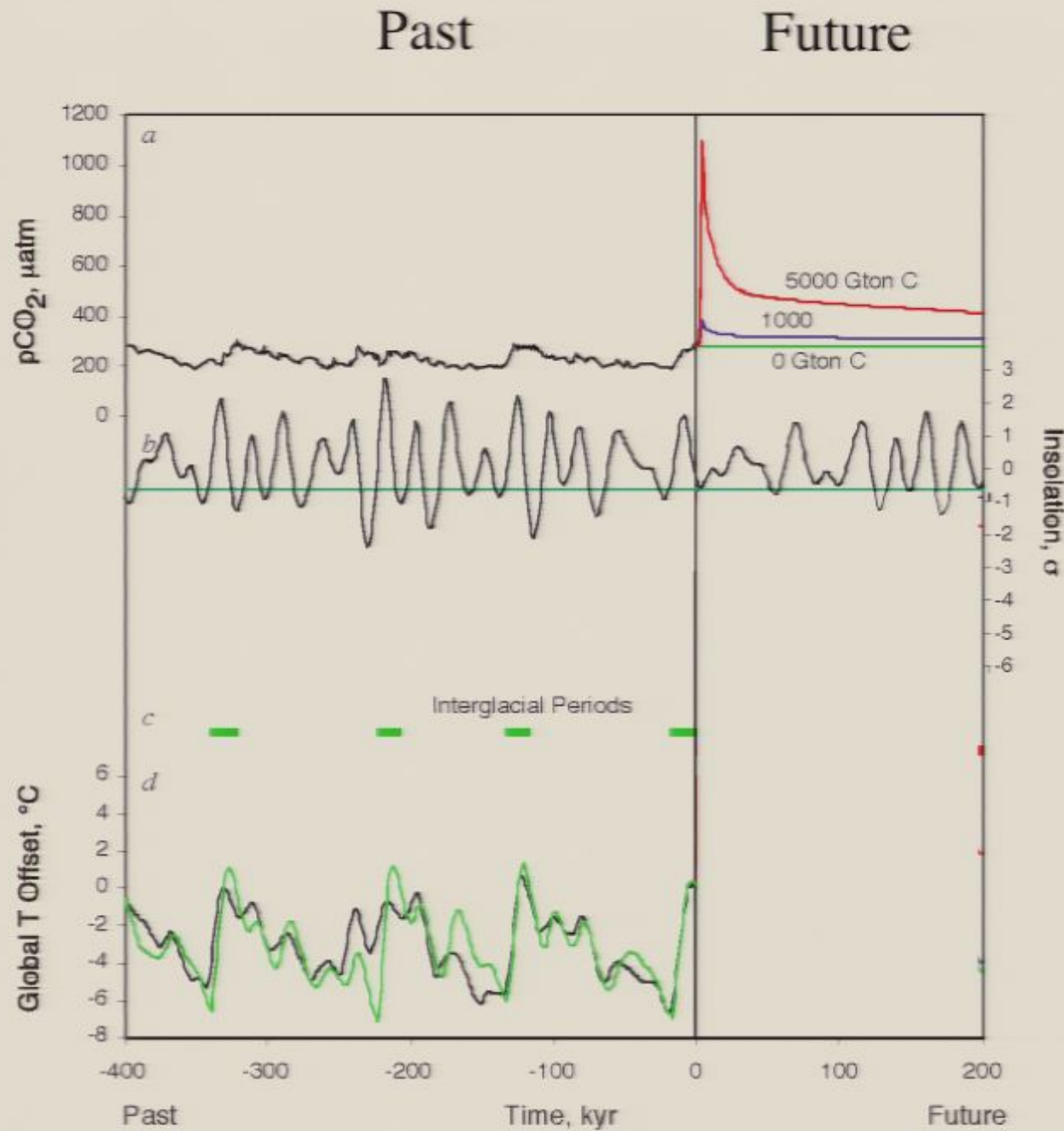
Paillard model



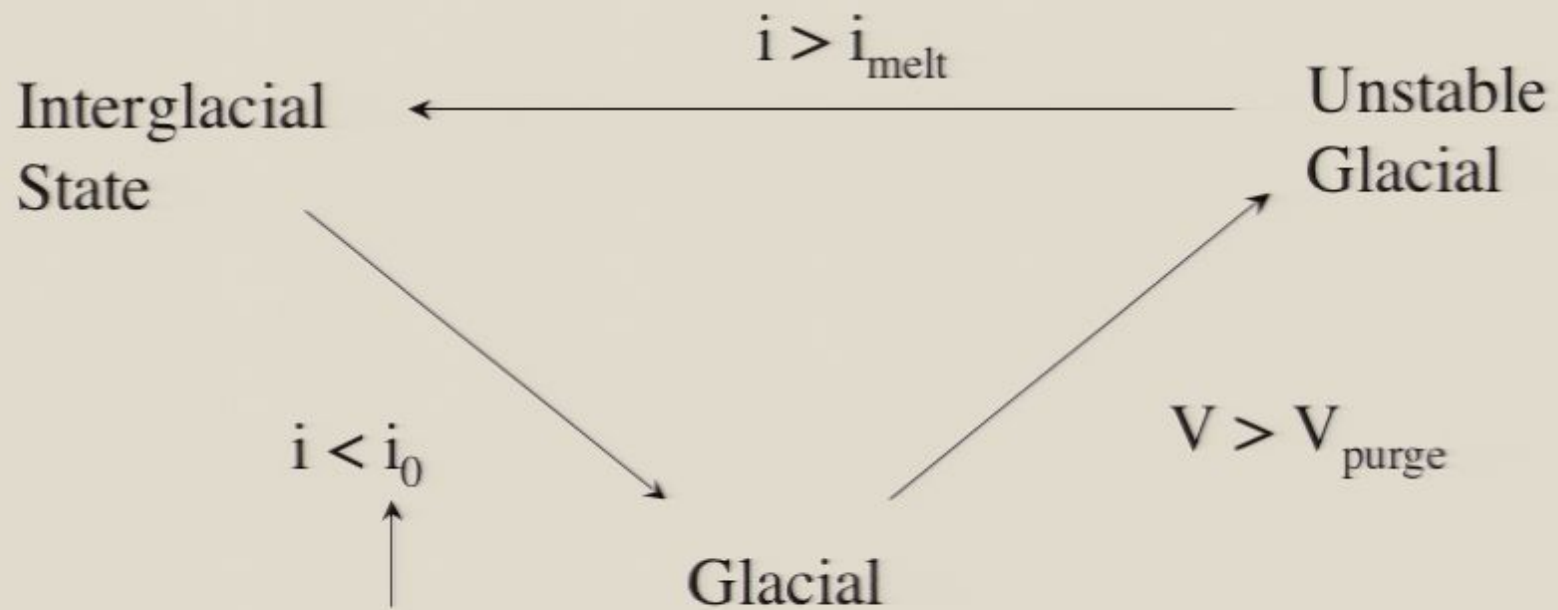
CLIMBER Model Hysteresis



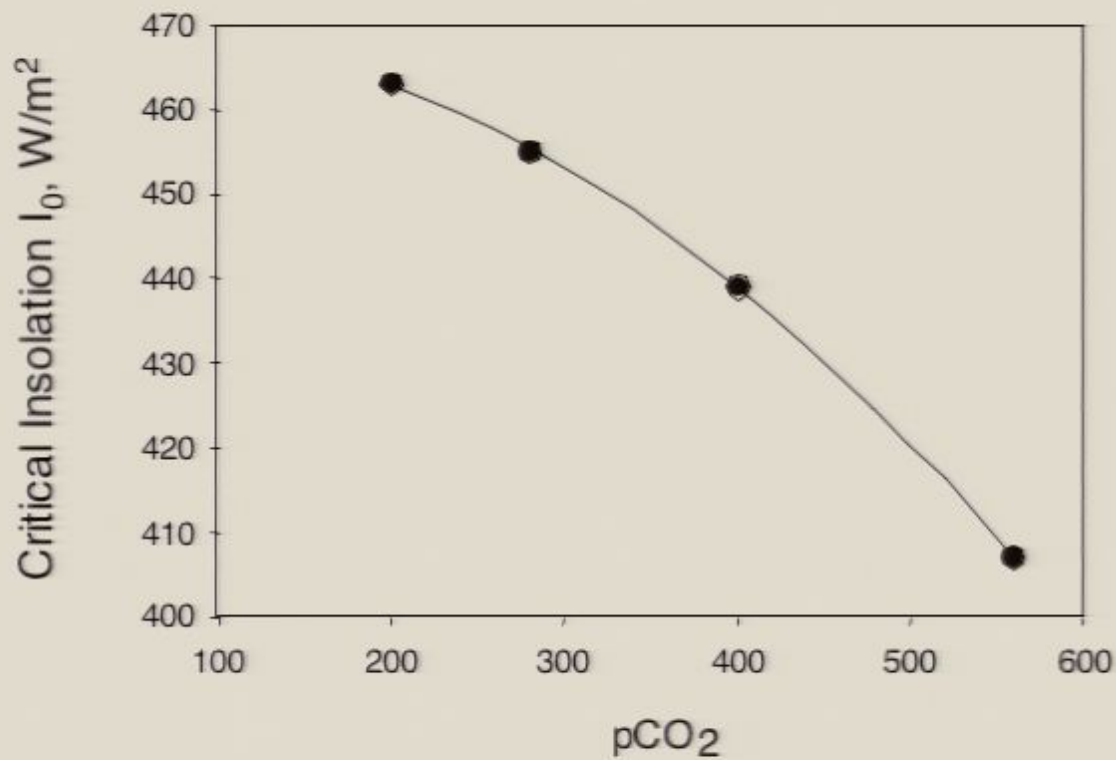
Interglacial mode until sunlight gets too dim.
Then suddenly an ice sheet forms, and it stays there
until sunlight gets even brighter than the original trigger.



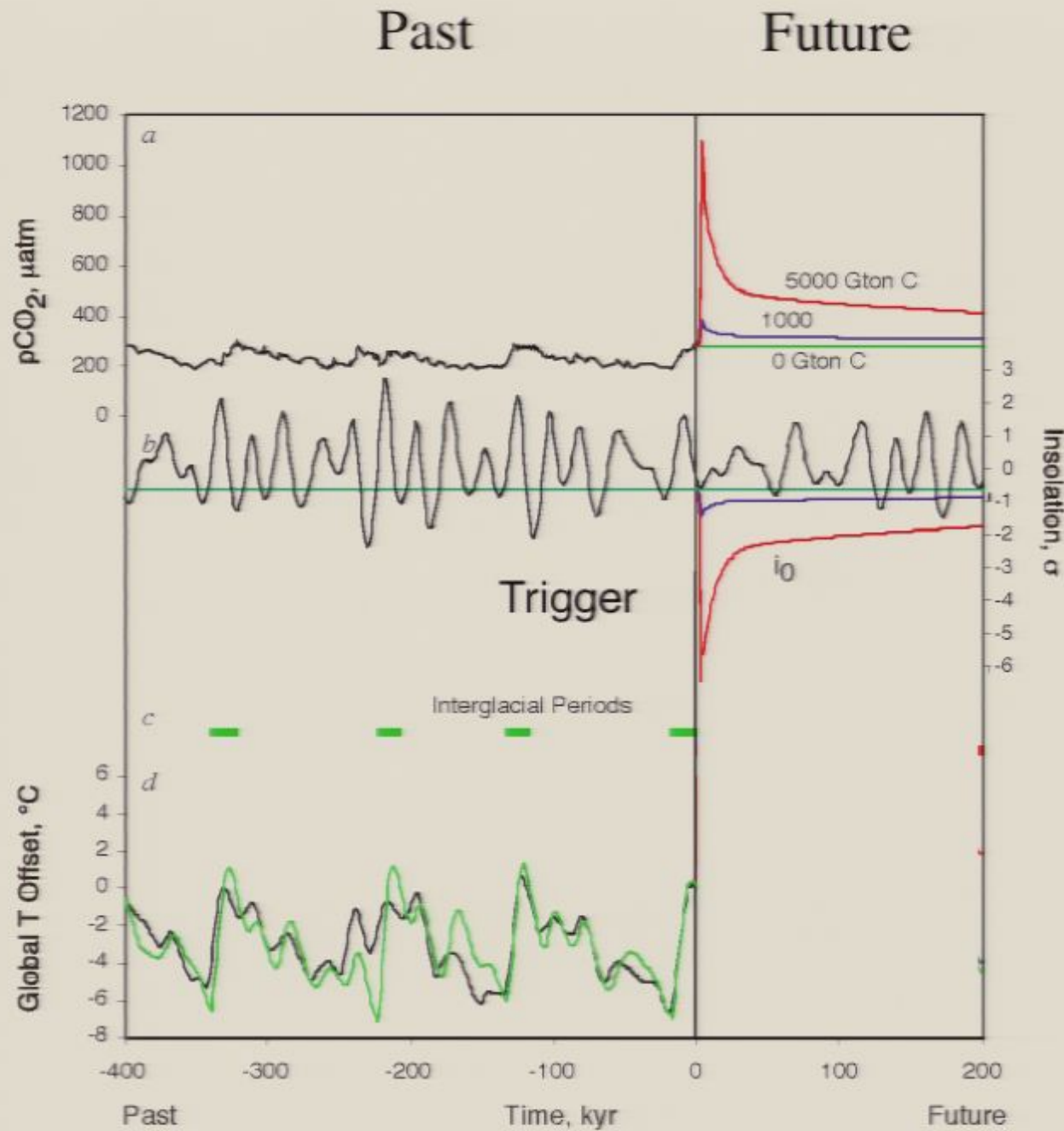
Paillard model



Baseline $p\text{CO}_2$
changes i_0 ,
moving the trigger



If CO₂ is higher, it takes a stronger sunlight change to trigger formation of an ice sheet.

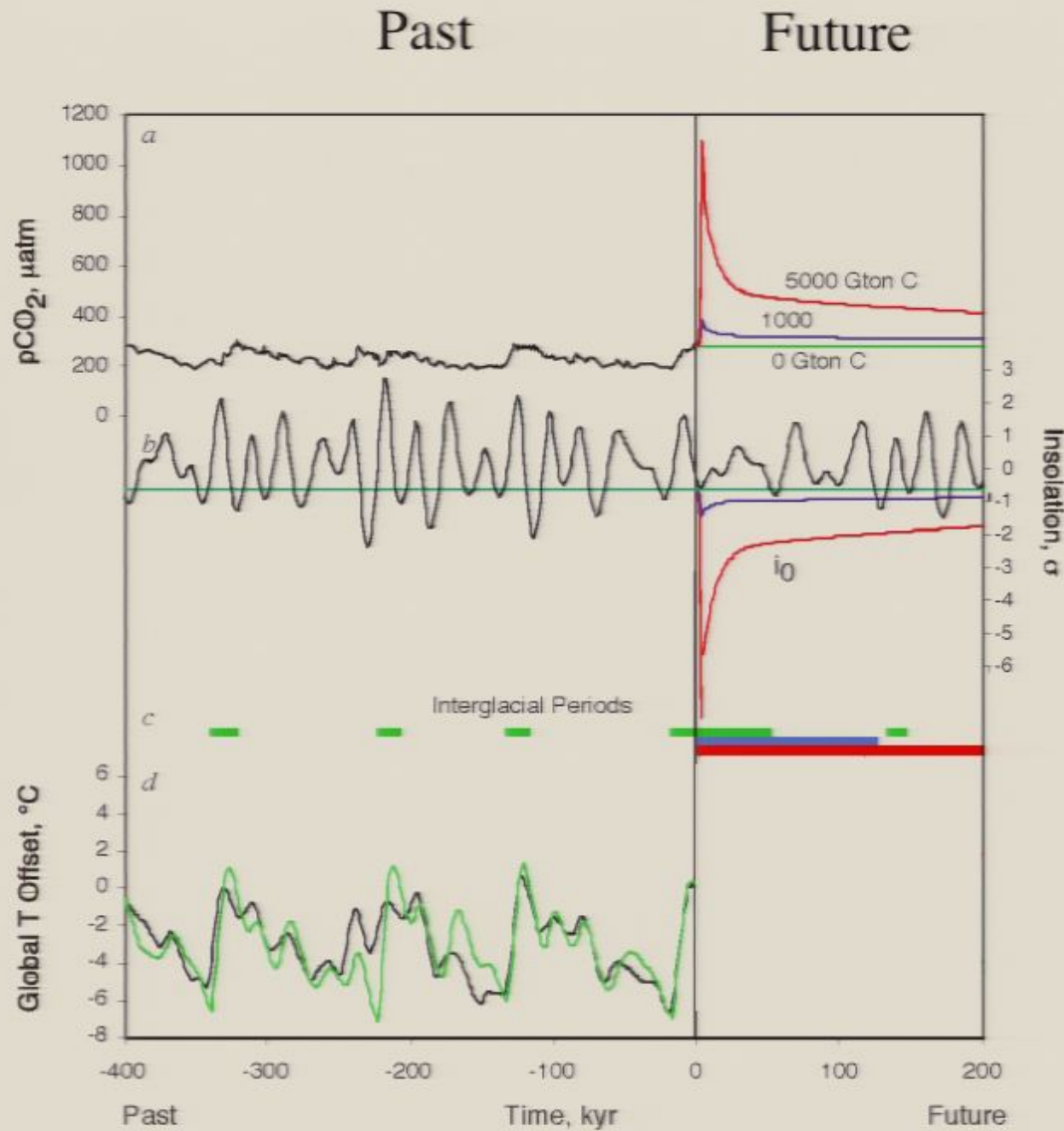


A Crude Temperature Estimate

$$\Delta T_{\text{total}} = 6^{\circ}\text{C} \cdot \Delta\text{ice} + 3^{\circ}\text{C} \cdot \frac{\ln[(p\text{CO}_2 - 278)/278]}{\ln(2)}$$



Includes the glacial/interglacial
pCO₂ feedback.



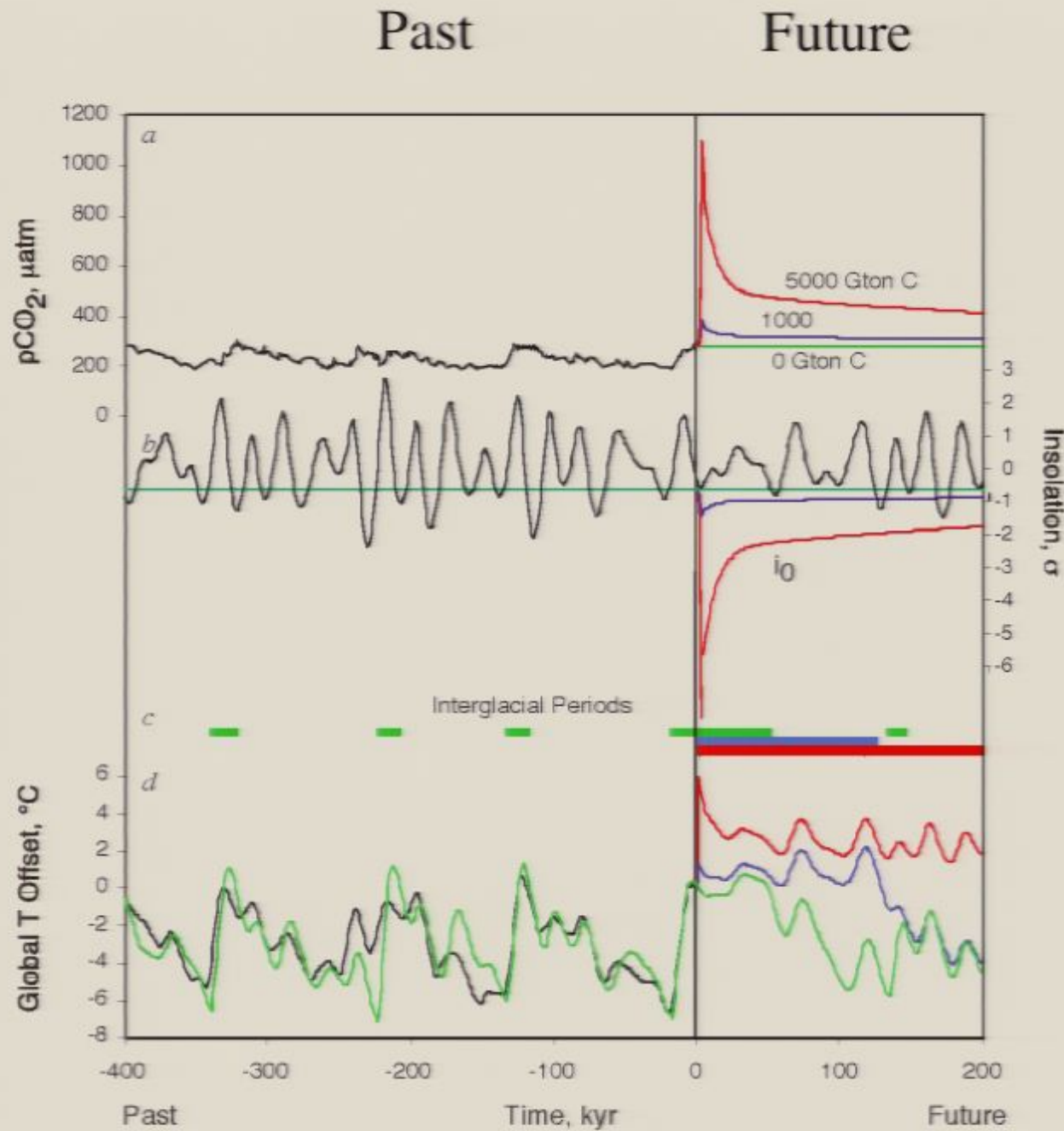
No Ice Age
for 500 Millennia

A Crude Temperature Estimate

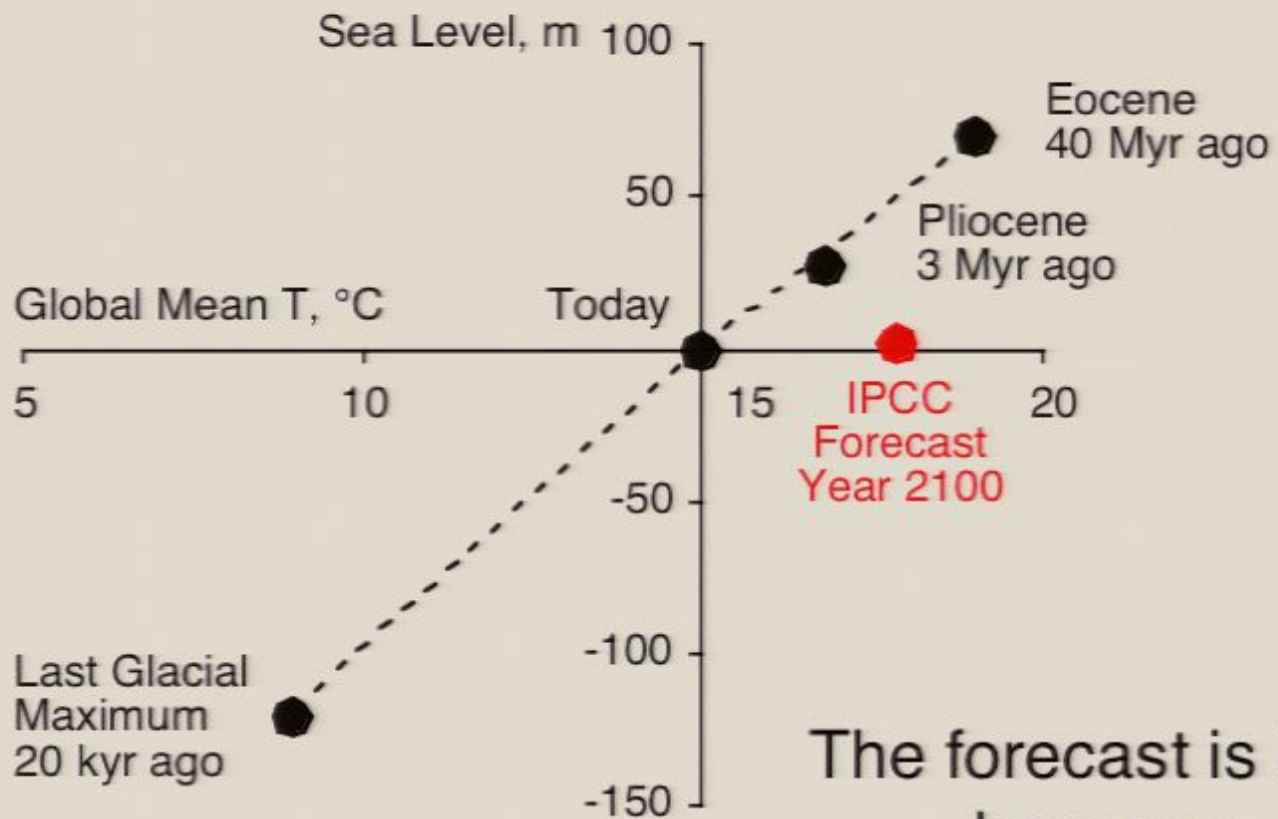
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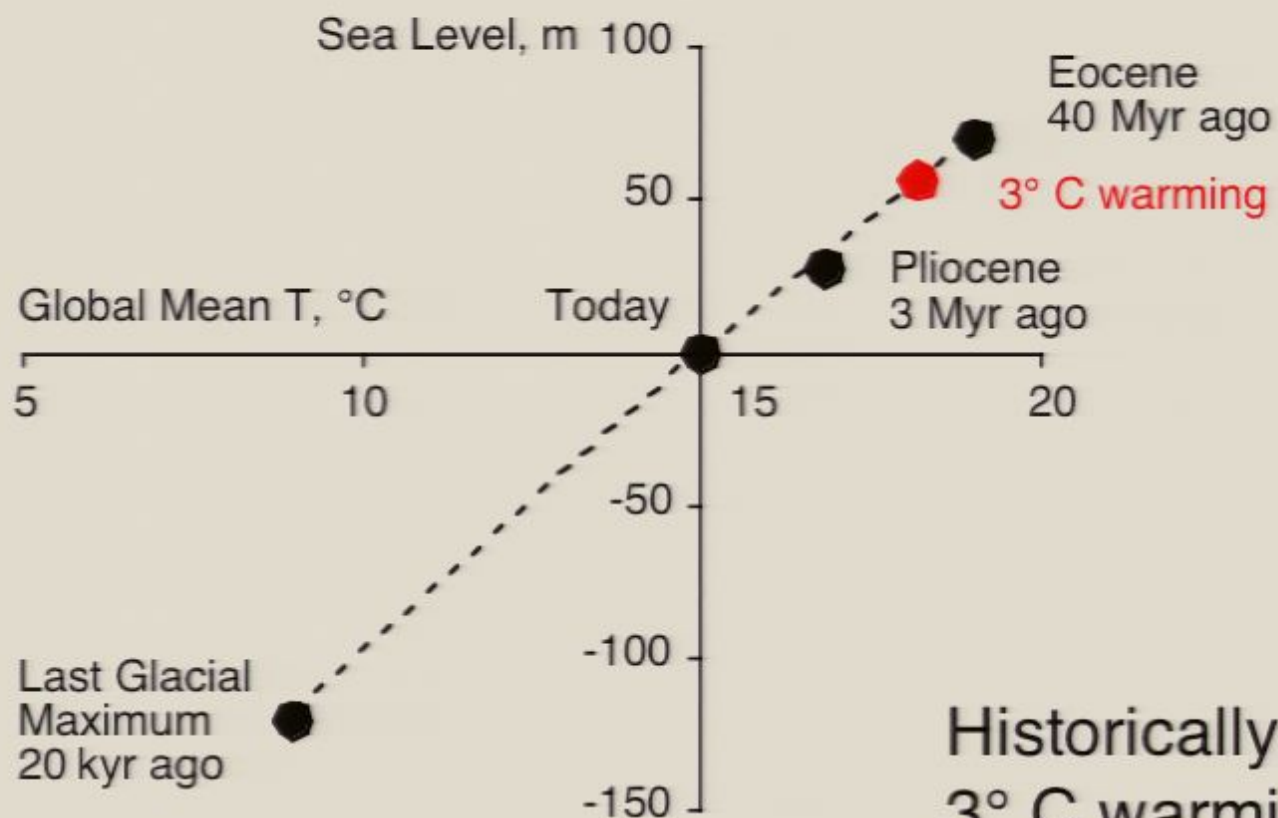


Sea Level



The forecast is low
because
melting ice is slow

Sea Level



Historically,
3° C warming
is a big deal.

Summary

A long tail (CO₂)

A movable trigger (nucleating an ice sheet)

- 12 [x] Orbital forcing and ice volume
- 13 [x] The spectrum of $d^{18}O$
- 14 [x] Precession and obliquity seem to work OK
- 15 [x] Eccentricity doesn't look so terrible, either. (On the face of it).
- 16 [x] Eccentricity
- 17 [x] Eccentricity needs a 20x amplifier
- 18 [x] The three cycles evolved in depend
- 19 [x] Calder's model
- 20 [x]
- 21 [x] Imbrie and Imbrie model
- 22 [x] Imbrie and Imbrie model
- 23 [x] Imbrie and Imbrie model
- 24 [x] Paillard threshold model
- 25 [x] Paillard model
- 26 [x]
- 27 [x] Gussy it up with ice growth/melt
- 28 [x] Paillard model gets the spectrum better
- 29 [x] Ice volume responds to N.H. insolation
- 30 [x] The World According to Paillard
- 31 [x] Paillard model
- 32 [x] CLIMBER Model Hysteresis
- 33 [x] Archer and Ganopolski, 2005
- 34 [x] Paillard model
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- 36 [x] Archer and Ganopolski, 2005
- 37 [x] Archer and Ganopolski, 2005
- 38 [x] A Crude Temperature Estimate
- 39 [x] Archer and Ganopolski, 2005
- 40 [x] Sea Level
- 41 [x] Sea Level
- 42 [x] Summary

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Summary

A long tail (CO_2)

A movable trigger (nucleating an ice sheet)

Click to add notes

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- 35 Yangtze Delta, China
- 36 Fossil Fuel Carbon and Ice Ages

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IPCC 2001 got this wrong

Table 1. Examples of greenhouse gases that are altered by human activities. (Based upon Chapter 1 and Table 1.1)

	CO ₂ (Carbon Dioxide)	CH ₄ (Methane)	N ₂ O (Nitrous Oxide)	CFC-11 (Chlorofluoro-carbon-11)	HFC-23 (Hydrofluoro-carbon-23)	CF ₄ (Perfluoro-methane)
Pre-industrial concentration	about 280 ppm	about 700 ppb	about 270 ppb	zero	zero	40 ppt
Concentration in 1998	345 ppm	1745 ppb	314 ppb	268 ppt	14 ppt	80 ppt
Rate of concentration change ^a	1.5 ppm/yr ^b	7.0 ppb/yr ^b	0.8 ppb/yr	-1.4 ppt/yr	0.55 ppt/yr	1 ppt/yr
Atmospheric lifetime	5 to 200 yr ^c	12 yr ^d	114 yr ^d	45 yr	260 yr	>50,000 yr

^a Rate has fluctuated between 0.7 ppm/yr and 2.8 ppm/yr for CO₂ and between 0 and 13 ppb/yr for CH₄ over the period 1982 to 1998.

^b Rate is calculated over the period 1982 to 1998.

^c No single lifetime can be defined for CO₂ because of the different rates of uptake by different removal processes.

^d This lifetime has been defined as an "equivalent time" that takes into account the indirect effect of the gas on its own residence time.

Since then, it's repeated everywhere

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IPCC 2001 got this wrong

Table 1. Examples of greenhouse-gases that are altered by human activities. (Based upon Chapter 1 and Table 4-1)

	CO ₂ (Carbon Dioxide)	CH ₄ (Methane)	N ₂ O (Nitrous Oxide)	CFC-11 (Chlorofluoro- carbon-11)	HFC-23 (Hydrofluoro- carbon-23)	CF ₄ (Perfluoro- methane)
Pre-industrial concentration	about 280 ppb	about 700 ppb	about 270 ppb	zero	zero	40 ppb
Concentration in 1998	365 ppb	1745 ppb	314 ppb	288 ppb	14 ppb	80 ppb
Rate of concentration change ^a	1.5 ppm/yr ^b	7.0 ppb/yr ^b	0.8 ppb/yr	-1.4 ppb/yr	0.55 ppb/yr	1.0 ppb/yr
Atmospheric lifetime	5 to 200 yr ^c	12 yr ^d	118 yr ^d	45 yr	280 yr	>50,000 yr

^a Rate has fluctuated between 0.7 ppm/yr (1958 to 1960) and 1.8 ppm/yr (1992 and 1993) for CO₂ and between 0 and 13 ppb/yr for CH₄ over the period 1982 to 1998.
^b Rate is calculated over the period 1980 to 1998.
^c The large lifetime can be defined for CO₂ because of the different rates of uptake by different natural processes.
^d The lifetime has been defined as an "equivalent time" that takes into account the indirect effect of the gas on its own residence time.

Slide 1: Global Warming in Geolo...

Since then, it's repeated everywhere

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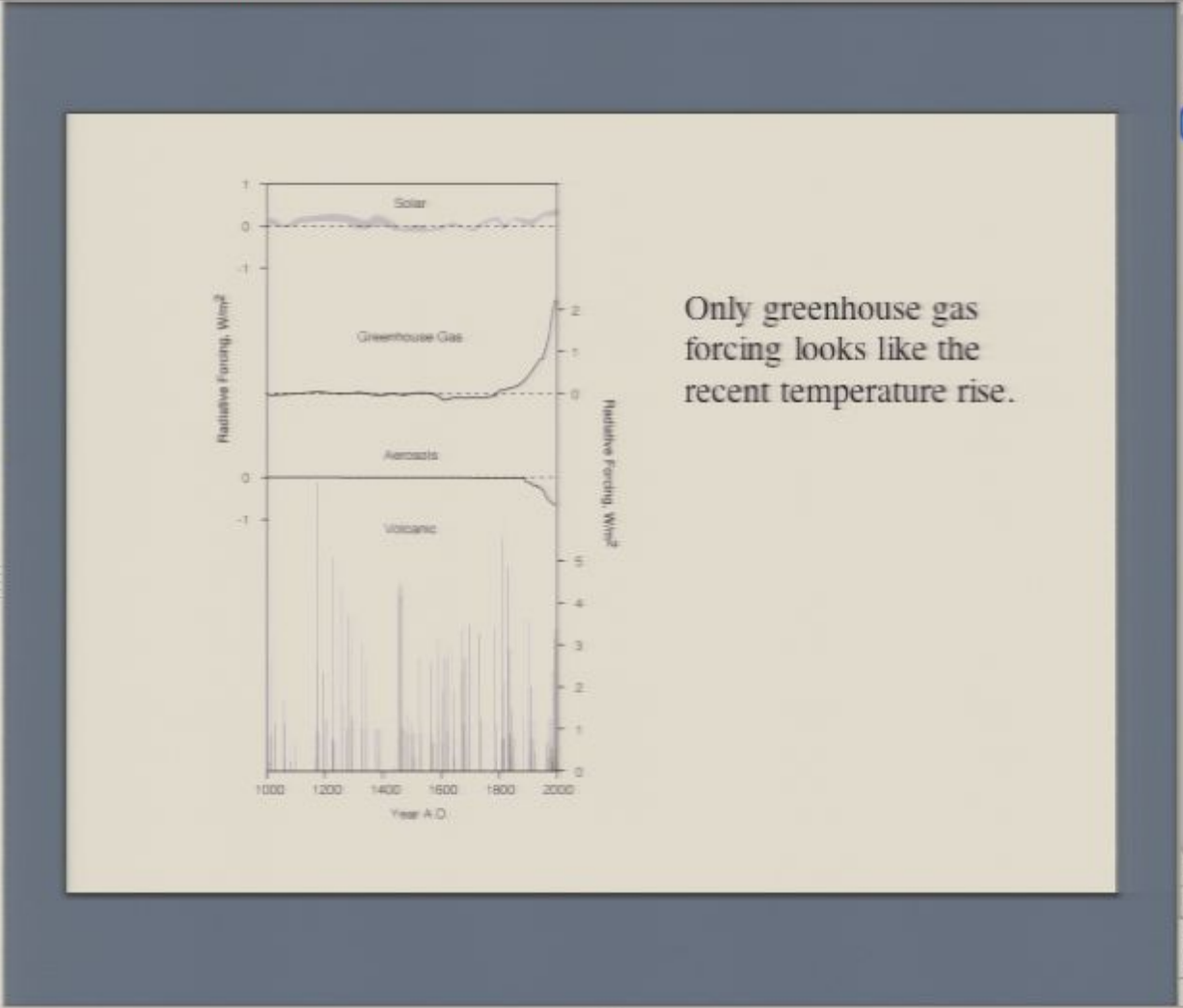


Ice sheets and glaciers are melting around the world

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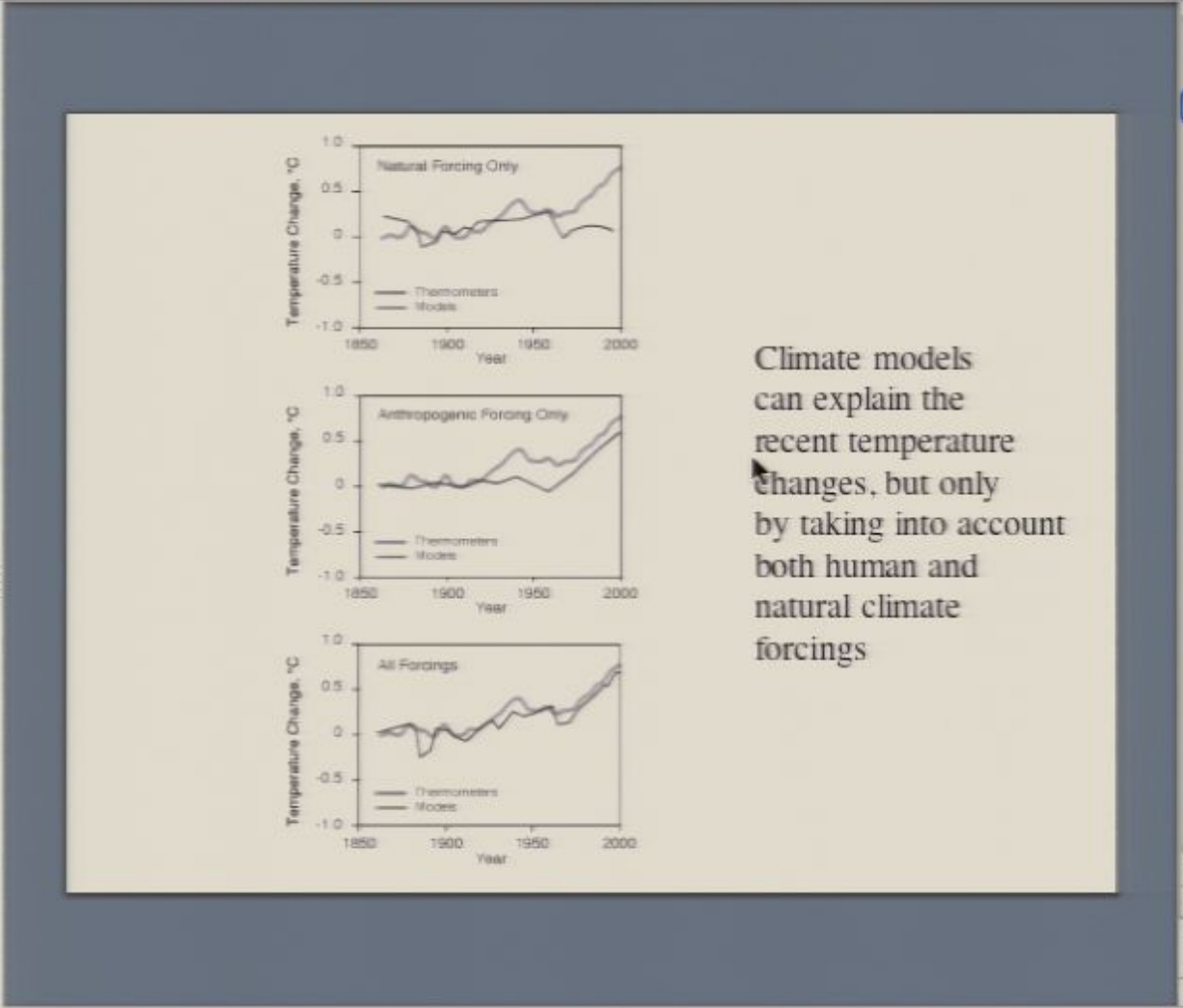


Only greenhouse gas forcing looks like the recent temperature rise.

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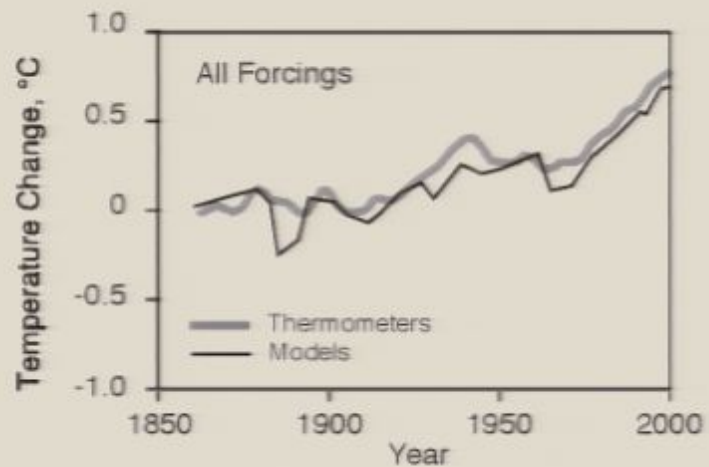
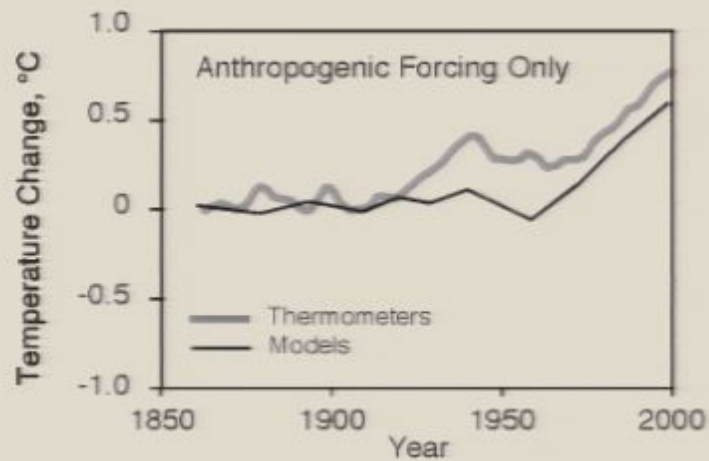
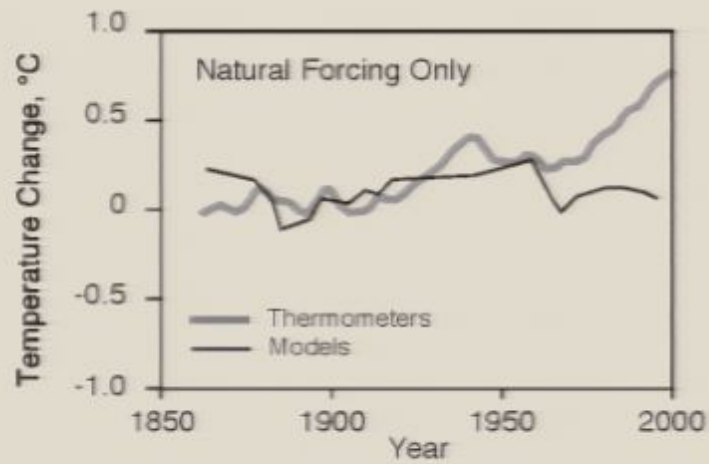
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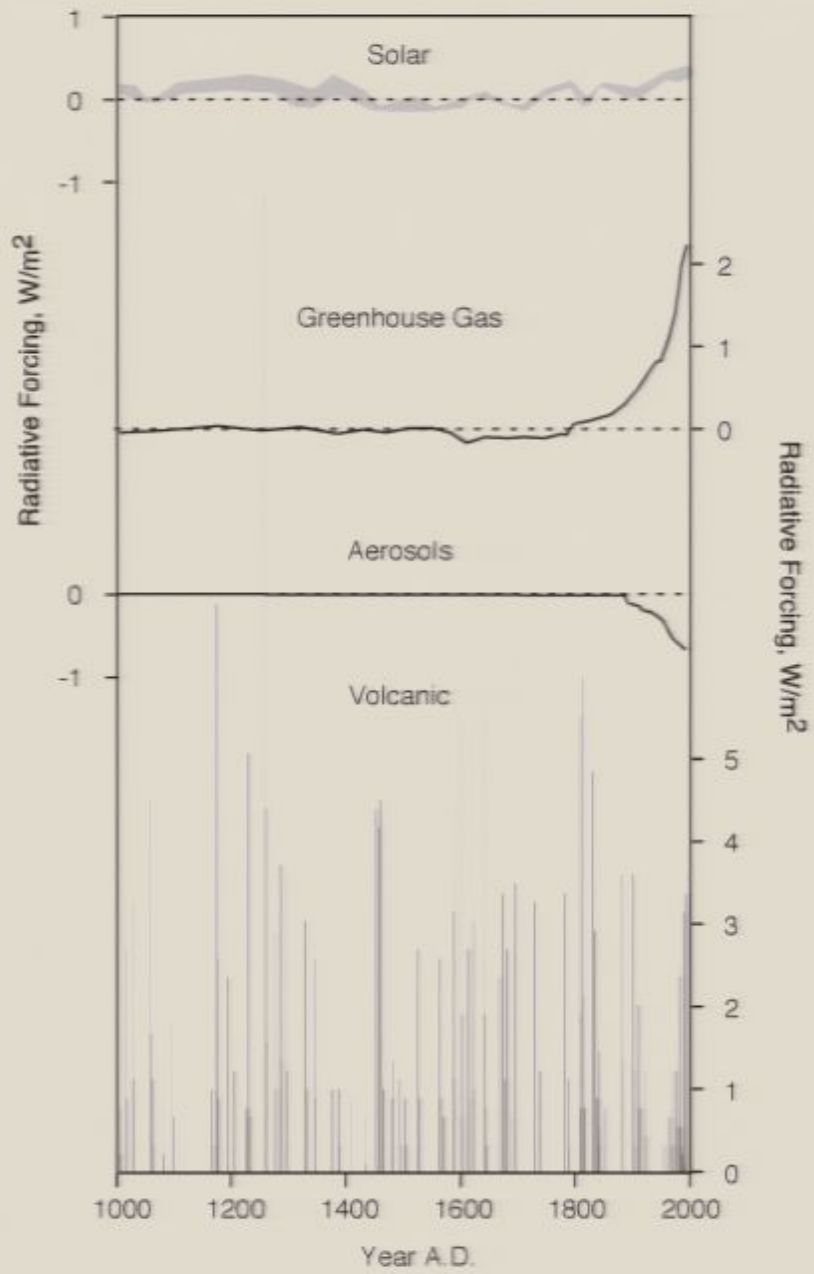


Climate models can explain the recent temperature changes, but only by taking into account both human and natural climate forcings

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